A Comparison of NGA-West2 Ground-Motion Models to Recent Chinese Data

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Abstract

Sichuan Province is one of the most seismically active areas in China. Several major earthquakes have struck this region in recent years, including the 2008 M_w 7.9 Wenchuan earthquake, the 2013 M_w 6.6 Lushan earthquake, and the 2017 M_w 6.5 Jiuzhaigou earthquake, causing vast socio-economic loss. These events have considerably expanded the ground-motion database for China, with high-quality records that can be used to improve ground-motion models (GMMs) for China and worldwide, and for advancing seismic hazard assessment. This study investigates the compatibility of advanced GMMs established by the NGA-West2 (Enhancement of Next Generation Attenuation Relationships for Western US) project for application to shallow crustal earthquakes in tectonically active regions and a local GMM for China with the recent Chinese strong-motion data. Specifically, this study quantitatively investigates the compatibility of these models with respect to magnitude scaling, source-to-site distance scaling, and site-effect scaling implied by the considered Chinese data. The results show that the Chinese data is characterized by slower attenuation with respect to the NGA-West2 GMMs. Regression coefficients are re-estimated for the considered GMMs based on the Sichuan data. These updated international and local GMMs can be applied for seismic hazard analysis in Sichuan. The comparisons between the extended ground-motion database for the Sichuan Province and the updated NGA-West2 GMMs represent an important first step toward the future development of new, improved GMMs for Sichuan and, ultimately, for China.

Keywords: Ground-motion models, China, NGA-West2, Intensity measures

1. Introduction

Sichuan Province is one of the most seismically active areas in China, with a long history of devastating earthquakes caused by the complex geographic and tectonic setting in the region. Specifically, Sichuan is located in southwest China, occupying most of the Sichuan Basin and the easternmost part of the Tibetan Plateau, between the Jinsha River on the west, the Daba Mountains in the north, and the Yungui Plateau to the south. Due to the collision between the Indian Plate and the Eurasian Plate, the Longmenshan fault was developed throughout the Cenozoic era [1] and caused both the 2008 M_w 7.9 Wenchuan earthquake [2] and the 2013 M_w 6.6 Lushan earthquake [3]. Recently, on 8 August 2017, the

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 M_w 6.5 Jiuzhaigou earthquake struck Sichuan, which was caused by the Kunlun fault in the region [4].

These moderate-to-large earthquake events have caused huge socio-economic loss. It is reported that the Wenchuan earthquake resulted in 88,000 fatalities or missing people, nearly 400,000 injuries, and five million people homeless [5]. Several buildings and infrastructure were damaged and severe cascading hazards (e.g., landslides and debris flows) were triggered by the Wenchuan earthquake [6]. The direct economic loss of the Wenchuan earthquake was estimated at US\$ 123 billion [7]. The Lushan earthquake, although characterized by a magnitude smaller than the Wenchuan earthquake, resulted in 193 fatalities, 25 missing people and over 12,000 injuries, causing extensive large-scale landslides and structural damage [3]. The economic loss of the Lushan earthquake was approximately US\$ 48.4 million [8]. The Jiuzhaigou earthquake, with a similar magnitude to the Lushan earthquake, caused 25 fatalities and 525 injuries due to the seismically induced geotechnical hazards [9]. The economic loss of the Jiuzhaigou earthquake was over RMB 8 billion (i.e., approximately US\$ 1.2 million) [10]. The earthquake-triggered landslides severely affected the Jiuzhaigou Valley, a World Natural Heritage Site selected by the United Nations Educational, Scientific, and Cultural Organization (UNESCO) [11]. However, the Jiuzhaigou earthquake did not cause substantial structural damage [9].

These destructive events have been well-recorded by the China National Strong Motion Observation Network System (NSMONS) managed by the China Earthquake Administration (CEA), which has been in operation since March 2008 [12]. The quantity and quality of these records have significantly improved the strong-motion database for China, promoting research in earthquake engineering and engineering seismology, in particular, for the characterizations of ground-motion data of moderate-to-large earthquakes and the development of new ground-motion models (GMMs) for probabilistic seismic hazard analysis (PSHA) in China.

This study focuses on the three recent events in Sichuan introduced above and investigates the compatibility of the NGA-West2 (Enhancement of Next Generation Attenuation Relationships for Western US) GMMs [13–17, hereafter referred to as ASK14, BSSA14, CB14, CY14 and I14] and a local Chinaspecific GMM [18, hereafter Huo89] to this Chinese strong-motion data. The model compatibility is investigated in terms of magnitude scaling, source-to-site distance scaling, and site-effect scaling. Moreover, the coefficients corresponding to the constant term, the geometric spreading term and the linear site response term in the considered GMMs are re-estimated based on the Sichuan data. The comparisons between the NGA-West2 GMMs and the extended ground-motion database for the Sichuan Province represent an important first step towards the future development of new, improved GMMs for Sichuan and, ultimately, for China.

This paper is organized as follows. We first review several studies comparing various GMMs to Chinese ground-motion data, followed by an introduction of the dataset used in this study, highlighting some issues in the site characterization. Then, we statistically evaluate the different horizontal groundshaking definitions, followed by a preliminary evaluation of the polarization in the dataset. A series of qualitative and quantitative analyses of the considered GMMs with respect to the Sichuan data in terms of magnitude scaling, source-to-site distance scaling, and site-effect scaling are carried out. Finally, some coefficients in the considered GMMs are re-estimated to address the specific features observed in the Sichuan data.

2. Literature Review

Engineering analysis of local and worldwide recent strong-motion data with respect to state-of-theart GMMs is beneficial for several reasons. Firstly, this allows one to evaluate whether a previously developed GMM is consistent with the newly available data, in order to improve the model performance. Moreover, one can assess the features of the new data and possibly implement modifications in GMMs if the new data is included. In addition, analyzing the compatibility between GMMs and ground-motion data can help to increase the confidence in model applications (e.g., seismic risk assessment exercises and the definition of target spectra for performance-based earthquake engineering or PBEE).

A set of ad-hoc GMMs for worldwide shallow crustal regions developed by the Pacific Earthquake Engineering Research Center (PEER) are usually used as state-of-the-art and state-of-practice benchmark GMMs for such analysis. Those models include the NGA-West (Next Generation of Ground-Motion Attenuation Models for the Western US) GMMs [19–23, hereafter referred to as AS08, BA08, CB08, CY08 and I08] developed based on a global strong-motion database consisting of 3,551 multi-component records from 173 shallow crustal earthquakes between 1935 and 2003, ranging in magnitude from 4.2 to 7.9 [24], and the updated NGA-West2 GMMs, developed based on a larger repository consisting of 21,336 (over 98%) three-component records from 599 events till 2011, ranging in magnitude from 3.0 to 7.9 [25]. Because of the large quantity and high quality of the NGA-West2 database, certain effects (e.g., the nonlinear site response) are well resolved in some of the NGA-West2 GMMs that could not be evaluated using only Chinese data. A general review of GMMs published prior to 2016 is given by Douglas and Edwards [26].

For this type of comparisons, the predictive variables required by the considered GMMs need to be estimated, including but not limited to, magnitude in terms of moment magnitude (M_w) and surface magnitude (M_s) ; source-to-site distance in km in terms of rupture distance (R_{rup}) , the closest distance to the surface projection of the fault rupture or Joyner-Boore distance (R_{JB}) , epicentral distance (R_{epi}) , and the horizontal distance from a site to the top edge of the rupture measured perpendicular to the strike of the fault (R_X) ; style-of-faulting; the average shear-wave velocity in the upper 30 m soil (V_{S30}) in m/s; the depth to the top of rupture (Z_{tor}) in km; the depth to $V_{S30} = 1.0$ km/s $(Z_{1.0})$ in km; the depth to $V_{S30} =$ 2.5 km/s $(Z_{2.5})$ in km; the fault dip in degrees; the rupture width in km; region-specific parameter for anelastic attenuation. Hence, these models are considerably more complicated than previous ones, often requiring several additional input parameters that are unknown. For instance, Kaklamanos et al. [27] introduced a framework for estimating the unknown source, path, and site parameters when implementing the NGA models in the engineering practice.

Several studies [e.g., 9, 12, 28–36] have investigated the compatibility of worldwide and Chinese GMMs to Chinese strong-motion data, in particular, to the aforementioned three major events in Sichuan. The main findings of these studies are summarized in Appendix and some common findings are as follows:

• In the case of the Wenchuan earthquake, the observed peak ground acceleration (PGA) and short-

period spectral ordinates are generally higher than the predictions of the reference GMMs over the considered distance range while long-period spectral ordinates are lower than those predicted by the reference GMMs over the distance range;

- In the case of the Lushan earthquake and the Jiuzhaigou earthquake, the observed PGA and shortperiod spectral ordinates are generally consistent with the predictions of the reference GMMs over the considered distance range while long-period spectral ordinates are lower than those predicted by the reference GMMs over the distance range;
- Directivity effect (i.e., the focusing of seismic wave energy along the fault in the direction of rupture) and hanging-wall effect (i.e., the increase in ground motion at short distances for sites on the hanging wall side of a rupture when compared to sites on the footwall side at the same closest distance) are detected in the Wenchuan earthquake and the Lushan earthquake;
- China ground motions attenuate more slowly with respect to source-to-site distance than the reference GMMs, possibly due to the higher values of the shear-wave quality factor Q for China than those implied by the reference GMM.

Although most of the considered studies focus on a single earthquake, Ji et al. [35] and Dangkua et al. [36] investigate the model compatibility using pooled Chinese data from multiple earthquakes. Their results show that the Chinese GMMs of Huo89 and Yu et al. [37] and NGA-West2 GMMs are generally consistent with the observations of PGA and short-period spectral ordinates over the considered distance range while these GMMs overpredict the long-period spectral ordinates over the considered distance range. It is worth pointing out that, although there are more recent China-specific GMMs [e.g., 37] compared to Huo89, they are only applicable to rock sites while Huo89 is applicable to both rock and soil sites. Because site effect has an important impact on ground motions, this study has chosen the Huo89 model, a widely used GMM for seismic-hazard assessment in China [35], for comparison. Moreover, Dangkua et al. [36] found that Chinese ground motions attenuate more slowly with respect to sourceto-site distance than the NGA-West2 GMMs. Furthermore, Dangkua et al. [36] observed some residual biases in the NGA-West2 and Chinese GMMs corresponding to the source-to-site distance scaling and magnitude scaling with respect to Chinese data, which implies that these GMMs may need adjustments to account for the specific features observed in the pooled Chinese data. However, it is noteworthy that the multi-event database used in these studies implicitly assumed homogeneous geological conditionals across a fairly large geographic area, which requires further investigation.

It is also shown in Appendix that a variety of data and methods have been used in the considered studies. For instance, most of the studies deal with data from NSMONS while Dangkua et al. [36] used a mixed dataset of NSMONS and NGA-West2. The processing methods applied to the ground-motion data in Wang et al. [29] and Bai [34] follow the NGA-West2 procedure, while Ren et al. [9], Wen and Ren [32] and Ji et al. [35] select different processing procedures; in both cases, the processing approaches may be different from the one used by NSMONS. Moreover, various definitions of the horizontal ground-shaking intensity measures (IMs) are used in these studies. For instance, the geometric mean definition

is used by most Chinese GMMs; the GMRotI50 definition (i.e., the median orientation-independent geometric mean, using period-independent rotation angle) proposed by Boore [38] is used in Bai [34]; and the RotD50 definition (i.e., the median orientation-independent non-geometric-mean measure, using period-dependent rotation angle) proposed by Boore [39] is used in Ren et al. [9]. Furthermore, the site classification used in each study consists of the two-category site classification provided by NSMONS (i.e., rock or soil), the five-category site classification (shown in Table 1) specified in the China Code for Seismic Design of Buildings (GB50011-2010) [40] based on V_{S20} (i.e., the average shear-wave velocity in the upper 20 m soil, in m/s) and soil depth, and the six-category site classification (A to F) defined in the National Earthquake Hazards Reduction Program (NEHRP) [41] based on V_{S30} . Since V_{S20} is the considered site-classification parameter in GB50011-2010, the V_{S30} value required by the international GMMs is estimated or assigned based on various standards or engineering judgements, resulting in different V_{S30} values for same sites used in different studies. In addition, several studies fit trend lines to residuals (e.g., total residuals, inter-event residuals or intra-event residuals) in order to analyze whether there is potential bias in GMMs. However, the method used to compute the inter-event residuals varies, such that, Dangkua et al. [36] follow the equations in Abrahamson and Youngs [42] while Wang et al. [29] use the average of total residual as the inter-event residuals for the Wenchuan earthquake.

[Table 1 about here.]

The above discussions suggest a general lack of consistency in comparing the Chinese data and the state-of-the-art GMMs across different studies. This may lead to model/data misinterpretation, biased results, and unsuitable GMMs used in performing PSHA for Sichuan Province. Therefore, this study presents a systematic way to compare GMMs to the ground-motion data of three most relevant recent Chinese earthquakes in Sichuan. A similar method has been used by Scasserra et al. [43] for Italian data. In addition, this study statistically investigates the different definitions of horizontal ground-shaking IMs. Moreover, this study assesses the improvement on model performance before and after accounting for the features observed in the Sichuan data, as further discussed in the following sections.

3. Dataset

The strong-motion data for the considered Sichuan earthquakes are collected from NSMONS and are processed according to standard practice by researchers at CEA (e.g., applying an acausal Butterworth filter with corner frequencies of 0.1 Hz and 30 Hz and baseline correction - personal communication). It is worth noting that the CEA processing procedure may be different from that used by the NGA-West2 regarding the selection of record and component-specific corner frequencies. In particular, the two-component horizontal records with $R_{JB} \leq 250$ km are selected in this study. The 250 km cut-off distance is within the applicable distance range of most NGA-West2 GMMs while the distance range of I14 GMM is extrapolated to 250 km. The geographic distribution of the recording stations and the M_w - R_{JB} distribution with the NSMONS site classification are presented in Figure 1. The metadata for each considered event is listed in Table 2.

[Figure 1 about here.] [Table 2 about here.]

The considered GMMs in this paper are the five NGA-West2 GMMs and the Huo89 GMM. The magnitude metric M_w is used by the NGA-West2 GMMs while M_s is used by Huo89. Given the finite-fault models in Table 2, several predictor variables required by these GMMs can be calculated accordingly. The region-specific parameters, e.g., region predictor in ASK14 and BSSA14, the coefficient of regional differences Δc_{20} in CB14, and the multiplicative adjustment factor γ in CY14 for anelastic attenuation are chosen for China accordingly.

3.1. Site classifications and V_{S30} of stations

The knowledge of V_{S30} is generally required by advanced GMMs (e.g., NGA-West2 GMMs). The V_{S30} accounts for the local site conditions and its accuracy has a huge impact on the GMM fitting and its applications. In fact, the local site conditions can modify the amplitude, frequency, and duration of the seismic shaking. As discussed in Section 2 and Appendix, various measurements/estimations of V_{S30} are used in the literature. Though similar results are observed in these studies, different V_{S30} values of the recording sites may shift the median predictions used as a reference in each comparison. Thus, this type of visual comparisons can only provide a general guidance on the overall goodness-of-fit while detailed analysis of residuals, as performed in this study, could provide more insights into the model performance.

To further elaborate on this issue, the V_{S30} estimated from two different methods is examined. Yu [44] obtained the V_{S20} soil profiles from in-situ borehole tests for 147 stations in Sichuan Province and Gansu Province. Assuming the soil medium is unchanged from the bottom of the borehole till 30 m depth, Yu extrapolated the V_{S20} values to estimate V_{S30} for all stations; 80 stations out of the 147 above are used in this study (V_{S30} values at these 80 stations are denoted by $V_{S30,Yu}$). Alternatively, the USGS V_{S30} Models and Data provides V_{S30} values ($V_{S30,USGS}$ hereafter) derived through approximated relationships with the topographic slope [45]. The $V_{S30,USGS}$ is, on average, higher than $V_{S30,Yu}$, especially in the range $V_{S30,USGS} \ge 500$ m/s (shaded area). According to GB50011-2010, a soil layer with a shear-wave velocity ≥ 500 m/s is regarded as bedrock. Yu stated that when the drilling depth (from ground surface to bedrock) is significantly ≤ 20 m, there is greater uncertainty in the extrapolated value.

[Figure 2 about here.]

For the purpose of this study, the $V_{S30,Yu}$ value at each site/station is used as the preferred value in the analysis and, if this data is not available, the $V_{S30,USGS}$ is used instead. It is worth noting that different V_{S30} estimation procedures were used in NGA-West2 for Chinese stations. Specifically, in the case of stations for which the V_S profiles are not available, values have been inferred V_{S30} from multiple proxies, such as the ground slope (as done here), a geotechnical proxy, and a terrain proxy [25]. However, the various estimation procedures result in a generally consistent site-class classification. However, as shown above, available information at seismic stations are frequently inadequate for a proper characterization of local site conditions and this can have a significant impact on the seismic hazard estimates at both large and local scale. Thus, there is an urgent need to invest in site characterization of the recording stations in China in order to improve the current knowledge on strong-motion seismology.

4. Horizontal ground-shaking definitions

Different horizontal ground-shaking definitions are used in various GMMs, as mentioned above. Several studies [e.g., 39, 46, 47] showed that the differences among these definitions from global databases are minimal (negligible at low periods and generally less than 5-8%). Moreover, Shahi and Baker [48] suggest the ratios of maximum-to-median IM across orientations (i.e., RotD100/RotD50, where RotD100 is the maximum orientation-independent non-geometric-mean measure, using period-dependent rotation angle proposed by Boore [39]) can be used as a proxy for the polarization of ground motions, which is important for analyzing the structural responses of 3-Dimensional (3-D) structural models in all orientations. Thus, this section first analyzes whether the differences between RotD50 and geometric mean definitions are statistically significant in the Sichuan data and then investigates the polarization in the considered dataset.

The median ratio between RotD50 and the geometric mean definition in terms of spectral acceleration (S_a) , $S_{a,\text{RotD50}}/S_{a,\text{GM}}$, computed from all the considered earthquakes are presented in Figure 3 together with its standard deviation. It is worth noting that the results are computed as the exponential of the mean and standard deviation of the natural logarithm of $S_{a,\text{RotD50}}/S_{a,\text{GM}}$ as a function period T. These estimates are useful for the conversion from a GMM for one definition to another, accounting for the propagation of uncertainty [46]. The median $S_{a,\text{RotD50}}/S_{a,\text{GM}}$ in Figure 3 (a) is slightly larger than unity (≤ 1.05), which suggests the two horizontal ground-shaking definitions give similar S_a amplitude. Moreover, the median $S_{a,\text{RotD50}}/S_{a,\text{GM}}$ she structural period T increases, reaching about 1.05 at T = 5 s. Beyer [46] argued that these increments may be due to "the stronger polarization of ground-motion waves". The standard deviation of $S_{a,\text{RotD50}}/S_{a,\text{GM}}$ in Figure 3 (b) increases with T from about 1.06 to 1.1 at T = 5 s, which is consistent with literature [e.g., 39]. However, these small values may not significantly increase the total standard deviations when converting GMM for RotD50 from geometric mean.

[Figure 3 about here.]

Besides the amplitude of different horizontal ground-shaking definitions, the distributions of different definitions are also critical for PSHA applications and the GMM conversion between different definitions. Jayaram and Baker [49] showed $S_{a,GMRotI50}$ follows a multivariate lognormal distribution. The same assumption is applied explicitly or implicitly when developing GMMs for other definitions. This section visually evaluates the IM distributions of RotD50 and geometric mean definitions by a quantile-quantile (Q-Q) plot, which is a graphical method used to determine whether the two samples come from the same distribution or not. If two samples come from a population with the same distribution, the data points should fall approximately along a 45-degree reference line. The Q-Q plot of $\ln(IM_{RotD50})$ versus $\ln(IM_{GM})$ in Figure 4 shows an approximately 45-degree straight line, suggesting that the difference in IM distributions of two definitions are not statistically significant. Thus, the differences between RotD50

and geometric mean definitions in Sichuan data regarding amplitude (i.e, point estimate) and distribution are insignificant. Therefore, to be consistent with NGA-West2 GMMs, RotD50 definition is used in the following sections for model comparisons and adjustments.

[Figure 4 about here.]

Regarding the polarization in ground motions, Shahi and Baker [48] suggest that, if RotD100/RotD50 $\approx \sqrt{2}$, the structural response is polarized in one orientation while if RotD100/RotD50 ≈ 1 , the structural response is similar in all orientations. As an example, Figure 5 presents the displacement trace and S_a over all orientations for representative stations/records. As shown in Figure 5 (a), the response at station 51GYZ in the Wenchuan earthquake shows little polarization with $S_{a,\text{RotD100}}/S_{a,\text{RotD50}} \approx 1.08$ while 51MZQ station in Figure 5 (b) shows a relatively strong polarization with $S_{a,\text{RotD100}}/S_{a,\text{RotD50}} \approx 1.33$. As shown in Figure 5 (c) and (d), the structural responses at station 51PJD in the Lushan earthquake and at station 62WYX in the Jiuzhaigou earthquake show some polarization with $S_{a,\text{RotD100}}/S_{a,\text{RotD50}} \approx 1.22$. Therefore, there is a certain degree of polarization in the Sichuan data and these data should be used with caution for the analysis of 3-D structural models. Moreover, the $S_{a,\text{RotD100}}/S_{a,\text{RotD50}}$ ratio is compared with the empirical model of Shahi and Baker [48] at representative periods, as shown in Figure 6, which shows that the ratios observed in the Sichuan data are consistent with the international model.

[Figure 5 about here.] [Figure 6 about here.]

5. Comparisons of the median predictions from GMMs

In this section, the Sichuan data are first visually compared with the median predictions of PGA, and S_a at two representative periods (i.e., 0.1 s and 1.0 s) from the considered GMMs, as shown in Figure 7. For illustrative purpose, V_{S30} is set to 450 m/s (i.e., stiff soil and Class C in NEHRP) because it is close to the median V_{S30} value 375 m/s of the Sichuan stations and I14 is only applicable for $V_{S30} \ge 450$ m/s. The distance metric for ASK14, CB14, CY14, and I14 is R_{rup} , R_{JB} for BSSA14 and R_{epi} for Huo89. No hanging wall effect is assumed and the region-specific parameters are chosen for China. The basin depth terms, $Z_{1.0}$ and $Z_{2.5}$, are empirically estimated from V_{S30} based on several NGA-West2 studies, as follows,

• Chiou and Youngs [16] model for $Z_{1.0}$ used in ASK14 and CY14:

$$Z_{1.0} = \frac{1}{1000} \exp\left[-\frac{7.67}{4} \ln\left(\frac{V_{S30}^4 + 610^4}{1360^4 + 610^4}\right)\right];\tag{1}$$

• Boore et al. [14] model for $Z_{1.0}$ used in BSSA14:

$$Z_{1.0} = \frac{1}{1000} \exp\left[-\frac{7.15}{4} \ln\left(\frac{V_{S30}^4 + 570.94^4}{1360^4 + 570.94^4}\right)\right];$$
(2)

• Campbell and Bozorgnia [15] model for $Z_{2.5}$ used in CB14:

$$Z_{2.5} = \exp(7.089 - 1.144 \ln V_{S30}). \tag{3}$$

It is well-recognized that these empirical estimations for the basin depths dependent on V_{S30} are developed based on the global dataset (mostly California data and excluding Japan data) and may not apply to Sichuan, China. Scasserra et al. [43] suggest that "if the average basin effect implied by the NGA GMMs is significantly in error, it would be expected to produce bias at long periods", which is discussed in the subsequent section. More in general, these visual comparisons provide a general guidance of the goodness-of-fit while detailed analysis of residuals will be carried out in the following sections.

As shown in Figure 7 (a), (d), and (g), regarding the Wenchuan earthquake, there is an overall good agreement between the predictions of the NGA-West2 GMMs and the observations in terms of PGA and short-period spectral ordinates over the considered distance range. However, some GMMs overestimate the long-period spectral ordinates, particularly, in the far field. Though the data in the near field is sparse, the considered GMMs are generally consistent with the data. These results are expected since this event is included in the NGA-West2 database for the calibration of the NGA-West2 GMMs. The Huo89 gives a similar median prediction as that of the NGA-West2 GMMs for short-period spectral ordinates but Huo89 overestimates the long-period spectral ordinates. Regarding the Lushan earthquake in Figure 7 (b), (e), and (h), the NGA-West2 GMMs are consistent with PGA and short-period spectral ordinates data in the near field but overestimate them in the far field, and the NGA-West2 GMMs slightly overestimate the long-period spectral ordinates over the distance range. The Huo89 model provides consistent PGA predictions with the observations while it overestimates spectral ordinates over the considered distance range for the Lushan earthquake. Finally, regarding the Jiuzhaigou earthquake in Figure 7 (c), (f), and (i), the NGA-West2 and Huo89 GMMs significantly overestimate the PGA and spectral ordinates over the considered distance range. These findings are consistent with the literature in Appendix. In addition, Figure 7 seems to suggest that the magnitude scaling of the considered GMMs is well-modeled for the three considered events. However, the visual comparison of Figure 7 shows the NGA-West2 and Huo89 GMMs may require some modifications to be fully suitable for the Sichuan region. Either overestimation or underestimation of the actual ground motions may result in biased PSHA and seismic risk assessment, and consequently, inaccurate estimates of potential earthquake-induced loss (e.g., for insurance purposes, risk management, etc.).

6. Overall GMM bias and standard deviation relative to the Sichuan data

The previous section has introduced qualitative, visual inspection-based comparisons between the median predictions from the considered GMMs and the recent Sichuan data. This section provides a quantitative analysis of the compatibility of NGA-West2 and Huo89 GMMs to this data. The residuals between the observed data and the median predictions from each considered GMM are evaluated at eight representative periods of 0 s (i.e., PGA), 0.1 s, 0.2 s, 0.3 s, 0.5 s, 1.0 s, 1.5 s, and 2.0 s, individually. The residuals between the data and a particular GMM referred to with index k are computed as in Eq. (4):

$$(r_{ij})_k = \ln(\mathrm{IM}_{ij})_{\mathrm{obs}} - \ln(\mathrm{IM}_{ij})_k, \qquad (4)$$

where $(r_{ij})_k$ is the total residual at site j for the event i given GMM k (in natural log units); $(IM_{ij})_{obs}$ is the observed IM at site j for the event i; $(IM_{ij})_k$ is the median prediction of IM at site j for the event

i given GMM k; $k=1, \cdots, 6$ corresponds to ASK14, BSSA14, CB14, CY14, I14, and Huo89, respectively. For simplicity, subscript k will be omitted in the rest of the paper. It is worth pointing out that whether the hanging-wall effect is included or not in the considered GMM is determined based on the location of stations with respect to the hanging wall.

The analysis of residuals with respect to magnitude scaling, source-to-site distance scaling and siteeffect scaling requires the knowledge of the inter- and intra-event residuals. A linear mixed-effect regression is performed to calculate these quantities as in Eq. (5):

$$r_{ij} = c + \eta_i + \varepsilon_{ij} \,, \tag{5}$$

where c is the constant coefficient representing a general offset (or bias) of the selected GMM k with respect to the Sichuan data; η_i is the inter-event residual of the event i representing the mean offset of the considered GMM to the data for event i (after adjusting for mean offset c, which is based on all events), which is assumed to follow a normal distribution with zero mean and τ standard deviation (in natural log units); ε_{ij} is the intra-event residual at site j for the event i, which is assumed to follow a normal distribution (in natural log units).

The average misfit between the considered GMMs and Sichuan data, denoted by the constant coefficient c, and its 95% confidence interval (CI) are shown in Figure 8 (a). It is shown that the predictions of the NGA-West2 and Huo89 GMMs are generally compatible with the observed PGA values and short-period spectral ordinates. However, there is a misfit between the predictions of these GMMs and the long-period spectral ordinates. The constant coefficient c is around zero for PGA and short-period spectral ordinates while negative for long-period spectral ordinates and this trend is similar among the six GMMs. In other words, the considered GMMs produce generally consistent predictions for Sichuan data in terms of PGA and short-period spectral ordinates while these GMMs overestimate the long-period spectral ordinates. This may be due to that the estimated basin depth based on the California data do not fully reflect the basin properties in Sichuan region, particularly at long periods, where the basin effects are most pronounced. The Huo89 model gives a larger c (in absolute value) with higher variability compared to the NGA-West2 GMMs, which may be due to the lack of information on various earthquake effects/regression variables in the Huo89 GMM.

[Figure 8 about here.]

Figure 8 (b), (c), and (d) present $c + \eta$ for three considered events, respectively. This quantity is equivalent to the average total residual of each event, which is defined as the inter-event residuals by Wang et al. [29], and it represents the misfit between the considered GMMs and a specific Sichuan event. Regarding the Wenchuan earthquake and the Lushan earthquake in Figure 8 (b) and (c), the six GMMs are generally consistent with the observations for PGA and short-period spectral ordinates while overestimating the long-period spectral ordinates. Regarding the Jiuzhaigou earthquake in Figure 8 (d), all six GMMs overestimate both the short- and long-period spectral ordinates across periods. The aforementioned results are consistent with the literature [e.g., 29].

The inter- and intra-event standard deviation, τ and ϕ , versus periods and their 95% CI are presented in Figure 8 (e) and (f), respectively. The estimated τ from the Sichuan dataset is around 0.5 across periods with large variability, which is possibly due to a limited number of events. The Huo89 gives a larger τ with higher variability than the NGA-West2 GMMs, which may be because the Huo89 GMM does not consider the style-of-faulting and other predictors, thus, has a higher variability across events. The trends of ϕ estimated from the Sichuan data against periods are consistent for all GMMs, however, ϕ is larger than expected, which may be because the information of complex earthquake effects is not available. In addition, the inter- and intra-event standard deviations predicted by the CB14 GMM given a reverse fault event in China with $M_W = 7.0$, $V_{S30} = 450$ m/s, $R_{rup} = 125$ km (i.e., the median source-to-site distance) are presented for comparison. It is shown that the results in this study follow a similar trend to the prediction; however, notable difference is observed, possibly due to the limited dataset used in this study. The aleatory variability of GMM is assumed to be dependent on magnitude, site condition (i.e., the effects of soil nonlinear response), and source-to-site distance (possibly due to "the epistemic uncertainty in regional attenuation rates" according to Abrahamson et al. [13]). Although different NGA-West2 GMMs have different assumptions and models for aleatory variability, they produce similar estimate of aleatory uncertainties and the finding mentioned above also holds for the comparisons with the predictions of ASK14, BSSA14, and CY14.

The inter-event residual $\eta = (\eta_i)$ is important to study the magnitude scaling. However, because this study only considers three events and two of them have a very similar magnitude, thus, the magnitude scaling cannot be properly captured by analyzing η . Nevertheless, the visual comparisons of Figure 7 seem to suggest the magnitude scaling is generally well-captured for the three considered events. Thus, this study will further focus on the compatibility of the considered GMMs regarding the source-to-site distance and the V_{S30} -based site effect implied by the recent Sichuan data.

7. Distance scaling

This section assesses the considered GMMs in characterizing the source-to-site distance scaling of the Sichuan data. This is achieved by examining trends of intra-event residuals as a function of the logarithm of source-to-site distance, similarly to Scasserra et al. [43]. To help illustrate trends, the linear regression as in Eq. (6) is used:

$$\varepsilon_{ij} = a_R + b_R \ln(R_{ij}) + \widetilde{\varepsilon}_{ij} \,, \tag{6}$$

where a_R and b_R are the regression coefficients for the distance scaling; R_{ij} is the source-to-site distance in km at sites j for the event i; $\tilde{\varepsilon}_{ij}$ is the remaining intra-event residual at site j for the event i that results from the fit of Eq. (6). Since I14 is applicable for "quasi-linear" sites with $V_{S30} \ge 450$ m/s, thus only 62 sites with $V_{S30} \ge 450$ m/s are used to assess the I14 GMM. The slope parameter b_R represents approximately the misfit of the distance scaling in the considered GMMs relative to the Sichuan dataset. The statistical *t*-test is used to study the significance level p of the distance dependency in the intra-event residuals. The null hypothesis to be tested is $H_0: b_R = 0$ (i.e., no bias in distance scaling). For example, if p < 0.05 (or 1 - p > 0.95), it is suggested there is significant evidence to reject the null hypothesis H_0 and the distance scaling of the selected GMM is biased with respect to Sichuan data.

The results of the analysis are plotted in Figure 9 together with a 95% CI of predictions. It is shown that there is bias in the distance scaling (i.e., 1 - p => 0.95) of CB14 model at long periods, and that of

Huo89 for PGA and long periods. The NGA-West2 and Huo89 GMMs with biased distance scaling give positive slope values, suggesting the decay rates with respect to distance in the Sichuan data are slower than those implied by the selected GMMs. Dangkua et al. [36] argued this may be due to the higher values of the shear-wave quality factor Q for China than those implied by the GMMs. It is worth noting that the absolute values of the slope are small in general, which may be due to the limited quantity of data. These results show that the distance scalings in some NGA-West2 GMMs and Huo89 GMM need to be adjusted to better capture the observed distance dependency in the Sichuan data.

[Figure 9 about here.]

The distance scaling in most of the available GMMs represents the attenuation of seismic energy with respect to source-to-site distance R, which is mainly characterized by the geometrical spreading term and the anelastic term. The geometrical spreading term models the decreasing energy with increasing distance and is commonly represented as $\ln \sqrt{R^2 + h^2}$, where h is the fictitious depth to be estimated. The anelastic term accounts for the material anelasticity that progressively attenuate the wave energy denoted by the R term. More recordings from well-recorded earthquakes available at larger distance are usually required to well constrain the anelastic term. Therefore, this study focuses on the geometrical spreading term in the considered GMMs, as shown in Table 3.

[Table 3 about here.]

To further examine the misfits with respect to distance in the considered GMMs, several coefficients associated with the geometric spreading term in these GMMs are re-estimated using the Sichuan data. As shown in Table 3, the constant coefficient, the slope coefficient of the magnitude-independent geometrical spreading term and the fictitious depth are re-evaluated, including: a_1 , a_2 , and c_4 in ASK14; e_a (additive constant term), c_1 , and h in BSSA14; c_0 , c_5 , and c_7 in CB14; c_1 , c_{4a} , and c_{RB} in CY14; α_{11} , α_{12} , β_{11} , and β_{12} in I14; a_0 and a_2 in Huo89. The other model coefficients are fixed at the published values. It is worth noting that the present study may not be well suited to determining regional adjustments on constant terms due to the limited number of events. Therefore, the results in terms of constant term should be considered with caution. The re-estimated coefficients with 95% CI are shown in Table 4. The published fictitious depth coefficients of ASK14 and CY14 are defined based on their preliminary evaluations, thus, c_{4M} in ASK14 and c_{RB} in CY14 are unchanged. The fictitious depth coefficients of CB14 at long periods are fixed at published values because the re-estimation gives impractical values, possibly due to the limited events.

In addition, the changes of the standard error of the estimate (SEE) before and after the re-estimation, Δs , is provided to illustrate the improvement of model performance, which is calculated as follows,

$$SEE = \sqrt{\frac{\sum_{i} \sum_{j} (\varepsilon_{ij})}{df}}, \ \Delta s = \frac{SEE_{r} - SEE_{p}}{SEE_{p}}$$
(7)

where df is the degree of freedom of the considered GMM (i.e., the number of observations minus the number of parameters to be estimated in the considered GMM); subscript r and p correspond to the re-estimated and published GMMs, respectively.

The model adjustments in Table 4 not only address the bias observed in the intra-event residuals (i.e., Figure 9), but also account for the general offset between the Sichuan data and NGA-West2 models (i.e., c in Figure 8), particular at long periods. As shown in Table 4, the constant coefficients, slope coefficients, and fictitious depth coefficients of NGA-West2 GMMs for short-period spectral ordinates are similar to the published values, which lie within the 95% CI of the refitted values. This result shows that the refitted coefficients are not significantly different from the published values, which is consistent with the findings in Figure 9. However, for long-period spectral ordinates, the refitted coefficients of NGA-West2 GMMs are significantly different from the published values, which may be due to the general offset c observed in Figure 8. Moreover, the values of Δs show that there is an improvement in model performance before and after re-estimation when calibrating the NGA-West2 GMMs using the local data, particularly, at long periods (i.e., improved by about 20%). Regarding Huo89 GMM, the re-estimated constant coefficients and slope coefficients are significantly different from the published values, which is again consistent with Figure 9. The Δs values for the Huo89 show that the model performances are improved significantly, particularly at long periods.

[Table 4 about here.]

In general, the re-estimation has improved the model performance, especially for those GMMs with bias in distance scaling. Moreover, the bias in distance scaling of the newly calibrated GMMs is negligible. Thus, in the following sections, the intra-event residuals ε_{ij} are re-computed using the re-estimated models and the updated intra-event residuals, denoted by $\tilde{\varepsilon}_{ij}$, will be studied with respect to the V_{S30} based site effect.

8. Site-effect scaling

The scaling of ground motions with respect to the V_{S30} parameter is analyzed in this section. We examine the trends of the updated intra-event residuals (i.e., no distance bias) as a function of the logarithm of V_{S30} , which is also used in Scasserra et al. [43]. The linear regression as in Eq. (8) is used, similarly to the previous section:

$$\widetilde{\varepsilon}_{ij} = a_V + b_V \ln(V_{S30,ij}) + \epsilon_{ij} \tag{8}$$

where $\tilde{\varepsilon}_{ij}$ is the updated intra-event residual excluding the distance bias at site j for the event i in Eq. (6); a_V and b_V are the model coefficients; ϵ_{ij} is the error term.

The slope parameter b_V represents approximately the misfit of the V_{S30} scaling in the considered GMMs with respect to the Sichuan data. The fitted trend lines of the updated intra-event residuals $\tilde{\epsilon}$ versus $\ln(V_{S30})$ are shown in Figure 10 together with 95% CI of predictions. The results show that there is statistical evidence to reject the null hypotheses H_0 : $b_V = 0$ (i.e., no bias in V_{S30} -based site-effect scaling) at 5% significance level for ASK14, CB14, and CY14 at long periods, for BSSA at PGA and long periods. However, I14 shows no bias in V_{S30} -based site-effect scaling. These results show that the V_{S30} -based site effect in some NGA-West2 GMMs need to be adjusted for the Sichuan data, in particular, for long-period spectral ordinates. The Huo89 GMM, though it is developed for soil site without site

effect term in its functional form, shows no bias in terms of V_{S30} with respect to the Sichuan data (i.e., 1 - p < 0.95), which implies that the simplified site category may be a good proxy to partially account for the site effect.

[Figure 10 about here.]

The site condition has a strong implication on engineering applications, as it can alter the ground motion characteristics (i.e., amplitude, duration, and frequency content). The site effect in GMMs is commonly modeled by the linear site response term and the nonlinear site response term, which account for the general site effects and the complex effects due to unusual site conditions, respectively. The linear site response term is typically represented as $\ln V_{S30}$, while the nonlinear site response term relates to the ground motions at the bedrock with a more complex functional form. To constrain the nonlinear site response term, it usually involves the 1-D analytical site response model, which is not available for the Sichuan region. Thus, this paper focuses on the linear site response term, as shown in Table 3

To address the misfit corresponding to the V_{S30} -based site effect, the slope coefficients of the linear site-effect term is re-estimated. Regarding the Huo89 GMM, a V_{S30} -based term (i.e., $a_5 \log(V_{S30})$) is added to account for the site effects. It is worth noting that if there is statistically significant evidence to reject the null hypothesis that an additive constant term is equal to zero, then this constant term would be included as well. The Δs values are provided to illustrate the improvement of the model performance. The re-estimated coefficients with 95% CI are shown in Table 5.

[Table 5 about here.]

As shown in Table 5, the slope coefficient of the linear site response terms in NGA-West2 GMMs are similar to the published values at short periods, however, these coefficients are significantly different from the published values for long-period spectral ordinates except I14. The Δs values of BSSA14 show that the model performance is improved by over 15% at the long period while the performance improvement for the other four NGA-West2 GMMs at long periods is around 5%. The Huo89 GMM is generally consistent with the published value at short periods while those for moderate periods are significantly different from the published values. The Δs of the Huo89 GMMs shows limited improvement after re-estimation at moderate periods. These results are consistent with the findings in Figure 10. These improvements at long periods address the bias caused by the higher values of shear-wave quality factor Q for China than that implied by the NGA-West2 GMMs [36]. Moreover, \hat{c} results for BSSA14 obtained in this study are similar to the China results in Figure 4 of Seyhan and Stewart [50], confirming the results in this study.

9. Conclusion

This study investigated the compatibility of state-of-the-art NGA-West2 GMMs for shallow crustal earthquakes in active regions and the China-specific Huo89 model with the recent strong-motion data in Sichuan Province, China. The 2008 M_w 7.9 Wenchuan earthquake, the 2013 M_w 6.6 Lushan earthquake, and the 2017 M_w 6.5 Jiuzhaigou earthquake are considered as case-study events. This paper started with a brief review of several studies comparing various GMMs to Chinese ground-motion data. It is shown that the long-period spectral ordinates in Sichuan data may be overestimated by both the international and local GMMs. Moreover, it is shown that a variety of data and methods, in particular, different estimations/measurements of site conditions at the recording stations are employed in the literature, which may be due to the inadequate site information. A visual comparison of V_{S30} obtained from two estimation methods shows possible misclassification of site condition, which suggests an urgent need for detailed investigation of site characterization.

To avoid bias caused by the different horizontal ground-shaking definitions used in the considered GMMs, the RotD50 and geometric mean definitions are statistically evaluated and it is found there is no significant difference between these two definitions. Moreover, by studying the ratio between the RotD50 definition and the geometric mean definition, certain polarization in the ground motions is detected in the Sichuan data, which implies that these data should be used with caution for analyzing the structural responses of 3-D models in all orientations. Furthermore, the median ratio between the RotD50 and geometric mean definitions of the spectral ordinates is compared to the empirical model of Shahi and Baker [48], which shows the trend observed in the Sichuan data is consistent with that implied by the international model.

The visual comparisons of the observations in Sichuan data and the median predictions from the considered GMMs have shown that the magnitude scaling in each GMM is well modeled for the three case-study events. However, it is shown that the NGA-West2 and Chinese GMMs may require some modification to be suitable for the Sichuan region. Finally, a series of quantitative analyses have been performed on the inter- and intra-event residuals to assess the model performance. By using a mixed effects procedure, the results again show that the considered GMMs tend to overpredict the long-period spectral ordinates. The further study of distance dependency in the intra-event residuals has shown that CB14 and Huo89 models showed positive trend with increasing distance, implying that slower attenuation in Sichuan possibly due to the higher shear-wave quality factor Q for China. Moreover, the V_{S30} -based site dependency is found in some NGA-West2 models, particularly, for long-period spectral ordinates. To address the misfits corresponding to the distance scalings and the site-effect scalings in the considered GMMs with respect to the Sichuan data, the coefficients of the constant terms, the geometric spreading term and the linear site response term in these GMMs have been re-estimated and these updated GMMs with these minor modification can be applied for the PSHA in Sichuan Province.

While this work has focused on Sichuan Province, the presented method is applicable elsewhere in China. Future work will formally evaluate data from other regions in a manner similar to what is described here.

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List of Figures

1	(a) Distribution of recording stations and epicentres $(\stackrel{\wedge}{\varkappa})$ for the considered events. (b)	
	M_w - R_{JB} distribution. The soil/rock sites are denoted by square (\Box)/cross (+), respec-	
	tively. The color version of this figure is available only in the electronic version of the	
	paper	23
2	$V_{S30,Yu}$ and $V_{S30,USGS}$ for 80 stations. The soil/rock site is shown by square $(\Box)/cross$ (+).	
	The color version of this figure is available only in the electronic version of the paper	23
3	The median and standard deviation of $S_{a,{ m RotD50}}/S_{a,{ m GM}}$ against the structural period T_{\cdot} .	24
4	The Q-Q plot of $\ln(IM_{RotD50})$ and $\ln(IM_{GM})$: (a) PGA; (b) $S_a(T = 1.0)$	24
5	Displacement trace and spectral acceleration for a structural period $T = 0.1$ s in all	
	directions computed using ground motion records from (a) Wenchuan earthquake $51GYZ$	
	station, (b) Wenchuan earthquake 51MZQ station, (c) Lushan earthquake 51PJD station,	
	(d) Jiuzhaigou earthquake 62WYX station.	25
6	Median $S_{a,\text{RotD100}}/S_{a,\text{RotD50}}$ ratios compared with the empirical model proposed by Shahi	
	and Baker [48]	25
7	Comparisons between the observations and median predictions of (a-c) PGA, (d-f) $S_a(T =$	
	0.1), (g-i) $S_a(T = 1.0)$ of Wenchuan, Lushan and Jiuzhaigou earthquakes, respectively,	
	assuming $V_{S30} = 450$ m/s. The soil/rock site is denoted by square (\Box)/cross (+), respec-	
	tively. The color version of this figure is available only in the electronic version of the	
	paper	26
8	(a) Offset parameter c, (b) $c+\eta$ for Wenchuan earthquake, (c) $c+\eta$ for Lushan earthquake,	
	(d) $c + \eta$ for Jiuzhaigou earthquake, (e) inter-event standard deviation τ , (f) intra-event	
	standard deviation ϕ for the six GMMs against T. The upper and lower bars represent	
	the 95% CI. The line in (e) and (f) is the standard deviation prediction of CB14 GMM	
	given a reverse fault event in China with $M_W = 7.0$, $V_{S30} = 450$ m/s, $R_{rup} = 125$ km.	
	The color version of this figure is available only in the electronic version of the paper	27
9	Variation of intra-event residuals against source-to-site distance for PGA, $S_a(T = 0.1)$,	
	$S_a(T = 1.0)$. The color version of this figure is available only in the electronic version of	
	the paper	28
10	Variation of intra-event residuals against shear-wave velocity V_{S30} for PGA, $S_a(T = 0.1)$,	
	$S_a(T = 1.0)$. The color version of this figure is available only in the electronic version of	
	the paper	29

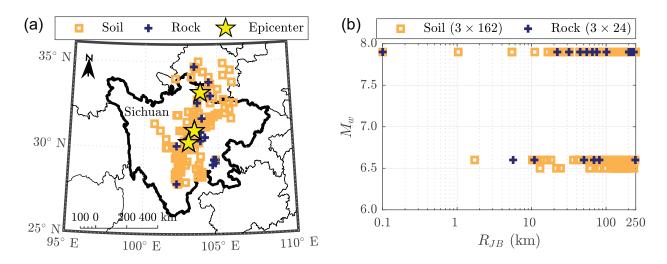


Figure 1: (a) Distribution of recording stations and epicentres $(\stackrel{\wedge}{\succ})$ for the considered events. (b) M_w - R_{JB} distribution. The soil/rock sites are denoted by square (\Box)/cross (+), respectively. The color version of this figure is available only in the electronic version of the paper

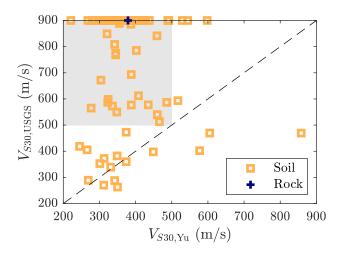


Figure 2: $V_{S30,Yu}$ and $V_{S30,USGS}$ for 80 stations. The soil/rock site is shown by square $(\Box)/cross$ (+). The color version of this figure is available only in the electronic version of the paper

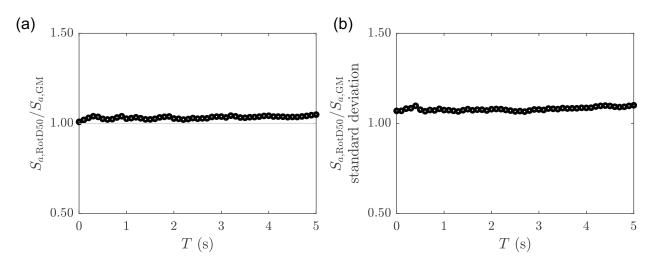


Figure 3: The median and standard deviation of $S_{a,\text{RotD50}}/S_{a,\text{GM}}$ against the structural period T.

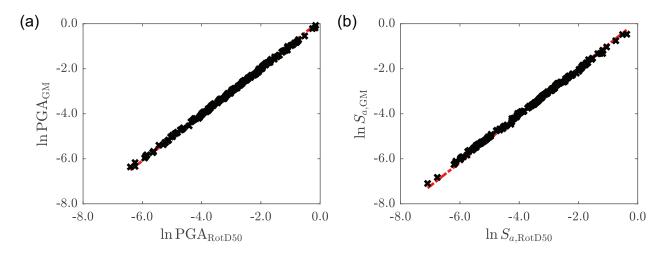


Figure 4: The Q-Q plot of $\ln(IM_{RotD50})$ and $\ln(IM_{GM})$: (a) PGA; (b) $S_a(T = 1.0)$.

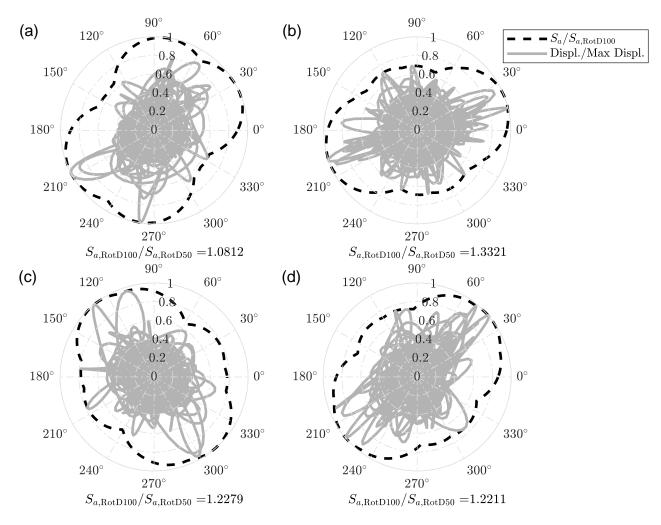


Figure 5: Displacement trace and spectral acceleration for a structural period T = 0.1 s in all directions computed using ground motion records from (a) Wenchuan earthquake 51GYZ station, (b) Wenchuan earthquake 51MZQ station, (c) Lushan earthquake 51PJD station, (d) Jiuzhaigou earthquake 62WYX station.

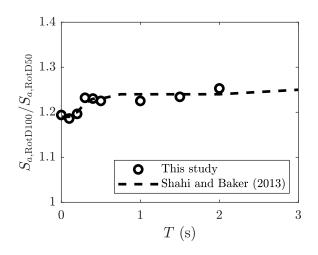


Figure 6: Median $S_{a,\text{RotD100}}/S_{a,\text{RotD50}}$ ratios compared with the empirical model proposed by Shahi and Baker [48]

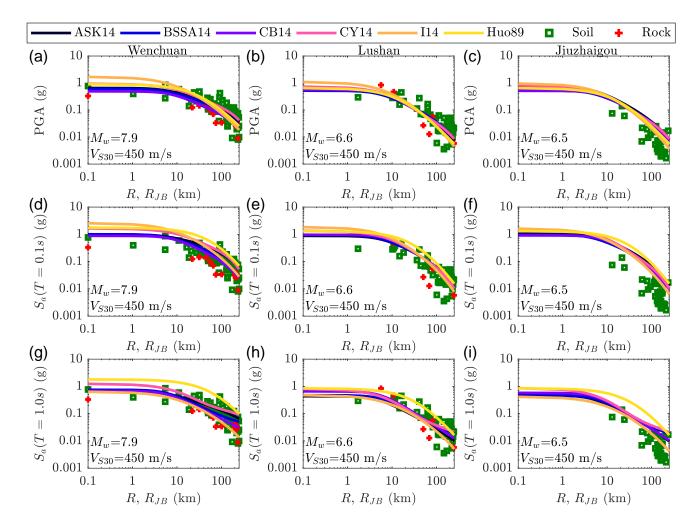


Figure 7: Comparisons between the observations and median predictions of (a-c) PGA, (d-f) $S_a(T = 0.1)$, (g-i) $S_a(T = 1.0)$ of Wenchuan, Lushan and Jiuzhaigou earthquakes, respectively, assuming $V_{S30} = 450$ m/s. The soil/rock site is denoted by square (\Box)/cross (+), respectively. The color version of this figure is available only in the electronic version of the paper

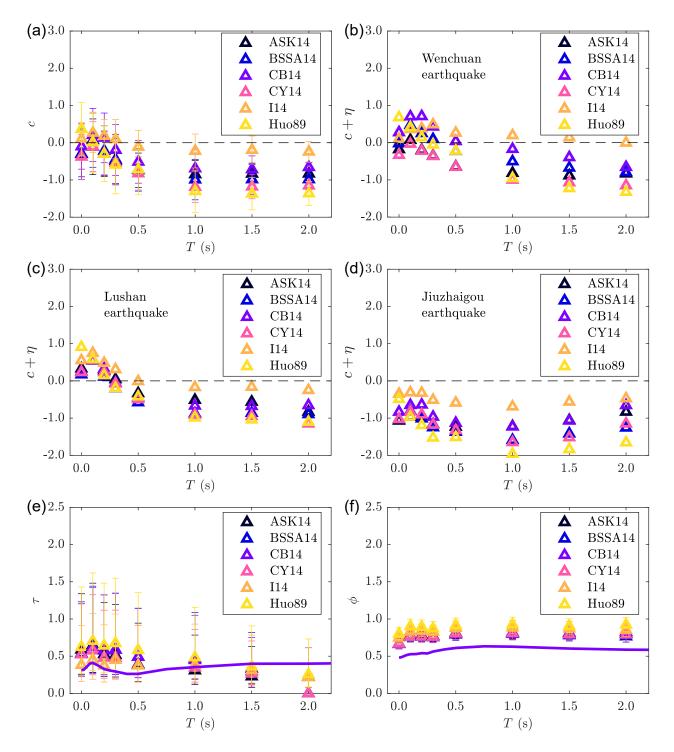


Figure 8: (a) Offset parameter c, (b) $c + \eta$ for Wenchuan earthquake, (c) $c + \eta$ for Lushan earthquake, (d) $c + \eta$ for Jiuzhaigou earthquake, (e) inter-event standard deviation τ , (f) intra-event standard deviation ϕ for the six GMMs against T. The upper and lower bars represent the 95% CI. The line in (e) and (f) is the standard deviation prediction of CB14 GMM given a reverse fault event in China with $M_W = 7.0$, $V_{S30} = 450$ m/s, $R_{rup} = 125$ km. The color version of this figure is available only in the electronic version of the paper

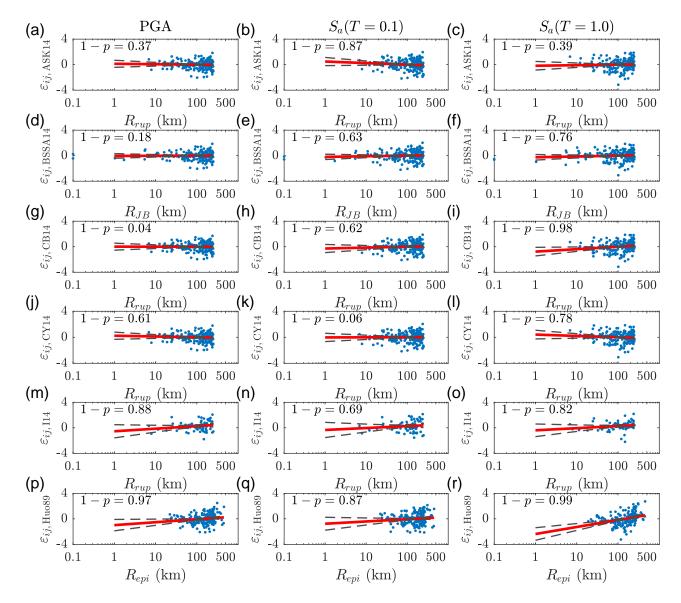


Figure 9: Variation of intra-event residuals against source-to-site distance for PGA, $S_a(T = 0.1)$, $S_a(T = 1.0)$. The color version of this figure is available only in the electronic version of the paper

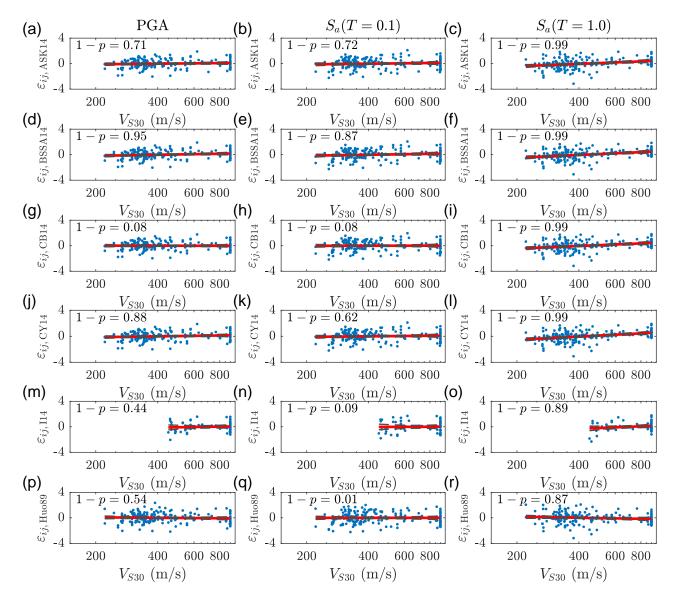


Figure 10: Variation of intra-event residuals against shear-wave velocity V_{S30} for PGA, $S_a(T = 0.1)$, $S_a(T = 1.0)$. The color version of this figure is available only in the electronic version of the paper

List of Tables

1	Site classification in the China Code for Seismic Design of Buildings (GB50011-2010) [40]	31
2	The metadata of selected events	31
3	The constant term, geometric spreading term, and linear site response term used in the	
	NGA-West2 and Huo89 GMMs	32
4	The re-estimated coefficients of distance scaling. 'n.c.' means no change to the published	
	value	33
5	The re-estimated coefficients of site-effect scaling. 'n.c.' means no change to the published	
	value.	35
A1	Summary table of relevant studies	37

$V_{S20} (m/s)$	I_0	I_1	II	III	IV
, 320 (1175)	Γ	Depth of	of upper s	soil layer (m)
$V_{S20} > 800$	0				
$500 < V_{S20} \le 800$		0			
$250 < V_{S20} \le 500$		< 5	≥ 5		
$150 < V_{S20} \le 250$		< 3	$3\sim 50$	> 50	
$V_{S20} \le 150$		< 3	$3 \sim 15$	$15\sim80$	> 80

Table 1: Site classification in the China Code for Seismic Design of Buildings (GB50011-2010) [40]

Table 2: The metadata of selected events

Event	Date and local time	Latitude $(^{\circ}N)$	Longitude $(^{\circ}E)$	Depth (km)	Faule style	M_w	Strike (°)	Dip (°)	Rake (°)	Length (km)	Width (km)
Wenchuan ^a	2008.05.12 14:28:04	31.0	103.4	14	Reverse	7.9	229.4	32.0	118.3	308	40
Lushan ^b	2013.04.20 08:02:46	30.3	103.0	13	Reverse	6.6	205.0	38.5	88.8	66	35
Jiuzhaigou ^c	2017.08.08 21:19:46	33.2	103.8	20	Strike-slip	6.5	148.5	68.9	-3.1	57	27

^a The finite-fault model is obtained from Wang et al. [2], 93 three-component records with $R_{JB} \leq 250$ km;

^b The finite-fault model is obtained from Wang et al. [3], 59 three-component records with $R_{JB} \leq 250$ km;

 $^{\rm c}~$ The finite-fault model is obtained from Wang et al. [4], 34 three-component records with $R_{JB} \leq 250$ km.

GMMs	Constant term ^a	Geometric spreading term	Linear site response term
ASK14	a_1	$\begin{cases} [a_2 + a_3 (M_w - M_1)] \ln (R) & M_w \ge 5.0\\ [a_2 + a_3 (5.0 - M_1)] \ln (R) & M_w < 5.0 \end{cases}$ where $R = \sqrt{R_{rup}^2 + c_{4M}^2}$ $c_{4M} = \begin{cases} c_4 & M_w > 5\\ c_4 - (c_4 - 1) (5 - M_w) & 4 < M_w \le 5\\ 1 & M_w \le 4 \end{cases}$	
BSSA14	none	$[c_1 + c_2 (M_w - 4.5)] \ln \left(\sqrt{R_{JB}^2 + h^2}/1.0\right)$	$\begin{cases} c \ln\left(\frac{V_{S30}}{760}\right) & V_{S30} \le V_c \\ c \ln\left(\frac{V_c}{760}\right) & V_{S30} > V_c \end{cases}$
CB14	<i>c</i> ₀	$(c_5 + c_6 M_w) \ln \left(\sqrt{R_{rup}^2 + c_7^2}\right)$	$c_{11} \ln \left(\frac{V_{S30}}{k_1}\right)$
CY14	c_1	$c_4 \ln\{R_{rup} + c_5 \cosh[c_6 \max(M_w - c_{HM}, 0)]\} + (c_{4a} + c_4) \ln\left(\sqrt{R_{rup}^2 + c_{RB}^2}\right)$	$\phi_1 \min\left[\ln\left(\frac{V_{530}}{1130}\right), 0\right]$
I14	$\begin{cases} \alpha_{11} & M_w \le 6.75 \\ \alpha_{12} & M_w > 6.75 \end{cases}$	$\begin{cases} (\beta_{11} + \beta_{21} M_w) \ln(R_{rup} + 10) & M_w \le 6.75 \\ (\beta_{12} + \beta_{22} M_w) \ln(R_{rup} + 10) & M_w > 6.75 \end{cases}$	$\xi \ln(V_{S30})$
Huo89	<i>a</i> ₀	$\frac{a_2 \log[R_{epi} + a_3 \exp(a_4 M_s)]}{M_{epi} \text{ is the surface magnitude; } R_{epi} \text{ is the surface magnitude; } R_{epi}$	none

Table 3: The constant term, geometric spreading term, and linear site response term used in the NGA-West2 and Huo89 GMMs

^a M_w is the moment magnitude; M_s is the surface magnitude; R_{rup} is the rupture distance; R_{JB} is the Joyner-Boore distance; R_{epi} is the epicentral distance; V_{S30} is the average shear-wave velocity of the upper 30 m soil; other parameters are estimated in the corresponding studies.

GMMs	T (s)	(Constant			Slope		Fict	itious dept	h	Δs^{b}
GIVINIS	1 (5)	Published	Refitted	\pm CI $^{\rm a}$	Published	Refitted	\pm CI	Published	Refitted	\pm CI	(%)
		a_1	\hat{a}_1		a_2	\hat{a}_2		c_4	\hat{c}_4		
	PGA	0.587	0.846	0.762	-0.790	-0.888	0.162	4.500	n.c.	-	-2
A CITZ 1 4	0.1	1.169	2.335	1.002	-0.790	-1.032	0.210	4.500	n.c.	-	-1
ASK14	0.5	1.571	0.969	0.751	-0.790	-0.803	0.162	4.500	n.c.	-	-20
	1	1.043	0.179	0.731	-0.790	-0.776	0.157	4.500	n.c.	-	-27
	2	0.329	-0.877	0.685	-0.790	-0.709	0.147	4.500	n.c.	-	-31
		e_a	\hat{e}_a		c_1	\hat{c}_1		h	\hat{h}		
	PGA	0	0.130	0.858	-1.134	-1.196	0.180	4.500	4.846	8.581	-2
DCCA14	0.1	0	0.408	0.945	-1.065	-1.103	0.199	4.130	5.024	10.002	-17
BSSA14	0.5	0	-0.749	0.825	-1.146	-1.109	0.176	5.340	3.701	6.070	-2
	1	0	-0.765	1.013	-1.193	-1.207	0.212	5.740	7.287	9.288	-21
	2	0	-1.132	0.880	-1.216	-1.167	0.185	6.540	7.178	8.304	-55
		c_0	\hat{c}_0		c_5	\hat{c}_5		c_7	\hat{c}_7		
	PGA	-4.416	-3.838	1.165	-2.773	-2.882	0.239	6.768	8.975	17.653	-1
CD14	0.1	-3.666	-2.506	1.556	-2.633	-2.783	0.315	7.294	15.999	18.949	-9
CB14	0.5	-8.379	-8.921	1.147	-2.296	-2.249	0.238	6.902	6.530	22.927	-6.0
	1	-11.011	-12.080	0.786	-2.158	-2.039	0.169	5.650	n.c.	-	-13.6
	2	-12.969	-14.701	0.698	-2.104	-1.869	0.150	6.632	n.c.	-	-23.7
		c_1	\hat{c}_1		c_{4a}	\hat{c}_{4a}		c_{RB}	\hat{c}_{RB}		
	PGA	-1.507	-1.059	1.179	-0.500	-0.653	0.243	50.000	n.c.	-	-6
CIV14	0.1	-0.561	0.022	1.425	-0.500	-0.619	0.292	50.000	n.c.	-	0
CY14	0.5	-1.651	-2.665	1.265	-0.500	-0.446	0.261	50.000	n.c.	-	-23
	1	-2.537	-2.940	1.221	-0.500	-0.647	0.252	50.000	n.c.	-	-38
	2	-3.415	-4.036	1.118	-0.500	-0.612	0.231	50.000	n.c.	-	-43
		α_{11}/α_{12}	$\hat{\alpha}_{11}/\hat{\alpha}_{12}$		β_{11}/β_{12}	$\hat{\beta}_{11}/\hat{\beta}_{12}$					

Table 4: The re-estimated coefficients of distance scaling. 'n.c.' means no change to the published value.

Continued on next page

GMMs	T (s)	(Constant			Slope		Fict	itious deptl	1	Δs^{h}
GIMINIS	1 (8)	Published	Refitted	\pm CI $^{\rm a}$	Published	Refitted	\pm CI	Published	Refitted	\pm CI	(%)
	DCA	7.089	7.110	1.805	2.994	2.893	0.379				
	PGA	9.014	9.107	2.520	2.949	2.949	0.542	-	-	-	-11.
	0.1	7.579	8.102	2.088	3.019	2.994	0.439				
	0.1	9.425	10.250	2.915	2.874	2.874	0.627	-	-	-	-15.
		9.214	8.038	1.726	2.856	2.592	0.363				. .
	0.5	11.448	11.852	2.411	2.351	2.351	0.518	-	-	-	-5.8
		7.045	6.314	1.565	2.734	2.601	0.329				
	1	9.857	9.948	2.185	2.068	2.068	0.470	-	-	-	-7.9
T 4 4	2	3.361	1.793	1.474	2.684	2.377	0.310				
I14	2	6.866	6.967	2.059	1.994	1.994	0.443	-	-	-	-4.7
		a_0	\hat{a}_0		a_2	\hat{a}_2					
	PGA	0.163	-0.202	0.601	-1.842	-1.576	0.266	-	-	-	-14
TT 0.0	0.1	1.700	0.997	0.849	-2.456	-2.116	0.361	-	-	-	-3
Huo89	0.5	1.219	-0.895	0.775	-2.135	-1.328	0.330	-	-	-	-16
	1	0.152	-2.338	0.733	-2.104	-1.251	0.313	-	-	-	-39
	2	-1.209	-3.782	0.669	-2.193	-1.332	0.287	-	-	-	-46

Table 4: continued from previous page

a CI is the confidence interval;

 $\mathrm{b}\,\Delta s$ is the changes of the standard error of the estimate (SEE) before and after the re-estimation.

GMMs	T (s)		Constant			Slope		Δs ^b
GIVIIVIS	1 (8)	Published	Refitted	\pm CI $^{\rm a}$	Published	Refitted	\pm CI	(%)
		Δa_1	$\Delta \hat{a}_1$		a_{10}	\hat{a}_{10}		
	PGA	0	n.c.	-	1.735	1.743	0.196	0
ASK14	0.1	0	n.c.	-	1.310	1.311	0.153	0
A5K14	0.5	0	n.c.	-	4.450	4.764	0.270	-1
	1	0	n.c.	-	4.300	4.652	0.255	-2
	2	0	-0.147	0.134	0.550	1.111	0.282	-4
		Δe_a	$\Delta \hat{e}_a$		С	\hat{c}		
	PGA	0	n.c.	-	-0.600	-0.575	0.165	0
BSSA14	0.1	0	n.c.	-	-0.487	-0.486	0.191	0
B55A14	0.5	0	n.c.	-	-0.969	-0.884	0.186	-24
	1	0	n.c.	-	-1.050	-0.663	0.183	-16
	2	0	n.c.	-	-1.039	-0.410	0.158	-18
		Δc_0	$\Delta \hat{a}_0$		c_{11}	\hat{c}_{11}		
	PGA	0	n.c.	-	1.090	1.037	0.139	0
CB14	0.1	0	n.c.	-	1.615	1.578	0.139	0
UB14	0.5	0	n.c.	-	2.355	2.525	0.328	0
	1	0	n.c.	-	1.447	1.808	0.324	-1
	2	0	n.c.	-	-0.514	-0.018	0.281	-3
		Δc_1	$\Delta \hat{c}_1$		ϕ_1	$\hat{\phi}_1$		
	PGA	0	n.c.	-	-0.521	-0.504	0.110	0
CV14	0.1	0	n.c.	-	-0.444	-0.440	0.125	-10
CY14	0.5	0	0.422	0.325	-1.004	-0.568	0.310	-2
	1	0	0.593	0.314	-1.094	-0.481	0.300	-4
	2	0	0.669	0.284	-1.115	-0.425	0.271	-6

Table 5: The re-estimated coefficients of site-effect scaling. 'n.c.' means no change to the published value.

Continued on next page

GMMs	T (s)		Constant			Slope		Δs $^{\rm b}$
GIVINIS	1 (5)	Published	Refitted	\pm CI $^{\rm a}$	Published	Refitted	\pm CI	(%)
		$\Delta \alpha$	$\Delta \hat{\alpha}$		ξ	ξ		
	PGA	0	n.c.	-	-0.854	-0.854	0.032	0
T1 4	0.1	0	n.c.	-	-0.757	-0.758	0.038	0
I14	0.5	0	n.c.	-	-1.023	-1.022	0.031	0
	1	0	n.c.	-	-1.009	-1.008	0.028	0
	2	0	n.c.	-	-0.851	-0.850	0.026	0
		Δa_0	$\Delta \hat{a}_0$		a_5	\hat{a}_5		
Huo89	PGA	0	n.c.	-	0	-0.001	0.022	0
	0.1	0	n.c.	-	0	-0.001	0.025	0
	0.5	0	1.198	0.928	0	-0.455	0.352	-2
Huo89	1	0	0.952	0.895	0	-0.362	0.339	-1
	2	0	n.c.	-	0	-0.001	0.021	0

Table 5: continued from previous page

a CI is the confidence interval;

b Δs is the changes of the standard error of the estimate (SEE) before and after the re-estimation.

Appendix

Table A1: Summary table of relevant studies

Reference	Dataset and IMs	Site classification or	Reference GMMs	Main findings
study		V_{S30} used		
Li et al.	• 569 single-component	Rock/ hard soil/ soft soil	PGA:	• PGA is higher than the predictions of the reference GMMs over
[12]	records for the		• California: Sadigh	the considered distance range;
	69		et al. [51]	• PGA attenuates more slowly with respect to R_{rup} than that im-
	quake, $R_{rup} \leq 500$		• Japan: Fukushima	plied by the reference GMMs;
			and Tanaka [52]	• Directivity effect is detected from the PGA distribution map;
	• PGA		• China: Wang et al. [53]	• Hanging-wall effect is detected from the PGA distribution map;
				• Vertical components of some near-field records are as large (regard-
				ing peak value) as, if not greater than, the horizontal components.
Li et al.	• 328 single-component	Code for Seismic Design	$PGA/S_a(1)$:	• PGA is higher than the predictions of the reference GMMs while
[28]	records for the	of Buildings (GB 50011-	• NGA-West: CB08	long-period spectral ordinate is lower than that of the reference
	Wenchuan earth-	2001) [54]		GMMs over the considered distance range;
	quake, $R_{rup} \leq 400$			• Directivity effect is detected from the IM distribution map;
	km			
	• PGA, $S_a(1)^a$			• Hanging-wall effect is detected from the IM distribution map;
				• No clear trend of ground motions with respects to site classification
Wang	• 72 two-component	$V_{S30} = 30/t$, where	PGA/PGV/	• PGA and short-period spectral ordinates are higher than the pre-
et al.	records for the	$t = \frac{H}{V_{s20}} + \frac{30-H}{500}$, and H	$S_a(0.02;10,20);$	dictions of the reference GMMs while long-period spectral ordi-
[29]	Wenchuan earth-	is the drilling depth in m	• NGA-West: AS08,	nates are lower than that of the reference GMMs over the consid-
	quake, $R_{rup} \leq 200$		BA08, CB08, CY08	ered distance range;
	km		V/H ratio:	
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Table	

• PGA, PGV^b , $S_a(0.02:10,20)^c$, V/H ratio ^d ratio ^d • Geometric mean• Geometric mean• Geometric mean• Geometric mean• Geometric mean• Geometric mean• Gametric mean• Gametric mean• FdA• FdA• PGA,• PGA,• Geometric mean• Geometric mean	VgaVgaShallow crust: Abra $S_a(0.02:10.20)^c$, V/H $S_{30}(0.02:10.20)^c$, V/HShallow crust: Abra $S_a(0.02:10.20)^c$, V/HBozorgnia and Camp- $ratio^d$ Bozorgnia and Camp- $Geometric meanEMM cited by WangGeometric meanet al. [29]Geometric meanEMM cited by WangGeometric meanEMGeometric meanEGO/Geometric meanEGO/Geometric meanEGO/Geometric meanEGOGeometric meanEGOGordEGOGOEGOGOEGOGOEGOGOEGO<$	Reference	Dataset and IMs	Site classification or	Reference GMMs	Main findings
PGA,PGVb, $(Shallow crust: Ahra- Sa, (0.02:10,20)^c, V/H)$ $(Sa, (0.02:10,20)^c, V/H)$ $Sa, (0.02:10,20)^c, V/H$ Bozorgnia and Camp- $ratio^d$ Bozorgnia and Camp- $(Cometric mean)$ $(Com$	PGA,PGVb, (1) (2)	study		V_{S30} used		
$S_a(0.02:10.20)^c$, V/Himmson and Silva [55], Bozorgnia and Camplaria bell [56]; Z. Yilmaz GMM cited by Wang et al. [29]• Geometric meanbell [56]; Z. Yilmaz GMM cited by Wang et al. [29]• Geometric meancommetric mean commetric et al. [29]• Gametric meancommetric bell [56]; Z. Yilmaz (SMM cited by Wang et al. [29]• Gametric meancommetric for solution• Gametric meancomplexity for solution• 64 wenchman earth- guake, $R_{rup} \leq 200$ PGA: tet al. [57]• 050/ soft soil ($V_{S30} < et al. [57]• 050/ soft soil (V_{S30} < tet al. [$	$S_a(0.02:10.20)^c$, V/Hhamson and Silva [55], Bozorgnia and Camp- bell [56]; Z. Yilmaz GMM cited by Wang et al. [29]• Geometric mean P_{ell}	Wang				• PGV has a similar standard deviation as that of the reference
ratiodBozorgnia and Camp- bell [56]; Z. Yilmaz GMM cited by Wang et al. [29]• Geometric meanGMM cited by Wang (GMM cited by Wang et al. [29]• Gametric meanEah. [29]• G4two-componentRock/ stiff soil ($V_{S30} >$ • 64two-componentRock/ stiff soil ($V_{S30} >$ • 61pasin:LocatisfortrecordsfortrecordsfortrecordsforWenchuanearth-250)soft soil ($V_{S30} >$ • NGA-West:huPGA:funde, $R_{rup} \leq 200$ wande, $R_{rup} \leq 200$ hu- NGA-West:hu- NGA-West:funde, $R_{rup} \leq 200$ er al. [57]quake, $R_{rup} \leq 200$ hu- NGA-West:hu- NGA-West:funde- NGA-West: $S_a(0.23:3), D_{s5-95}^{\circ}$ • Geometric mean- NGA-West:- Geometric mean- NGA-West:	ratiodBozorgnia and Campbell [56]; Z. Yilmaz bell [56]; Z. Yilmaz GMM cited by Wang et al. [29]•Geometric mean GMM cited by Wang et al. [29]• GMM cited by Wang et al. [27]• GMM cited by Wang et al. [57]• GMA -West: AB0S, CY08• GGM : $Game$ • GGM : $Game$ • GGM -West: $AB0S, CY08•Geometric mean•Geometric mean$	et al.			hamson and Silva [55],	GMMs (median prediction comparison is unknown);
• Geometric mean • Geometric mean • GMM cited by Wang • et al. [29] • et al. [20] • et al. [57] • NGA-West: km • PGA, • NGA-West: * AB08,CY08 • Geometric mean • Geometric mean • Geometric mean	• Geometric meanbell [56]; Z. Yilmaz GMM cited by Wang et al. [29]• $Gometric mean$ GMM cited by Wang et al. [29]• GMM cited by Wang et al. [29]•• GMM cited by Wang et al. [20]•• GMM cited by Wang et al. [20]•• GMM cited by Wang et al. [57]•• GMM cited by Wang et al. [57]•• GMM cited by Wang et al. [57]•• GMA cite fun•• GMA cite $S_a(0.2:3,3), D_{si-95}^{\circ}$ •• GCM cuttic mean $S_a(0.2:3,3)$;	9] con-	ratio ^d		Bozorgnia and Camp-	 W/H motion in towns of DCA and short nowind smoothed and instance and
GMM cited by Wang (GMM cited by Wang et al. [29]64two-component64two-componentRock/ stiff soil $(V_{530} > PGA:$ records for the250)/ soft soil $(V_{530} < PGA:$ records for theWenchuanearth-250)/ soft soil $(V_{530} < PGA:$ et al. [57]quake, $R_{rup} \leq 200$ quake, $R_{rup} \leq 200$ MB08, CY08hm $VGA-West:$ hm $PGA,$ $PGV,$ $e PGA,$ $PGV,$ $e rot.S_a(0.2:3.3), D_{s5-95}^{e}\bullet Geometric meanS_a(0.2:3.3); D_{s5-95}^{e}\bullet Geometric meanS_a(0.2:3.3); D_{s5-95}^{e}\bullet Red West:S_a(0.2:3.3); D_{s5-95}^{e}$	GMM cited by Wang et al. [29]64two-component for64two-component Rock/ stiff soil ($V_{530} >$ Fordsforrecordsfortrup250)/ soft soil ($V_{530} <$ PGA: ruprecordsfortrup250)quake, $R_{rup} \leq 200$ quake, $R_{rup} \leq 200$ et Al. [57]quake, $R_{rup} \leq 200$ thm $Menchuanearth-funrandom (V_{530} < 0.23,3), D_{s5-95}^{\circ}\sigma (0.23,3), D_{s5-95}^{\circ}\sigma (0.23,3), D_{s5-95}^{\circ}\sigma (0.23,3), D_{s5-95}^{\circ}\sigma (0.23,3), D_{s5-95}^{\circ}\sigma (0.23,3), D_{s5-95}^{\circ}\sigma (0.23,3), D_{s5-95}^{\circ}\sigma (0.23,3); D_{s5-95}^{\circ}$	inued	• Geometric mean		bell [56]; Z. Yilmaz	I have than the predictions of the reference GMMs in the near-field
et al. [29]64two-componentRock/ stiff soil ($V_{S30} > PGA$: records for the250)/ soft soil ($V_{S30} < 0$ Tecords for the250)/ soft soil ($V_{S30} < 0$ Wenchuanet al. [57]quake, $R_{rup} \leq 200$ quake, $R_{rup} \leq 200$ fmAB08,CY08PGA, BGA, $S_a(0.2.3,3), D_{s5-95}^{\circ}$ • Geometric mean $S_a(0.2.3,3):$ $S_a(0.2.3,3):$	et al. [29]• 64two-componentPock/ stiff soil ($V_{S30} >$ Per At.recordsforthe250)/ soft soil ($V_{S30} <$ • 64worchuanrecordsforthe250)/ soft soil ($V_{S30} <$ • 64worchuanearth-250)/ soft soil ($V_{S30} <$ • 64worchuanearth-250)/ soft soil ($V_{S30} <$ • 76A-West:hm• PGA,PGA, $S_a(0.23,3), D_{s5-95}^{\circ}$ • Geometric meanSa(0.23,3):• Geometric meanSa(0.23,3):• 600				GMM cited by Wang	while that are consistent in the far field. The V/H ratios in terms
64two-componentRock/ stiff soil ($V_{530} >$ PGA:• 64two-componentRock/ stiff soil ($V_{530} >$ PGA:recordsforthe250)/ soft soil ($V_{530} <$ PGA:recordsforthe250)/ soft soil ($V_{530} <$ • Taipei basin:Wenchuanearth-250)/ soft soil ($V_{530} <$ • Taipei basin:Wenchuanearth-250)et al. [57]quake, $R_{rup} \leq 200$ et al. [57]• NGA-West:hmAB08,CY08PGV:AB08,CY08• PGA,PGV,• NGA-West:AB08,CY08• Geometric mean $S_a(0.2:3,3)$: $S_a(0.2:3,3)$:	64two-componentRock/ stiff soil ($V_{S30} >$ PGA:• 64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe250/ soft soil ($V_{S30} <$ • Taipei basin:Umenhuanearth-250/ soft soil ($V_{S30} <$ • Taipei basin:Lohquake, $R_{rup} \leq 200$ 250)soft soil ($V_{S30} <$ • Taipei basin:Lohquake, $R_{rup} \leq 200$ 250)et al. [57]• NGA-West:km250)et al. [57]• NGA-West:AB08,CY08kmSa(0.2:3,3), D_{s5-05}^{e} • NGA-West:AB08,CY08• Geometric meanSa(0.2:3,3):• NGA-West:AB08,CY08				et al. [29]	of long-period spectral ordinates are higher than the predictions
64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe250)/ soft soil ($V_{S30} >$ PGA:recordsforthe250)/ soft soil ($V_{S30} <$ et al. [57]wenchuanearth250)et al. [57]et al. [57]quake, $R_{rup} \leq 200$ et al. [57]et al. [57]htm $S_a(0.2:3,3), D_{s5-95}$ et al. [57]e. Geometric mean $S_a(0.2:3,3)$: $S_a(0.2:3,3)$:	64two-componentRock/ stiff soil ($V_{S30} >$ PGA:• 64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe $250 /$ soft soil ($V_{S30} <$ •Wenchuanearth- $250 /$ soft soil ($V_{S30} <$ •Taipei basin:LohWenchuanearth- $250 /$ soft soil ($V_{S30} <$ •Taipei basin:LohWenchuanearth- $250 /$ soft soil ($V_{S30} <$ •Taipei basin:LohWenchuanearth- $250 /$ soft soil ($V_{S30} <$ •Taipei basin:LohWenchuanearth- $250 /$ soft soil ($V_{S30} <$ •Taipei basin:LohWenchuanearth- $250 /$ soft soil ($V_{S30} <$ •NGA-West:AB08,CY08MenchuenS_a(0.2:3,3), D_{s5-95}^{\bullet} •NGA-West:AB08,CY08AB08,CY08•Geometric meanS_a(0.2:3,3):S_a(0.2:3,3):S_a(0.2:3,3):S_a(0.2:3,3):					of the reference GMMs over the considered distance range;
64two-componentRock/ stiff soil ($V_{330} >$ PGA:recordsforthe250)/ soft soil ($V_{330} <$ orTrecordsforthe250)/ soft soil ($V_{330} <$ orWenchuanearth-250)soft soil ($V_{530} <$ orWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanearth-250)ororWenchuanforfororWenchuanforforforSa(0.2:3,3)D singerforSa(0.2:3,3)forforSa(0.2:3,3)forforSa(0.2:3,3)forforSa(0.2:3,3)for<	64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe250)/ soft soil ($V_{S30} >$ PGA:recordsforthe250)/ soft soil ($V_{S30} >$ et al. [57]Wenchuanearth-250)soft soil ($V_{S30} <$ et al. [57]quake, $R_{rup} \leq 200$ et al. [57]et al. [57]quake, $R_{rup} \leq 200$ et al. [57]et al. [57]Guake, $R_{rup} \leq 200$ et al. [57]et al. [57] $VBCA.West:$ $AB08, CY08$ km $AB08, CY08$ $radotic mean$ $S_a(0.2:3,3), D_{s5-95}^e$ \bullet Geometric mean $S_a(0.2:3,3):$					• Directivity effect is detected from IM distribution map;
64two-componentRock/ stiff soil ($V_{S30} >$ PGA:• 64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe 250 / soft soil ($V_{S30} <$ •Wenchuanearth- 250 / soft soil ($V_{S30} <$ •Taipei basin:Wenchuanearth- 250 / soft soil ($V_{S30} <$ ••Wenchuanearth- 250 / soft soil ($V_{S30} <$ ••Wenchuanearth- 250 /soil ($V_{S30} <$ •Wenchuanearth- 250 /soil ($V_{S30} <$ •Wenchuan $Sa0/$ er al. [57]•Wenchuan $Sa0/$ er al. [57]•Wenchuan $Sa0/$ $Sa0/$ •Sa(0.2:3,3). D_{s5-95}° $AB08, CY08$ Sa(0.2:3,3). $Sa(0.2:3,3).$ $Sa(0.2:3,3).$	64two-componentRock/ stiff soil (V_{S30} >PGA:recordsforthe 250 / soft soil (V_{S30} >PGA:recordsforthe 250 / soft soil (V_{S30} >et al. [57]Wenchuanearth- 250)soft soil (V_{S30} >et al. [57]Wenchuanearth- 250)et al. [57]quake, $R_{rup} \leq 200$ PGA .et al. [57]quake, $R_{rup} \leq 200$ PGA . PGA .Model, $R_{rup} \leq 200$ PGA . PGA .ender, $R_{rup} \geq 200$ PGA . PGA .e					• Hanging-wall effect is detected from IM distribution map;
64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe250)/ soft soil ($V_{S30} <$ • Taipei basin:Wenchuanearth-250)/ soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth-250)soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth-250)et al. [57]• NGA-West:AB08,CY08Menchric mean $S_a(0.2:3,3), D_{s5-95}^{\circ}$ • NGA-West:AB08,CY08AB08,CY08• Geometric mean $S_a(0.2:3,3); S_a(0.2:3,3):S_a(0.2:3,3):S_a(0.2:3,3):S_a(0.2:3,3):$	64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe 250 / soft soil ($V_{S30} <$ PGA:recordsforthe 250 / soft soil ($V_{S30} <$ et al. [57]Wenchuanearth- 250)soft soil ($V_{S30} <$ et al. [57]quake, $R_{rup} \leq 200$ et al. [57]et al. [57]quake, $R_{rup} \leq 200$ et al. [57]et al. [57]km PGA , PGV , $AB08, CY08$ km $S_a(0.2:3,3), D_{s5-95}^{e}$ PGV : $S_a(0.2:3,3), D_{s5-95}^{e}$ $AB08, CY08$ Geometric mean $S_a(0.2:3,3):$					
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• 64 two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe 250 // soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 // soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 // soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 // soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 // soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 // soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 // soft soil ($V_{S30} <$ • Taipei basin:LohWenchuan $earth-250// soft soil (V_{S30} <• NGA-West:AB08,CY08MenchuenS_a(0.2:3,3), D_{s5-95}• NGA-West:AB08,CY08AB08,CY08Geometric meanS_a(0.2:3,3):S_a(0.2:3,3):S_a(0.2:3,3):S_a(0.2:3,3)$	• 64 two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe 250 / soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 / soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 /soft soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 / 250 /et al. [57]•Wenchuanearth- 250 / $et al. [57]$ •NGA-West:Ich $R_{rup} \leq 200$ $R_{rup} \leq 200$ $R_{rup} \leq 200$ •NGA-West:Ich R_{rup} R_{rup} R_{rup} R_{rup} R_{rup} Ich R_{rup}					while long-period spectral ordinates attenuate more slowly than
64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe 250 /soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 /soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 /soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 / 250 /soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 / 250 /soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 / 260 /soil ($V_{S30} <$ • Taipei basin:LohWenchuanearth- 250 / 260 /soil ($V_{S30} <$ • NGA-West:Mank PGV , PGV : PGV : PGV : PGV :Sa(0.2:3,3), D_{s5-95}^{e} PGV : PGV : PGV :Geometric mean $S_a(0.2:3,3)$: $S_a(0.2:3,3)$:	64two-componentRock/ stiff soil ($V_{S30} >$ PGA:recordsforthe 250 / 50 /soil ($V_{S30} <$ •Taipei basin:LohWenchuanearth- 250 / 250 /soil ($V_{S30} <$ •Taipei basin:LohWenchuanearth- 250 / 250 /soil ($V_{S30} <$ •Taipei basin:LohWenchuanearth- 250 / 250 / $al. [57]$ •Hallquake, $R_{rup} \leq 200$ hu PGN •NGA-West:hu R_{rup} S_{00} PGN •NGA-West:hu $S_a(0.2:3,3), D_{s5-95}$ PGV :•NGA-West: $S_a(0.2:3,3), D_{s5-95}$ • PGN :• $AB08, CY08$ •Geometric mean $S_a(0.2:3,3)$: $S_a(0.2:3,3)$: $S_a(0.2:3,3)$					that implied by the reference GMMs
recordsforthe 250 /soil (V_{S30} • Taipei basin:LohWenchuanearth- 250)et al. [57]et al. [57]quake, $R_{rup} \leq 200$ 250 • NGA-West:AB08,CY08kmAB08,CY08PGV:AB08,CY08• PGA,PGV,• NGA-West:AB08,CY08• Geometric mean $S_a(0.2:3,3), D_{s5-95}^e$ • NGA-West:AB08,CY08	recordsforthe250)/ soft soil($V_{S30} <$ • Taipei basin:LohWenchuanearth-250)soft soil($V_{S30} <$ • Taipei basin:LohWenchuanearth-250)soft soil• NGA-West:quake, $R_{rup} \leq 200$ AB08, $CY08$ PGA,PGA:kmPGA,PGV;PGV:PGV:soft observationPGV;PGV:PGV:functionPGV;PGV:PGV:functionPGV;PGV:PGV:functionPGU:PGV;PGV:functionPGU:PGV;functionPGU:PGV:functionPGU:PGU:functionPGU:PGU:functionPGU: <td>Wen</td> <td>·</td> <td>Rock/ stiff soil (V_{S30} ></td> <td>PGA:</td> <td>• PGA and short-period spectral ordinates are higher than the pre-</td>	Wen	·	Rock/ stiff soil (V_{S30} >	PGA:	• PGA and short-period spectral ordinates are higher than the pre-
Wenchuanearth- earth-250)et al. [57]quake, $R_{rup} \leq 200$ km• NGA-West: AB08,CY08hm• PGA, PGV,PGV: PGV:• PGA, $S_a(0.2:3,3), D_{s5-95}^{\rm e}$ • NGA-West: AB08,CY08• Geometric mean $S_a(0.2:3,3)$: $S_a(0.2:3,3)$;	Wenchuanearth- earth-250)et al. [57]quake, $R_{rup} \leq 200$ km• NGA-West: AB08,CY08hm• PGA, $S_a(0.2:3,3), D_{s5-95}^e$ • NGA-West: PGV: • NGA-West: AB08,CY08• Geometric mean· Geometric mean $S_a(0.2:3,3)$:• Salos $S_a(0.2:3,3)$:	et al.	for	$250)/$ soft soil (V_{S30} <		dictions of the reference GMMs while long-period spectral ordi-
\leq 200 PGV, PGV, $^{-95^{e}}$ • NGA-West: $^{-95^{e}}$ • NGA-West: AB08,CY08 $^{-36}$ • NGA-West: AB08,CY08 $S_a(0.2:3,3):$	\leq 200 PGV, PGV, $^{-95^{e}}$ BGV: $^{-05^{e}}$ AB08,CY08 PGV: AB08,CY08 AB08,CY08 $S_a(0.2:3,3):$	[30]		250)	et al. [57]	nates are lower than that of the reference GMMs over the consid-
PGV, PGV: $^{-95}^{e}$ • NGA-West: AB08,CY08 $S_a(0.2:3,3)$:	PGV, PGV: PGV: PGV: PGV: PGV: PGV: AB08,CY08 $S_a(0.2:3,3)$:		ke, $R_{rup} \leq$		• NGA-West:	ered distance range, in particular, in the far field;
PGV, $^{-95^{e}}$ • NGA-West: AB08,CY08 $S_{a}(0.2:3,3)$:	PGV, PGV. PGV: considered distance range, in particular, in $^{-95^{\circ}}$ • $NGA-West$: • D_{S5-95} is higher than the predictions of th AB08,CY08 the considered distance range, in particula $S_a(0.2:3,3)$:		km		AB08, CY08	• PGV is lower than the predictions of the reference GMMs over the
$^{-95^{e}}$ • NGA-West: • AB08,CY08 • $S_a(0.2:3,3)$:	$^{-95^{\rm e}}$ • NGA-West: $AB08,CY08$ • D_{S5-95} is higher than the predictions of the considered distance range, in particula the considered distance range, in particula				PGV:	considered distance range, in particular, in the far field;
AB08,CY08 $S_a(0.2:3,3):$	AB08,CY08 AB08,CY08 Calendary and the production of the considered distance range, in particula $S_a(0.2:3,3)$:		$S_a(0.2:3,3), D_{s5-95}^{e}$		• NGA-West:	• D_{or} or is higher than the neadictions of the reference GMMs over
			• Geometric mean		AB08,CY08 $S_n(0.2:3,3)$:	the considered distance range, in particular, in the far field;

Reference	Dataset and IMs	Site classification or	Reference GMMs	Main findings
study		V_{S30} used		
Wen			• NGA-West: CY08	• Ground motions attenuate more slowly with respect to R_{rup} than
et al.			• California: Sadigh	that implied by the reference GMMs;
[30] con-			et al. [51]	• Directivity effect is detected from IM distribution map;
Daning			s5-95:	• Hanging-wall effect is detected from the near field data;
			• Shallow crust: Bom- mer et al. [58], Kemp-	• PGA on rock is lower than the prediction of reference GMMs while
			ton and Stewart [59]	PGA on soil is higher than that of the reference GMMs. However,
				this trend is not clear for PGV and D_{S5-95} .
Lu et al.	• 93 two-component	• V_{S30} is estimated by	PGA:	• PGA and short-period spectral ordinates are higher than the pre-
[31]	records for the	empirical method	• NGA-West: AS08,	dictions of the reference GMMs over the considered distance range
	Wenchuan earth-	based on the	BA08, CB08, CY08	while long-period spectral ordinates are consistent with that of the
	quake, $R_{rup} \leq 300$	horizontal-to-vertical	• Japan: Zhao et al. [60]	reference GMMs;
	km	response spectral ratio		• PGA and spectral ordinates attenuate more slowly with respect to
	• PGA, $S_a(0.1:5,11)$	[60]		R_{rup} than that implied by Zhao et al. [60] GMM but consistent
	• Cometrie mean	• $V_{S30}=30/t$ where		with the NGA-West GMMs;
		$H \ge 20, t=1; H<20,$		• No clear trend of ground motions with respect to V_{S30} is observed.
		$t = \frac{\pi}{V_{S30}} + \frac{10}{750}$		But long-period spectral ordinates at sites with $V_{S30}<300$ m/s
				are generally lower than the predictions of the reference GMMs.
Wen	• 140 three-component	Rock $(V_{S30} = 760)$ /Soil	PGA:	• PGA is generally consistent with the predictions of the GMMs of
and Ren	records for the Lushan	$(V_{S30} = 350)$	• NGA-West: AS08,	Yu and Wang [61], Huo89 and BSSA14 while it is higher than the
[32]	earthquake, $R_{rup} \leq 200$		BA08, CB08, CY08	predictions of the GMMs of Yu et al. [37], Lei et al. [62] over the
	km		▲ China: Huo80 Vii and	considered distance range, which is possibly due to the site effects.
			Wang [61],	PGA at the hanging wall and the footwall in the near

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Reference	Dataset and IMs	Site classification or	Reference GMMs	Main findings
study		V_{S30} used		
Wen	• PGA, PGV, $S_a(0.2:2,4)$		Yu et al. [37], Lei et al.	field is consistent with the predictions of the NGA-West GMMs;
and Ren [39] _{con-}	• Geometric mean		[62]	• Short-period spectral ordinates are consistent with the predictions
tinued			• NGA-West2: BSSA14	of the Huo89 and BSSA14 GMMs in the near field while short- period spectral ordinates are higher than these two GMMs in the
			• China: Huo89	far field. The long-period spectral ordinates are lower than the
			PGV:	predictions of the Huo89 and BSSA14 GMMs, in particular, in the
			• China: Huo89, Yu	far field.
			et al. [37]	• PGV is consistent with the predictions of the GMMs of Yu et al.
				[37] and Huo89, though Huo89 overpredicts PGV in the near field.
				• Hanging-wall effect is detected from near field data;
				\bullet Ground motions attenuate more slowly with respect to R_{JB} than
				that implied by BSSA14 possibly due to higher shear-wave quality
				factor Q.
Xie	• 69 two-component	NEHRP site clas-	$PGA/PGV/S_a(0.2:3,4):$	• PGA is generally consistent with the predictions of reference
et al.	records for the Lushan	sification: class B	• NCA-West· AS08	GMMs over the considered distance range;
[33]	earthquake, $R_{JB} \leq 350$	$(V_{S30}{=}760)/$ class C	E	• PGV is lower than the medictions of the reference GMMs over the
	km	$(V_{S30}=560)/$ class D $(V_{C30}=270)$		considered distance range;
	• PGA, PGV, $S_a(0.2:3,4)$			• The short-period spectral ordinates are generally consistent with
	• Geometric mean			the predictions of the reference GMMs over the considered dis-
				tance range while long-period spectral ordinates are lower than
				that of the reference GMMs over the considered distance range
				particularly in the far field;
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Reference	Dataset and IMs	Site classification or	Reference GMMs	Main findings
study		V_{S30} used		
Xie				• Hanging-wall effect is detected from near field data;
et al.				• Spectral ordinates observed at sites of NEHRP class D are higher
[33]				than the prediction of the reference GMMs possibly due to the bicher shear-wave quality factor O
Bai [34]	• 49 two-component	Infered from topographic	$PGA/S_a(0.06:10,n/a):$	• PGA and spectral ordinates are generally consistent with the pre-
	records for the Lushan	slope $[45]$	• NGA-West2: ASK14,	dictions of the reference GMMs over the considered distance range;
	earthquake, $R_{rup} \leq 200$ km		BSSA14, CB14, CY14, I14	• Directivity effect is detected from PGA distribution map;
	• $PGA, S_a(0.06:10, n/a)$			• Hanging-wall effect is detected from PGA distribution map.
	• GMBat150			
Ren	• 42 two-component	• If available, obtain	PGA/SA:	• PGA and short-period spectral ordinates are consistent with the
et al. [9]	records for the Ji-	metadata of the	• NGA-West2: ASK14,	prediction of the reference GMMs over the considered distance
	uzhaigou earthquake,	stations from NGA-	BSSA14	range while long-period spectral ordinates are lower than that of
	$R_{rup} \leq 400 \ {\rm km}$	West2 database;	• Europe: Bommer	the reference GMMs over the considered distance range, in partic-
	• PGA, PGV, D_{s5-95} ,	• Otherwise, V_{S30} is in-	et al. [63]	ular, in the far field;
	$S_a(0.04{:}5{,}\mathrm{n/a})$	ferred from empirical	• Italy: Bindi et al. [64]	• PGV is lower than the predictions of the reference GMMs over the
	• RotD50	relationships between	PGV:	considered distance range;
		GB500111-2010 classifi- c_{afion} and V_{con}	• NGA-West2: ASK14,	• D_{s5-95} is generally consistent with the prediction of the reference
		CAUTOIL ALLA V.230.	BSSA14	GMMs while it is slightly lower than the predictions of the refer-
			D_{s5-95} :	ence GMMs in far field;
			• shallow crust: Afshari	• Ground motions attenuate more slowly with respect to R_{rup} than
			and Stewart [05]	that implied by the reference GMMs possibly due to higher shear-
				wave quality factor Q .
				Continued on next page

Main findings	 PGA and short-period spectral ordinates are consistent with the predictions of the reference GMMs while long-period spectral ordinates are lower than that of the reference GMMs; This paper mainly focuses on developing correlation models for China. 	• • •	• Magnitude-dependency in the residuals of PGA is observed in some NGA-West2 GMMs and the China GMM. Magnitude-dependency in the residuals of short-period spectral ordinates are observed in some NGA-West2 while that in the residuals of long-period spectral ordinates are observed in all NGA-West2 GMMs;
Reference GMMs	PGA/S _a (0.01:2,13) • China: Huo89	 PGA/S_a(0.2:2,3): China: Yu et al. [37] NGA-West2: ASK14, BSSA14, CB14, CY14, I14 NGA-East: Goulet et al. [66] 	
Site classification or V_{S30} used	Soil only	 If available, V_{S30} is collected from the NGA-West2 database; Otherwise, if two-category site classification from China database is available, set V_{S30}=760 for rock and V_{S30}=360 for soil; 	 Otherwise, if site cate- gory from China Code [40] is available (i.e., soil thickness).
Dataset and IMs	• 1534 two-component records from 767 earthquakes from 2007 to 2014 in China, $5 \le M_s \le 7.5$, $R_{epi} \le 200$ km • PGA, $S_a(0.01:2,13)$	• 1517 two-component records from 10 sets of earthquakes (including multiple events) from 1985 to 2014 in and near China collected from multiple data sources, $4 \leq M_w \leq 7.9$, $R_{rup} \leq 1500$ km	 PGA, S_a(0.2:2,3) Geometric mean
Reference study	Ji et al. [35]	Dangkua et al. [36]	

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Reference Dataset and IMs	Site classification or	Reference GMMs	Main findings
study	V_{S30} used		
Dangkua	and V_{S20} is known),		• Distance-dependency in the residuals of PGA is observed in some
et al.	assign rock or soil site		NGA-West2 GMMs and the China GMM while that in the resid-
[36] con-	conditions based on		uals of spectral ordinates are observed in all NGA-West2 GMMs;
tinued	the site descriptions		• Ground motions attenuate more slowly with respect to R_{rup} than
	• Otherwise, assumes		that implied by the NGA-West2 GMMs possibly due to higher
	the site is rock site.		shear-wave quality factor Q .
a $S_{\alpha}(T)$ denotes the spectral acceleration of a structural period T:	of a structural period T :		

• h ž (T) BC

b PGV is peak ground velocity;

 $c S_a(T_1:T_2,n_T)$ denotes the spectral accelerations of n_T structural periods between T_1 and T_2 that are used for comparison;

d $\mathrm{V/H}$ ratio is the vertical-to-horizontal response spectral ratio;

e D_{s5-95} is the significant duration of 5%-95% definition.