

1 **Role of particle characteristics on the compression behaviour of gap-graded sands**

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32 **Abstract**

33

34 Recent research has investigated the compression in gap graded mixtures of sands with
35 combined mineralogy focusing on the key factors that might imply the occurrence of
36 convergent or non-convergent paths in compression (i.e. transitional or non-transitional
37 behaviour). From previous work, the mineralogy of the matrix composed by larger grains
38 seems to determine the possibility of the occurrence of transitional behaviour. Hence, if
39 there is a strong and stiff matrix made of quartz sand particles, which are either larger
40 than or at least of similar size to the other component, then non-convergent compression
41 paths (i.e. transitional behaviour) are likely to occur. As a further confirmation of this
42 hypothesis, this technical note presents the results of oedometer tests on the same range of
43 mixtures of a quartz sand and a carbonate sand as was used in previously published work,
44 but with the larger component being of the stiffer and stronger quartz sand. In agreement
45 with the hypothesis, transitional behaviour occurred.

46

47 **Keywords**

48 Laboratory tests; Sands; Compressibility

49

50

51 **Introduction**

52 In recent literature, the several papers on ‘transitional’ behaviour of reconstituted and natural
53 intermediate graded soils have highlighted the occurrence of non-convergent paths in
54 compression (e.g. Martins et al., 2001, Nocilla et al., 2006; Nocilla & Coop, 2008; Ventouras &
55 Coop, 2009; Ponzoni et al., 2014). In gap graded soils, even if the influence of particle features
56 has been investigated (i.e. particle nature, granulometry and mineralogy), a clear picture has not
57 yet emerged about what causes it (e.g. Martins et al., 2001; Ferreira and Bica, 2006; Shipton et
58 al., 2006; Carrera et al., 2011; Shipton and Coop, 2012; Shipton and Coop, 2015).

59

60 Analysing existing research on transitional behaviour in gap graded sand or sand-silt mixtures,
61 Ponzoni et al. (2017) highlighted that non-convergent compression paths and/or non-unique
62 Critical State Lines (CSLs) were detected for mixtures in which the matrix was made of larger
63 particles of quartz, no matter whether other component was of the same mineralogy or not. In
64 order to test this observations, albeit indirectly, they carried out an investigation involving a
65 series of oedometer and triaxial tests to determine the behaviour of gap graded mixtures with
66 larger particles of carbonate and smaller particles of quartz. In recent literature the importance
67 has been highlighted of packing in sandy silt mixtures (e.g. Chang at al. , 2015) along with the
68 existence of a transitional content of fines for which the mixtures pass from the coarse-grain
69 dominant region to a fine-grain dominant region (e.g. Chang at al., 2016; Zuo and Baudet,
70 2015). Using high stresses in order to be sure whether the compression paths would eventually
71 converge or not, Ponzoni et al. (2017) showed **the convergent** compression paths **for their**
72 **mixtures made of larger grains of carbonate. In detail, comparing their results with non**
73 **convergent compression paths of mixtures tested by Shipton & Coop (2012), which were**
74 **featured by a similar binary mineralogy (i.e. quartz and carbonate) but larger particles of**
75 **quartz, for all the they highlighted the convergence of the mixtures they tested did**
76 **converge**, confirming that the mineralogy of the larger grains seemed to determine the
77 occurrence or not of the transitional mode of compression behaviour. Indeed, the convergence

78 of the compression paths for all mixes of Ponzoni et al. (2017) of different proportions seemed
79 to confirm that the presence of a fabric made of larger particle of carbonate that is easier to
80 erase in mixtures in which quartz particles are always the smaller component, possibly allowing
81 particle arrangement, hence the destructurement of the fabric and a unique Normal Compression
82 Line (NCL) to be defined. In their investigation, as for the particle breakage, the values of the
83 overall relative breakage (Br) for the mixes tested were quite low, even if the one-dimensional
84 Normal Compression Lines (1D NCLs) appear to be unique, confirming that while significant
85 breakage can occur in soils with convergent behaviour it is not a prerequisite.

86 As a consequence, the authors hypothesised that a much more robust fabric that would be
87 difficult to erase should be present in a mixture with a quartz matrix when the other component
88 of the mixture has either smaller particles or particles of similar dimension, and either of the
89 same mineralogy or of a weaker type. In this case, the other, non-quartz component can break
90 severely, inhibiting breakage of the larger quartz particles, which might prevent a unique 1D
91 NCL being defined. Taking this into account, this technical note presents a new series of
92 oedometer tests that were conducted to confirm the role of the mineralogy of the larger grains.
93 As supposed by Ponzoni et al. (2017), transitional behaviour should be expected for gap graded
94 mixtures of sands with larger grains of a stronger type. However, such data did not exist in their
95 work for the same mineralogy of mixtures. Hence, in order to compare with what was
96 previously investigated, this study used the same mineralogy as Ponzoni et al. (2017) but with
97 the reverse composition, i.e. larger particles of the same stronger type (quartz).

98

99 **2. Materials, procedures and laboratory tests**

100 Ponzoni et al. (2017) found convergent behaviour in compression for quartz-carbonate mixtures
101 of Leighton Buzzard sand (LBS) and Carbonate sand (CS) with a low ratio R between d_{max} and
102 d_{min} , where the d_{max} is the mean value of the larger particle size distribution, and d_{min} is the
103 mean value of the smaller. Their mixtures had an R value equal to 3.02 and the matrix made of
104 larger particles of weaker carbonate grains. More often, in the literature, R is estimated as $R^* =$

105 D50/d50 (e.g. Cabalar and Hasan, 2013; Zuo and Baudet, 2015) which provides a rigorous
106 definition of this value. In Table 1, both values have been reported. The Leighton Buzzard sand
107 (LBS) used by Ponzoni et al. (2017) consisted of strong and sub-rounded particles, characterised
108 by high sphericity, a specific gravity G_s equal to 2.66 and grain dimensions used by them
109 between 0.212 and 0.80mm, while the Carbonate sand (CS) had angular and weak shelly
110 carbonate particles, with low sphericity, a specific gravity G_s equal to 2.71 and grain
111 dimensions between 0.212 and 0.3mm (Table 1).

112 This research therefore focused on the reverse combination with an identical quartz-carbonate
113 mineralogy with a low ratio R equal to 5.7 but larger particles of the stronger type (i.e. quartz).
114 High pressure one-dimensional compression tests were performed on specimens of mixtures
115 that were created artificially by mixing two soils of these different mineralogies. The Leighton
116 Buzzard sand (LBS) again consists of strong and sub-rounded particles, characterised by high
117 sphericity, a specific gravity G_s equal to 2.66 but grain dimensions between 0.85 and 2mm. The
118 Carbonate sand (CS), obtained by crushing a weak limestone, has sub-angular and weak
119 particles, with low sphericity, a specific gravity G_s equal to 2.72 and grain dimensions between
120 0.075 and 0.425mm. Single particle strengths σ_f have been measured (Tso and Wang, 2015) and
121 characteristic strengths of 406 MPa and 39 MPa were found for the LBS and for CS,
122 respectively. Before testing, the commercial granulometric distributions were confirmed with
123 particle size analyses by means of sieving. All four sand grain characteristics for the current
124 research and that of Ponzoni et al. (2017), are summarised for comparison in Table 1, while the
125 particle size distributions are shown in Fig. 1.

126

127 Five “mixtures” were considered. Two of them were 100% one mineral and three were mixtures
128 of the two soils; 80% Quartz/20% Carbonate, 50% Quartz/50% Carbonate, and 20%
129 Quartz/80% Carbonate by dry weight. Twenty-four oedometer Tests were carried out using a
130 conventional 38mm diameter fixed ring oedometer (up to stresses of 12.4MPa) and a 30mm
131 diameter floating ring oedometer, used to reach higher pressures (up to 20.5MPa) while

132 minimising wall friction. Wet compaction was used in order to create groups of samples with
133 the same initial specific volumes and control their repeatability.

134

135 The accuracy in the measurements of the initial specific volume, v_i , is of essential importance
136 otherwise apparent non-convergence could simply arise from poor accuracy. The initial specific
137 volumes, v_i , should be calculated from different methods based on measurements as
138 independent of each other as possible, although some inter-dependency is necessarily present
139 (Rocchi and Coop, 2014). For the oedometer tests presented in this study, when possible,
140 measurements were based on the initial and final dry unit weights and on the final water
141 contents when water loss could be prevented. Discarding any clearly anomalous data, the
142 absolute value of the maximum difference between the mean of the initial specific volumes and
143 the single calculated values was, on average, less than 0.01. The mean standard deviation
144 between methods was also equal to 0.01. G_s was not the same for the two sands so that the
145 specific volumes have to be determined by means of the phase equations.

146

147 **3. Results**

148 Compression data for the five mixtures are shown in Fig. 2. The oedometer tests show that at
149 higher stress levels it is possible to identify a unique one-dimensional normal compression line
150 (1D-NCL) in the v - $\log \sigma'v$ plane for the 100% one mineral samples, as expected (Coop and Lee,
151 1993; Coop, 2005) and highlighted in Fig. 2a. Despite this, the remaining three mixes (20%Q,
152 and 80%Q in Fig. 2b and 50%Q in Fig. 2c) showed a non-convergent compression behaviour.
153 The non-convergence of the compression paths for these mixtures, including those with a large
154 amount of weaker grains (20%Q), highlights a compression behaviour which is distinctly
155 different from that seen by Ponzoni et al. (2017), for which unique normal compression lines
156 occurred no matter what their proportions, as compared in Fig.2c, where compression paths of
157 mixtures for the same mix proportions (50%Q) are shown for the two programmes.

158 The mixtures of Fig. 2b and 2c are therefore examples of transitional soils, despite a similar low
159 R ratio and mineralogy as those of Ponzoni et al. (2017), in which the larger particles were of
160 carbonate. For the single mineral samples (0%Q and 100%Q), the convergence of compression
161 paths eventually occurs for effective vertical stresses that depend on the mineralogy and the
162 corresponding particle strengths, which are significantly different. Hence, for mixtures with a
163 quartz content of 100% the convergence occurs for effective vertical stresses of about 10MPa,
164 while if the soil is made of carbonate particles only (0%Q) the convergence is reached for
165 effective vertical stresses smaller than 8 MPa. The slope chosen for each 1D-NCL was set to a
166 value of 0.48, since the lines appear to be parallel, similarly to Ponzoni et al. (2017).

167

168 The comparison for single mineralogy samples, given in Fig. 2a shows that the influence of the
169 mineralogy on the location of the compression lines is that the 1D-NCL moves towards higher
170 stresses for the quartz. The relative positions confirm what was previously observed (Ponzoni et
171 al., 2017; Leleu and Valdes, 2007) and as observed by McDowell & Bolton (1998), for each
172 mineral the smaller particles are stronger and give a higher 1D-NCL, resulting in the 1D-NCLs
173 being closer to each other for the current tests than those of Ponzoni et al. (2017).

174

175 According to Ponzoni et al., 2014, the degree of convergence of the oedometer compression
176 data can be quantified by the coefficient m which is calculated from a graph plotting the initial
177 specific volumes against that at a common maximum stress (12,4MPa). For consistency the
178 initial specific volume was taken in each test at 20kPa vertical stress (v_{20}). For soils with fully
179 convergent paths, like for example a uniform sand (e.g. Coop & Lee, 1993), the gradient of the
180 data on this graph, defined as m , would be zero and for soils with perfectly parallel compression
181 paths $m=1$. The value of m for the oedometer curves in Fig.3a is clearly increasing with the
182 amount of quartz content up to the a maximum of 0.6 when the proportion of the two
183 mineralogies is balanced (50%Q) and tends to decrease toward the single mineralogy specimens
184 (100% and 0%Q) where the unique 1D-NCL can be identified again. For the convergent

185 behaviour of the mixes of Ponzoni et al. (2017), values of m equal to 0 can be assumed. The
186 ratio C_s/C_c between the swelling index C_s and the compression index C_c , in Fig. 3a, shows
187 higher values the more transitional is the soil (i.e. higher m) indicating a more gradual yielding
188 behaviour, which is another typical feature of transitional behaviour.

189

190 *3.1 Particle breakage*

191 Figure 3b shows an analysis of the particle breakage that was carried out on the five mixtures in
192 terms of Hardin's (1985) relative breakage Br . The final gradings were measured by means of
193 sieves. The convex trend of the line that interpolates the Br values seems to be the mirror image
194 of the trend of m in Fig. 3a.

195

196 If the relative breakage is compared with the values of Ponzoni et al. (2017) in Fig. 3b, the
197 overall breakage is similar, even when the percentage of quartz is low and even if the stress
198 levels are not quite the same. Although this confirms that the amount of overall breakage is not
199 a good guide as to whether transitional behaviour may occur or not, for this research, when
200 transitional behaviour occurs the quartz content seems to be an influencing factor for Br . The
201 amount of breakage for the larger quartz particles is low. This is probably due to the particle
202 strength of quartz and, even if in Fig. 4 there is breakage for the 100%Q, for the 50% Q
203 mixtures it is clear how breakage involves the carbonate particles only. However, it is possible
204 that there is an effect of density on breakage (Altuhafi & Coop, 2011): a close inspection of the
205 values of initial specific volume v_i on Fig. 3b, indicates a general trend for which denser
206 samples show higher breakage, in contrast to what was expected, but this may add to the scatter
207 of values.

208

209 **4. Conclusions**

210

211 According to Ponzoni et al. (2017), transitional behaviour should be expected for gap graded
212 mixtures featuring larger grains of a stronger type, no matter in which proportion. Hence, this
213 technical note has presented a new series of oedometer tests that were carried out investigating
214 the behaviour of gap graded mixtures with larger particles of quartz but using the same
215 mineralogy as Ponzoni et al. (2017).

216

217 As shown in Fig.3, the non-convergence of compression paths for all the three mixtures tested,
218 included those with a large amount of stronger grains, highlights a compression behaviour
219 which is very different from that detected by Ponzoni et al. (2017), for which unique NCLs were
220 encountered for all mixtures. The amount of overall breakage was found not to be a good guide
221 as to whether transitional behaviour may occur or not. However, for these mixtures the quartz
222 content might be an influencing factor for Br and the trend for Br is the mirror image of the
223 trend of the of degree of convergence m.

224

225 The non-convergence of the compression paths at high stresses for the three mixes seems to
226 confirm the presence of a fabric that is difficult to erase, in mixtures in which quartz particles
227 are the larger component. The absence of breakage after test for the quartz component in
228 mixtures shown in Fig. 4 can be a further confirmation of this assumption.

229

230 By comparing the behaviour of these mixtures with those previously investigated with a weaker
231 matrix made of carbonate sand particles, that are of larger size to the other component, the
232 mineralogy of the larger grains seemed to determine the possibility of the occurrence of the
233 transitional mode of compression behaviour. Further research should be carried out on other
234 binary mixtures of artificial soils investigating the relative breakage of the single components in
235 the mixture, if possible, in order to assess any influence in determining transitional or non-
236 transitional behaviour.

237

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240 the tests.

241

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Soil	Sand	Mineralogy*	Gs*	Shape ⁺	Grading [mm]	R	d ₅₀	R*
Ponzoni et al. (2017)	Crushed Limestone	Calcium carbonate (CaCO ₃ : 99%)	2.72	Low sphericity sub-angular	0.71-2.36	3.02	1.56	3.25
	Leighton Buzzard Sand	Quartz (SiO ₂ : 100%)	2.66	High sphericity sub-rounded	0.212-0.8		0.48	
This research	Crushed Limestone	Calcium carbonate (CaCO ₃ : 99%)	2.72	Low sphericity sub-angular	0.075-0.425	5,7	0,274	5.18
	Leighton Buzzard Sand	Quartz (SiO ₂ : 100%)	2.66	High sphericity sub-rounded	0.85-2		1,42	

+classification based on Powers (1953). * Thermogravimetric analyses and Gs investigation carried out by Ponzoni et al. (2017).

300

301 Table 1. Mixture properties: comparison with Ponzoni et al. (2017).

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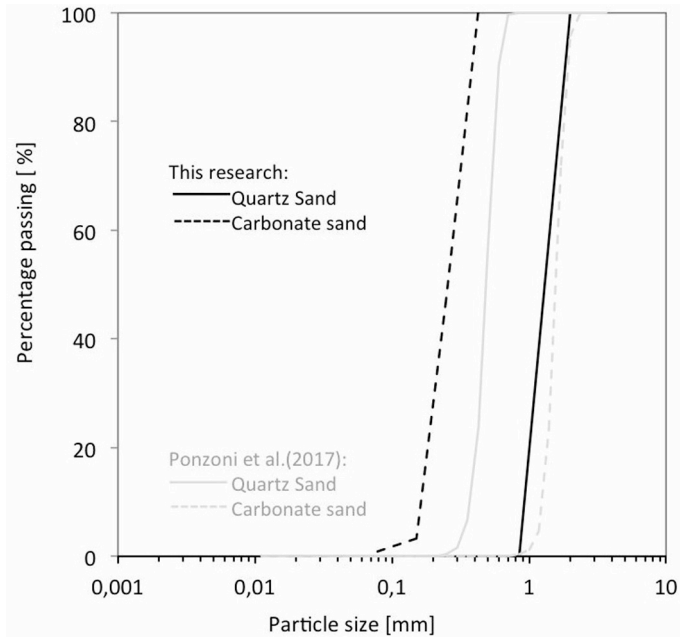
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313 **Figure captions**

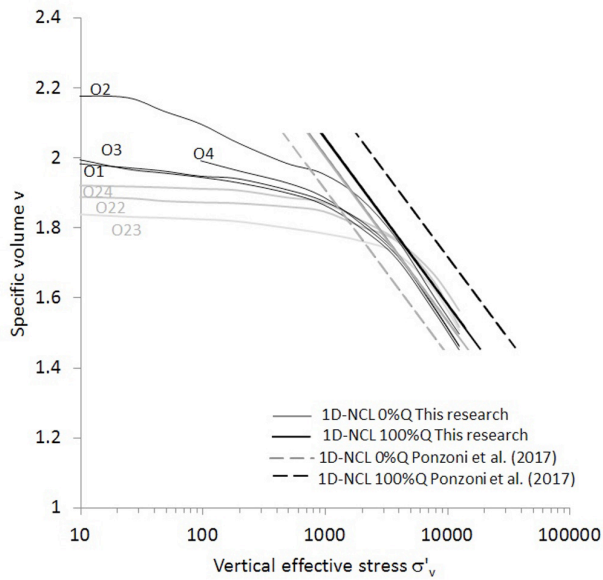


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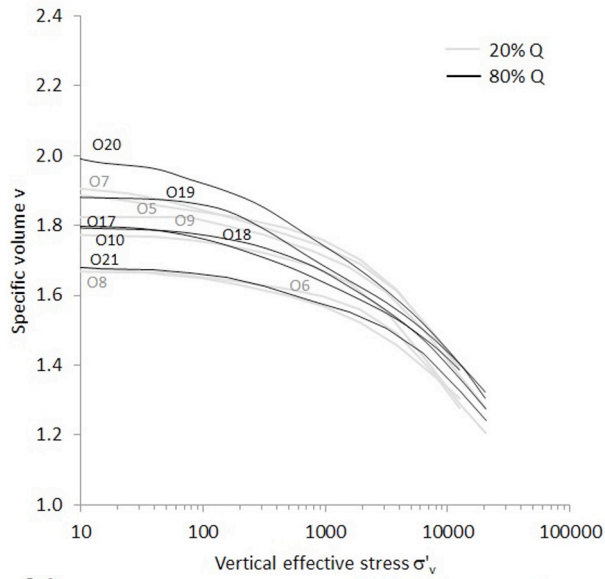
315 Figure 1. Particle size distributions of sands tested in this research and by Ponzoni et al. (2017).

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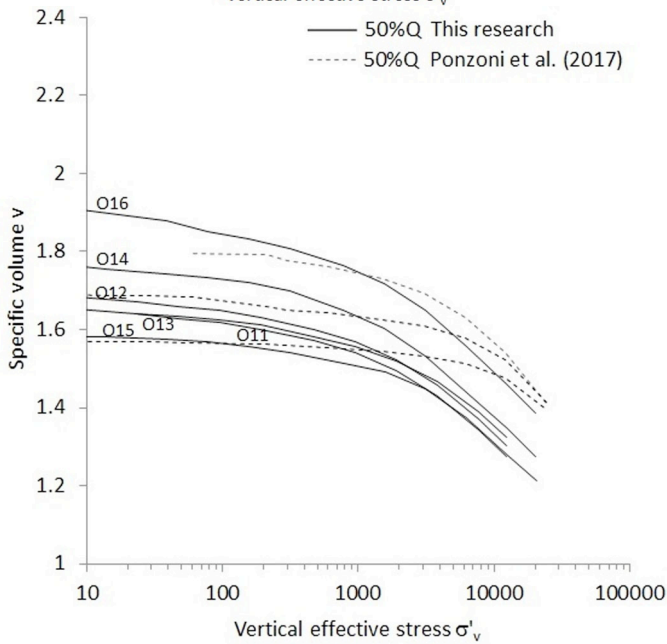
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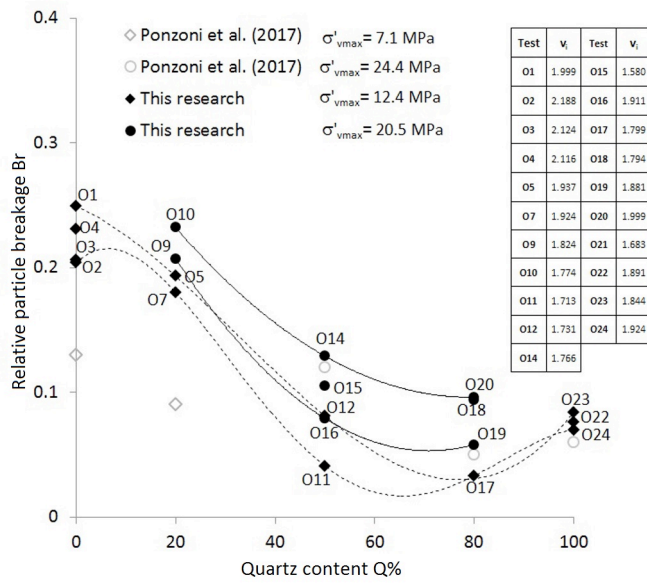
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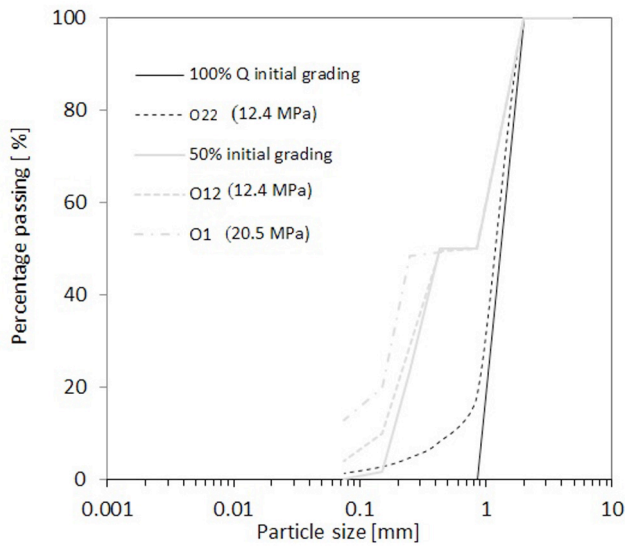
321 Figure 2. Oedometer compression curves for mixtures of quartz and carbonate sands: (a) 0%
 322 quartz and 100% quartz in this research and by Ponzoni et al., 2017; (b) 20% quartz and 80%
 323 quartz; (c) 50% quartz and comparison of compression behaviour between mix with 50%Q of
 324 this research and mix with 50% Q of Ponzoni et al. 2017;

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326

327 Figure 3. (a) Degree of convergence m and C_s/C_c ratio versus quartz content ($Q\%$). Comparison
 328 with Ponzoni et al. 2017. (b) Relative particle breakage (Br) versus quartz content ($Q\%$) and
 329 comparison with Ponzoni et al. 2017.



330

331 Figure 4. Particle size distribution before and after the test for 100% Q and 50% Q .

332