

Dynamic uplift, recycling, and climate control on the petrology of passive-margin sand (Angola)

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ABSTRACT

The subequatorial Angolan continental margin offers excellent conditions to test textbook theories on the composition of passive-margin sediments generated in different climatic and tectonic regimes. We use here comprehensive petrographic, heavy-mineral, geochemical and zircon-geochronology datasets on modern fluvial, beach, shelfal, and deep-marine sands and muds collected from hyperarid northern Namibia to hyperhumid Congo to investigate and assess: a) how faithfully sand mineralogy reflects the lithological and time structures of source rocks in a tectonically active rifted margin; b) in what climatic and geomorphological conditions the mark of

chemical weathering becomes strong and next overwhelming; and, c) to what extent the effect of weathering can be isolated from quartz dilution by recycling of older siliciclastic strata and other physical controls including hydraulic sorting and mechanical wear. A new refined classification of feldspatho-quartzose and quartzose sands and sandstones is proposed.

First-cycle quartzo-feldspathic to feldspar-rich feldspatho-quartzose sand eroded from mid-crustal granitoid gneisses of the Angola Block exposed in the dynamically uplifted Bié-Huila dome is deposited in arid southern Angola, whereas quartz-rich feldspatho-quartzose to quartzose sand characterizes the lower-relief, less deeply dissected, and more intensely weathered rifted margin of humid northern Angola. Pure quartzose, largely recycled sand is generated in the vast, low-lying hyperhumid continental interiors drained by the Congo River. The progressive relative increase of durable minerals toward the Equator results from three distinct processes acting in accord: active tectonic uplift in the arid south, and progressively stronger weathering coupled with more extensive recycling in the humid north. The quartz/feldspar ratio increases and the plagioclase/feldspar ratio decreases rapidly in first-cycle sand generated farther inland in the Catumbela catchment, reflecting stronger weathering in wet interior highlands. Discriminating weathering from recycling control is difficult in northern Angola. Although textural features including deep etch pits even on relatively resistant minerals such as quartz and microcline or rounded outline and abraded overgrowths provide valuable independent information, recycling remains as a most elusive problem in provenance analysis of terrigenous sediments.

Keywords: Provenance analysis; Classification of feldspatho-quartzose sands; Raman-counting of deep-water silt; U-Pb zircon geochronology; Weathering of detrital minerals; Quartzarenite problem.

"The present is the key to the past"

Archibald Geikie, *The Founders of Geology*

"From this gateway Moment a long eternal lane stretches backward: behind us lies an eternity. Must not what ever can happen, already have happened, been done, passed by before? And if everything has already been here before, what do you think of this moment, dwarf? "

Friedrich Nietzsche, *Thus Spoke Zarathustra*, *On the Vision and the Riddle*.

1. Introduction

1 Paleotectonic and paleoclimatic interpretations based on the composition of ancient sandstone suites
 2 often rely on assumptions seldom verified in a sufficient number of modern case studies ([Basu,](#)
 3 [1985](#)). The effect of chemical weathering in the weathering-prone climatic conditions found in
 4 equatorial and subequatorial settings has been investigated extensively in South America
 5 ([Franzini and Potter, 1983](#); [Johnsson et al., 1988, 1991](#); [Potter, 1994](#)), but less so in Africa (e.g.
 6 [Dupré et al., 1996](#); [Schneider et al., 2016](#)), which is an excellent location in which to test classic
 7 textbook theories on passive-margin sediments generated at diverse latitudes and in diverse climatic
 8 and tectonic regimes.

9 Sand deposited along passive margins is generally viewed as characterized by feldspatho-quartzose
 10 composition, with quartz/feldspar ratio controlled on the one hand by tectonic activity favoring
 11 rapid erosion of exhumed basement rocks and preservation of less stable detrital components, and
 12 on the other hand by chemical weathering that determines instead their progressive alteration and
 13 final breakdown ([Folk, 1980](#); [Dickinson, 1985](#)). Passive margins can be distinguished into volcanic
 14 (magma-rich) and magma-poor ([Franke, 2013](#)). Sediment sources in the latter are ancient crystalline
 15 basements exposed on rifted-margin shoulders and their cover strata including pre-rift, syn-rift and
 16 post-rift siliciclastic and carbonate rocks. Volcanic detritus derived from continental flood basalts
 17 and associated felsic products may be common and locally dominant in the former ([Garzanti et al.,](#)
 18 [2014a](#)).

19 Coastal Angola, located at subequatorial latitudes, is characterized by hyperarid climate in the south
 20 passing progressively northward to semiarid, humid, and eventually hyperhumid in the Congo ([Fig.](#)

21 1). The geology of the Angolan rifted margin is relatively simple and homogeneous, with the
22 Angola Block consisting of largely mid-Paleoproterozoic granitoid basement in the southern and
23 central parts and the metamorphic West Congo Belt of Neoproterozoic age in the north. After Early
24 Cretaceous rifting leading to the opening of the southern Atlantic Ocean, the region has undergone
25 successive and also recent stages of dynamic uplift (Giresse et al., 1984; Hudec and Jackson, 2004;
26 Al-Hajri et al., 2009; Guiraud et al., 2010; MacGregor, 2013; White, 2016; Green and Machado,
27 2017). Coastal Angola represents therefore an exceptionally suitable natural laboratory in which to
28 investigate how the detrital signature of daughter sand may respond to the lithological and time
29 structures (i.e., the diverse age patterns of source areas obtained by different geochronological
30 methods) of parent rocks, and to assess to what extent sand mineralogy can be modified during the
31 sedimentary cycle by chemical weathering as well as by diverse physical factors (i.e., mechanical
32 breakdown, hydraulic sorting, and recycling).

33 This study illustrates an integrated petrographic, mineralogical, geochemical and geochronological
34 dataset on modern sands collected along the entire Angolan continental margin, which complements
35 data obtained with the same methods and following the same criteria during previous studies on
36 sediment generation in southwestern Africa (Vermeesch et al., 2010; Garzanti et al., 2012a, 2014a,
37 2014b, 2014c, 2015, 2017; Dinis et al., 2016, 2017), to which the reader is referred as aspects
38 specific to adjacent regions are concerned.

39

40 **2. The Angolan rifted margin**

41

42 *2.1. Basement rocks*

43 The oldest rocks of subequatorial Africa belong to the Archean to Paleoproterozoic Kalahari and
44 Congo cratons, welded together by a series of orogenies including the major Namaqua arc-
45 accretion and continent-collision event between ~ 1.1 and 1.05 Ga (De Waele et al., 2008; Jacobs et
46 al., 2008). West Gondwana was finally amalgamated during the Neoproterozoic to Cambrian

47 Damara ("Pan-African") orogeny (Gray et al., 2008; Heilbron et al., 2008; Vaughan and Pankhurst,
48 2008; Frimmel et al., 2011).

49 In western Africa, the Congo craton is represented by the Angola Block, the core of which consists
50 of felsic Eburnean (~ 2 Ga) plutonic and metamorphic rocks (Fig. 1D; De Carvalho et al., 2000;
51 McCourt et al., 2013). Along its northeastern limit, Neoproterozoic granites, gneisses and migmatites
52 occur together with gabbros, norites and charnockites (Liberian-Limpopo massifs; Delhal et al.,
53 1976; De Carvalho et al., 2000). In southern Angola to northernmost Namibia, where it is
54 represented by the Epupa metamorphic unit, the Angola Block is intruded by one of the largest
55 anorthosite bodies on Earth, dated at 1.37-1.38 Ga and including norite and gabbro (Mayer et al.,
56 2004; Drüppel et al., 2007).

57 The West Congo and Kaoko belts were formed during the Neoproterozoic-Cambrian orogeny and
58 both display progressively increasing metamorphic grade from only mildly deformed
59 Neoproterozoic foreland units in the east to medium-grade or even high-grade rocks in the west.
60 The West Congo Belt, stretching from southwestern Gabon to northwestern Angola, comprises
61 Paleoproterozoic basement gneisses intruded by ~ 2 Ga granites and thrust NE-ward onto a
62 Neoproterozoic succession beginning with a thick volcano-sedimentary unit including a bimodal
63 magmatic suite dated between 1.0 and 0.9 Ga (Tack et al., 2001). The overlying West Congolian
64 Group, only gently folded and unmetamorphosed in the east, includes siliciclastic strata with
65 intercalated pillow-basalts and diamictites, capped by a stromatolite-bearing carbonate platform
66 followed in turn by siliciclastic syn-orogenic to post-orogenic sediments (Frimmel et al., 2006). The
67 Kaoko Belt, located at the southwestern tip of the Congo craton and representing the transpressive
68 northern arm of the Damara orogen, comprises the Epupa Paleoproterozoic basement and an arc
69 terrane with high-grade metasediments and granitoids close to the coast (Goscombe et al., 2003;
70 Goscombe and Gray, 2007).

71

72 *2.2. Syn-rift to post-rift sedimentary successions*

73 Rifting and initial sea-floor spreading in the South Atlantic progressed northward during the Early
74 Cretaceous (Moulin et al., 2010). Tectonic evolution was controlled by oceanic fracture zones
75 extending through continental crust commonly along Precambrian structures and locally leaking
76 mantle-derived magmas (Mohriak and Rosendhal, 2003). The conjugate Brazilian and African
77 rifted margins were thus segmented into transform, oblique and normal-rifted tracts separated by
78 transverse fault zones (Guiraud et al., 2010). Along the Angolan coast, Precambrian basement is
79 non-conformably overlain by a several km-thick Lower Cretaceous to Neogene succession
80 accumulated in distinct depocenters (Lower Congo, Inner and Outer Cuanza, Benguela, and Namibe
81 basins) and subdivided into syn-rift continental deposits followed by evaporites and post-rift marine
82 sediments (Hudec and Jackson, 2002; Chaboureau et al., 2013).

83 The up to 5 km-thick pre-salt strata of Neocomian to Aptian age include feldspar-rich fluvio-
84 lacustrine conglomerates, sandstones and shales, lacustrine carbonates, and interbedded lava flows
85 and ash deposited during tectonic extension (Bate et al., 2001; Sabato-Ceraldi and Green, 2017).
86 Salt deposits mostly consisting of massive halite were deposited at mid-Aptian times from north of
87 the Walvis Ridge to Gabon. Some tens of m-thick onshore, evaporites reach up to 1.4 km offshore
88 and overstep the underlying pre-salt wedge both landward and seaward, suggesting direct
89 deposition onto oceanic crust formed close to sea-level (Marton et al., 2000). Halokinesis impacted
90 greatly sedimentation at subsequent stages and determined the structural style of the Angolan
91 passive margin (Fort et al., 2004). The post-salt succession is up to 5 km-thick. Fully marine
92 conditions returned in the Albian, with deposition of platform carbonates passing to deep-water
93 mudstones offshore. Siliciclastic sedimentation, mainly confined to the Benguela and Namibe
94 basins in the Albian (Quesne et al., 2009; Gindre-Chanu et al., 2015), became dominant in the
95 Upper Cretaceous and Cenozoic (Séranne and Anka, 2005). Prograding clastic wedges and repeated
96 erosional gaps in the coastal plain and shelf document several phases of inner-margin uplift during
97 the Neogene (Jackson et al., 2005). In the vast Kalahari basin, stretching ~ 2200 km in the
98 hinterland from the Congo to South Africa, up to 450 m-thick largely fluvial sediments were

99 deposited possibly since the Late Cretaceous, and a huge sand sea was formed by eolian reworking
100 during drier periods of the Plio-Pleistocene to early Holocene (Haddon and McCarthy, 2005).

101

102 *2.3. Syn-rift to post-rift volcanic rocks*

103 Syn-rift volcanism reaching climax around 132 Ma is best represented by the Etendeka continental
104 flood basalts and quartz latites of northern Namibia, situated landward of the Walvis Ridge (Renne
105 et al., 1996; Ewart et al., 2004). The Angolan margin is instead mainly volcanic-poor (Séranne and
106 Anka, 2005; Péron-Pinvidic et al., 2017), although Lower to Upper Cretaceous tholeiitic to alkaline
107 bimodal volcanic products are found in several localities of the Namibe and Cuanza basins (Fig.
108 1D; Masse and Laurent, 2016). In the Namibe basin, tholeiitic basalts overlain by felsic stratoid
109 flows and travertine rest unconformably on Precambrian basement or pre-rift Karoo sandstones,
110 whereas basaltic flows overlying ammonite-bearing strata are dated as Santonian (84.6 ± 1.5 Ma;
111 Alberti et al., 1992; Strganac et al. 2014; Gindre-Chanu et al., 2016). In the Sumbe area of central
112 Angola, ~ 100 m-thick tholeiitic lavas dated at ~ 132 Ma and altered pyroclastic rocks directly
113 overlie the Precambrian basement, which is intruded by tholeiitic dykes subparallel to the coast.
114 Sodic alkaline and transitional products dated at ~ 91 Ma are intercalated with Upper Cretaceous
115 marine strata (Marzoli et al., 1999). Volcanism took place at the landward extension of the Cuanza
116 seamount chain built on oceanic crust, which acted as a sedimentary barrier between the Benguela
117 and Cuanza basins (Hudec and Jackson, 2002). Reactivation of deep lithospheric faults during Early
118 Cretaceous rifting was associated with kimberlite and carbonatite magmatism concentrated along
119 the 50-90 km-wide Lucapa tectonic corridor extending to the NE of the Namibe basin (Comin-
120 Chiamonti et al., 2011; Castillo-Oliver et al., 2016).

121

122 *2.4. Climatic gradients and weathering conditions*

123 Angola is situated between ~ 17°30'S and ~ 6°S, roughly at the same distance from the Tropic of
124 Capricorn and the Equator and thus right in the middle of the subequatorial belt. The coastal zone,
125 stretching N/S for ~ 1450 km between the mouths of the Cunene and Congo rivers, experiences a

126 latitudinal climatic gradient. Conditions change progressively from hot desert in coastal Namibia
127 and southern Angola to hot semi-arid in the coastal Benguela region, and next to tropical savanna
128 northward toward the humid Congo (Fig. 1C). The aridity of the southern region results from the
129 influence of quasi-stationary anticyclonic conditions on most austral Africa coupled with the
130 Benguela upwelling system responsible for low sea-surface temperatures and low-humidity
131 southerly winds (Shannon and Nelson, 1996; Lancaster, 2002). The cold and nutrient-rich Benguela
132 Current flows equatorward all along the Atlantic coast of southern Africa as the eastern branch of
133 the subtropical gyre, until it converges with the warm and nutrient-poor southward-flowing Angola
134 Current representing the eastern part of the Angola gyre (Lass and Mohrholz, 2008). Aridity
135 decreases progressively north of the Angola-Benguela Front, located between 20°S and 14°S and
136 moving northward in July-September and southward in January-March (Kostianoy and Lutjeharms,
137 1999; Hardman-Mountford et al., 2003). Along the coast, average annual precipitation varies from
138 < 100 mm in the hyperarid south, where rainfall is rare and exceeded by water generated from
139 condensation and fog (locally called "cassimbo"), to 400-600 mm north of Luanda. Rainfall
140 increases more rapidly up to ~1500 mm inland, where climate becomes humid subtropical or
141 temperate-highland tropical with dry winters at high elevation. The isohyets trend approximately
142 E/W in the continental interior, where the latitudinal gradient is much stronger than along the coast
143 (Fig. 1C).

144 Average annual temperatures vary little through much of coastal Angola, from 21-27°C north of
145 10°S to 20-24°C in the south, and drop to ~15°C only in the Moçamedes Desert and in the
146 highlands, which reach 2619 m a.s.l. at Mount Moco in the headwaters of the Balombo River
147 (Diniz, 2006). The warm rainy season, longer at lower latitudes and in the highlands, starts between
148 September and November and lasts until March to May depending on the region, with rainfall peak
149 in February to March. Weak southerly to south-westerly winds prevail all over the year in the
150 coastal region, where north of Namibe eolian sand transport becomes negligible.

151 Different weathering conditions controlled by the two broadly perpendicular N/S and E/W climatic
152 gradients, and consequently different intensities of pedogenic processes, are reflected by clay-

153 mineral assemblages and chemical-weathering proxies in river muds (Dinis et al., 2017). Kaolinite,
154 supplied in abundance to the coast by most major Angolan rivers, is largely generated in the wet
155 hinterland, and particularly on ancient flat surfaces decreasing in elevation westward and separated
156 by escarpments of variable relief. Expansive clays, instead, are mainly formed in Meso-Cenozoic
157 basins located in dryer areas along the coast. As a consequence, the smectite/kaolinite ratio is ~ 1 in
158 arid southern Angola, decreases rapidly in the semiarid Benguela region, and is very low in the
159 more humid north with the exception of muds carried by the Longa, Cuanza, and Bengo rivers
160 cutting across the Meso-Cenozoic Cuanza basin in their lower course.

161

162 2.5. River systems

163 Apart for the Congo (4,000,000 km², length 4700 km), the largest Angolan rivers draining into the
164 Atlantic Ocean are the Cunene ($\sim 110,000$ km², ~ 1050 km) and the Cuanza ($\sim 150,000$ km², ~ 960
165 km). The Longa, Queve, Curoca, and Mebridege are ~ 300 km-long and drain $\sim 20,000$ km² each;
166 the Catumbela and Coporolo are 200-250 km-long and drain $\sim 15,000$ km² (Fig. 1B). Of similar
167 length are the Loge, Dande, Bengo, and Bero, which also drain more than 10,000 km² each, and
168 another dozen ephemeral rivers draining between 2000 and 7000 km². Three major rivers are
169 sourced in the very same area between the cities of Huambo and Katchiungo at ~ 1800 m a.s.l. on
170 the dynamically uplifted Bié-Huila dome: the Queve draining northward, the Cunene draining
171 southward, and the Cubango branch of the Okavango draining southeastwards across the Kalahari
172 (Fig. 1). The Cuanza is instead sourced in the Kalahari and flows northwards around the Angola
173 Block. These river courses may have formed soon after opening of the South Atlantic, although the
174 modern configuration with radial drainage of the Bié-Huila dome was acquired much more recently
175 in response to dynamic uplift, as documented by morphometric analysis (Pritchard et al., 2009;
176 Anka et al. 2010). Few rivers have regular concave equilibrium profiles. Most display youthful,
177 staircase profiles with long flat segments separated by very steep tracts, reflecting the stepped
178 remnant plateaus typical of the African landscape (Diniz, 2006). Convex profiles indicative of

179 retrogressive erosion characterize rivers draining the Kalahari plateau or coastal uplifts (Leturmy et
180 al., 2003).

181 Sizeable dams or weirs were built on the Cunene (Matala 1954, Gove 1975, Calueque 1976,
182 Ruacana 1980), Cuanza (Cambambe 1963, Capanda 2004), and Catumbela rivers (Lomaum 1965).
183 Several were abandoned, damaged severely or destroyed during the civil war (1975-2002), and have
184 been rehabilitated recently. Others include the Mabubas Dam on the Dande and the Kiminha Dam
185 on the Bengo. Based on sediment volumes stored in the Cambambe reservoir and measured in the
186 late Eighties, the Cuanza solid load amounts to $620 \pm 100 \cdot 10^3$ ton/a (43% mostly fine sand with
187 some gravel, 44% coarse silt, 13% clay to medium silt; Holisticos, 2012). This corresponds to low
188 sediment yields and erosion rates (4 ± 1 ton/km² a, < 0.002 mm/a).

189

190 2.6. Coastal geomorphology and marine processes

191 The Atlantic coast of Angola is wave-dominated and microtidal (maximum significant wave height
192 ~ 1.75 m; mean tide amplitude ~ 1 m). A powerful southerly swell, originated by persistent stormy
193 winds far away in the Southern Ocean between 40°S and 60°S, causes periodic high-energy wave
194 events called "*calema*" and extensive northward sand transport all along the Atlantic shore of
195 southern Africa. Scarcely affected by either local winds or the Benguela Current, the direction of
196 littoral transport is controlled by the oblique incidence of the wave front to the coast, which
197 determines the asymmetry of deltas and the formation of sand spits in proximity of coastal re-
198 entrants where waves are refracted and sand accumulates offshore (Guilcher et al., 1974).

199 Shelf topography and width exert a major role on sediment transport. The southern Angola oblique-
200 rifted to transform-rifted margin has a narrow coastal plain delimited offshore by a steep continental
201 slope and inland by a high-elevation escarpment with relief up to 1.5 km (Feio, 1981; Lopes et al.,
202 2016). The normal-rifted margin of northern Angola is much wider and has gentler continental
203 slope and hinterland relief without a major escarpment. Such a marked topographic difference
204 reflects both the tectonic structure inherited from Lower Cretaceous rifting and prominent recent

205 uplift of the Bié-Huila dome in southwestern Angola ([Moulin et al., 2005](#); [Pritchard et al., 2009](#);
206 [Guiraud et al., 2010](#)).

207 The continental shelf, 25-30 km-wide to the -150 m isobath offshore of the Cunene mouth, narrows
208 rapidly north of the Baia dos Tigres spit to < 10 km along most of the coast from Tombua to
209 Benguela. Shelf width is reduced to ~ 1 km at the head of submarine canyons offshore of the
210 Curoca and Bero river mouths, which intercept littoral drifting sand thus terminating the ~ 1800
211 km-long Orange littoral cell ([Garzanti et al., 2017](#)), as well as between north of the Coporolo mouth
212 and the Baia Farta spit. The shelf widens again to ≥ 20 km north of Benguela, and widens further
213 north of Luanda to ~ 70 km south of the Congo river mouth.

214

215 **3. Sampling and analytical methods**

216

217 All along the coast of Angola, in June 2015 we have sampled 26 fluvial sand bars exposed at low-
218 water level from all major rivers draining into the Atlantic Ocean, 43 beaches, and 2 dunes.
219 Sediment samples collected offshore of the Congo, Tapado, Balombo, Coporolo and Cunene
220 mouths as well as south of the Cuanza seamount chain and on the Walvis Ridge at water depths
221 between -30 and -3000 m b.s.l. were retrieved from the MARUM repository in Bremen. Offshore
222 samples were collected a few dm below seafloor during Meteor expeditions M6/6, M20/2, and
223 M41/1 ([Wefer et al., 1988](#); [Schulz et al., 1992, 1998](#)), or represent the upper meters of cores drilled
224 during DSDP Leg 40 and ODP Leg 175 ([Fig. 1A](#)). Numerous other river, beach, and dune samples
225 were collected in various years along the coast of Namibia, in the Cunene catchment and southern
226 Moçamedes Desert, in central Angola, in the terminal tract of the Congo River, and in the Republic
227 of Congo. This extensive set of 136 samples overall allowed us to monitor detrital signatures in
228 response to changing lithology and age of source terranes along the Atlantic passive margin of
229 southern Africa, as well as compositional trends controlled not only by latitudinal climatic zonation
230 from hyperarid Namibia to hyperhumid Congo but also by the sharp inland rainfall gradient. Full
231 information on sampling sites is provided in [Appendix Table A1](#) and Google Earth™ file

232 [Angolamargin.kmz](#). Main parameters characterizing river catchments are given in [Appendix Table](#)
233 [A2](#).

234

235 *3.1. Petrography and heavy minerals*

236 Petrographic composition of each sand sample was determined by counting 400 to 450 points in
237 thin section by the Gazzi-Dickinson method ([Ingersoll et al., 1984](#)). Sands are classified by their
238 main components exceeding 10% QFL (e.g., in a feldspatho-quartzose sand $Q > F > 10\%QFL > L$;
239 [Garzanti, 2016](#)). An adjective reflecting the dominant rock-fragment type may be added freely (e.g.,
240 plutoniclastic). Among feldspatho-quartzose sands, feldspar-rich ($Q/F < 2$; plagioclase-rich if
241 plagioclase/K-feldspar > 2 , K-feldspar-rich if K-feldspar/plagioclase > 2) and quartz-rich ($Q/F > 4$)
242 compositions are here formally distinguished ([Fig. 2A](#)). Cross-hatched microcline and untwinned
243 K-feldspar including perthite are called for simplicity microcline and orthoclase throughout the text.
244 Metamorphic grains were classified by protolith composition and metamorphic rank; average rank
245 for each sample was expressed by the Metamorphic Indices MI and MI*, ranging respectively from
246 0 (detritus from sedimentary and volcanic rocks) and 100 (detritus from very low-grade
247 metamorphic rocks) to 500 (detritus from high-grade metamorphic rocks; [Garzanti and Vezzoli](#)
248 [2003](#)).

249 From the widest possible size-range obtained by wet sieving (mostly $< 500 \mu\text{m}$ for clean beach and
250 dune sands and $15\text{-}500 \mu\text{m}$ for positively skewed river sands), heavy minerals were separated by
251 centrifuging in Na-polytungstate (2.90 g/cm^3) and recovered by partial freezing with liquid
252 nitrogen; 200 to 250 transparent heavy minerals were point-counted at suitable regular spacing on
253 grain mounts to obtain real volume percentages ([Galehouse, 1971](#)). On finer-grained offshore
254 samples, heavy-mineral analyses were carried out by Raman point-counting ([Andò et al., 2011](#)) on
255 the $> 5 \mu\text{m}$ (silty clays) or $>15 \mu\text{m}$ fraction (sandy silts and silty sands) obtained by wet sieving.
256 The ZTR index, expressing the "chemical durability" of the suite ([Garzanti, 2017](#)), is the sum of
257 zircon, tourmaline and rutile over total transparent heavy minerals ([Hubert, 1962](#)). The hornblende
258 colour index (HCI) varies from 0 in detritus from greenschist-facies and lowermost amphibolite-

259 facies rocks to 100 in detritus from granulite-facies rocks (Andò et al., 2014). Heavy-mineral
260 concentration, calculated as the volume percentage of total (HMC) and transparent (tHMC) heavy
261 minerals, ranges from extremely poor ($HMC < 0.1$), poor ($0.5 \leq HMC < 1$) and moderately poor (1
262 $\leq HMC < 2$), to rich ($5 \leq HMC < 10$), very rich ($10 \leq HMC < 20$) and extremely rich ($20 \leq HMC <$
263 50); placer sands are defined by $HMC \geq 50$. The Source Rock Density (SRD) index, defined as the
264 weighted average density of extrabasinal terrigenous grains, was used to detect hydraulic-controlled
265 concentration of denser minerals (Garzanti and Andò, 2007). In all analysed samples corrosion
266 features were assessed systematically by three different operators on ~ 25,000 transparent heavy-
267 mineral grains, following the classification of surface textures in Andó et al. (2012). Significant
268 minerals are listed in order of abundance throughout the text. Key compositional parameters are
269 summarized in Table 1. The complete petrographic, heavy-mineral and surface-texture datasets are
270 provided in Appendix Tables A3, A4, and A5.

271

272 3.2. Geochemistry

273 Chemical analyses of 26 river, 23 beach and 2 dune sands were carried out at ACME Laboratories
274 (Vancouver) mostly on a quartered aliquot of the 63-2000 μm class obtained by wet-sieving.
275 Following a lithium metaborate/tetraborate fusion and nitric acid digestion, major oxides and
276 several minor elements were determined by ICP-ES and trace elements by ICP-MS (see Appendix
277 A for specific information on the adopted analytical protocol).

278 To estimate weathering we used several chemical indices, including the Chemical Index of
279 Alteration [$CIA = 100 \cdot Al_2O_3 / (Al_2O_3 + CaO + Na_2O + K_2O)$; Nesbitt and Young, 1982] and the
280 Weathering Index [$WIP = 100 \cdot (CaO/0.7 + 2Na_2O/0.35 + 2K_2O/0.25 + MgO/0.9)$; Parker, 1970],
281 calculated using molecular proportions of mobile alkali and alkaline earth metals corrected for CaO
282 in apatite. Instead of correcting the CIA for CaO in carbonates based on mineralogical data, which
283 may result in significant error (Garzanti and Resentini, 2016), we preferred to use the CIX, a simple
284 modification of the CIA not considering CaO [$CIX = 100 \cdot Al_2O_3 / (Al_2O_3 + Na_2O + K_2O)$;
285 Garzanti et al., 2013]. Weathering intensities were also calculated for each single mobile element by

286 comparing its concentration to that of non-mobile Al in our samples and in the Upper Continental
287 Crust standard: $\alpha^{AlE} = (Al/E)_{sample}/(Al/E)_{UCC}$ (Garzanti et al., 2013, modified after α values of
288 Gaillardet et al. 1999; UCC standard after Taylor and McLennan, 1995; Rudnick and Gao, 2003).
289 Rare earth elements (REE) were normalized to CI carbonaceous chondrites (McDonough and Sun
290 1995). Main weathering indices are shown in Table 1. The complete geochemical dataset is
291 provided in Appendix Table A6.

292

293 3.3. Detrital geochronology

294 Detrital zircons were identified by Automated Phase Mapping with a QEMSCAN[®] WellSite[™]
295 instrument (Vermeesch et al., 2017) on the heavy-mineral separates (mostly < 500 or 15-500 μ m
296 class) of 50 selected samples. U-Pb ages were determined at the London Geochronology Centre
297 using an Agilent 7700x LA-ICP-MS (laser ablation-inductively coupled plasma-mass spectrometry)
298 system, employing a NWR193 Excimer Laser operated at 11 Hz with a 20 μ m spot size and 2.5-3.0
299 J/cm² fluence.

300 No cathodo-luminescence (CL) imaging was done, and the laser spot was always placed "blindly" in
301 the interior of zircon grains. Textural information acquired by CL-imaging can be useful for detrital
302 geochronology. For example, the presence of oscillatory zoning in zircon may indicate an igneous
303 origin, whereas the presence of complex patterns indicates a multi-phase growth history. This
304 knowledge can be very useful for studies whose aim is to identify specific igneous or metamorphic
305 sediment sources. But in other studies the textural information provided by CL-imaging creates
306 additional challenges. If one's main goal is to compare different age distributions with each other
307 like fingerprints, then it is important that all samples are treated in exactly the same way. The
308 availability of textural information makes it more difficult to decide which part of a grain should be
309 analysed: core, rim, or both? In such cases, it is better to always place the ablation spots in the
310 middle of the grain. A "blind dating strategy" is the easiest way to ensure consistency between
311 samples. Data reduction was performed using GLITTER 4.4.2 software (Griffin et al., 2008). We
312 used ²⁰⁶Pb/²³⁸U and ²⁰⁷Pb/²⁰⁶Pb ages for zircons younger and older than 1100 Ma, respectively. No

313 common Pb correction was applied. Grains with $> +5/-15\%$ age discordance were discarded, and
314 2967 concordant ages were obtained overall. The complete geochronological dataset is provided in
315 [Appendix B](#).

316

317 **4. Detrital signatures of river sands**

318

319 *4.1. Cunene River*

320 The Cunene River, sourced in Bié-Huila highlands, flows across the fossil Kalahari dunefield and
321 once drained towards the Etosha Pan of northern Namibia as the Okavango River drains today into
322 the Makgadikgadi Pan of Botswana ([Fig. 1B](#)). Captured only recently by a coastal stream ([Haddon
323 and McCarthy, 2005](#)), the lower Cunene cuts deeply across the rifted-margin shoulder where the
324 Paleoproterozoic Epupa basement, the Mesoproterozoic Cunene anorthosite, and finally the Kaoko
325 northern branch of the Damara orogen are exposed ([Becker et al., 2006](#)). In the terminal tract the
326 river marks the northern limit of the Cunene coastal dunefield and eventually reaches the Atlantic
327 Ocean south of the Moçamedes Desert. Sand mineralogy changes repeatedly downstream ([Fig. 2E](#)),
328 which implies higher erosion rates in the lower tract than in the upper to middle course. At Ruacana
329 Falls at the end of the middle course, sand composition is quartz-rich feldspatho-quartzose with
330 dominant monocrystalline quartz, orthoclase $>$ plagioclase \geq microcline and a poor amphibole-
331 epidote heavy-mineral suite. Quartz grains recycled from Kalahari dunes (e.g., supplied by the
332 Caculuar tributary and Mucope subtributary) are about twice as abundant as first-cycle detritus
333 from basement rocks. At Epupa Falls, sand is plagioclase-rich feldspatho-quartzose and markedly
334 enriched in polycrystalline quartz, clinopyroxene and hypersthene, revealing dominant supply from
335 crystalline basement including the Cunene anorthosite and associated mafic rocks (e.g., drained by
336 the Omuhongo tributary; [Fig. 3](#)). Sand composition changes again in the terminal tract because of
337 extensive mixing with dune sand windblown from the coast and mostly derived ultralong-distance
338 from the Orange River ([Garzanti et al., 2014c](#)). In Cunene sand at Epupa Falls, zircon-age spectra

339 are unimodal with broad slightly asymmetric Mesoproterozoic peak centered at ~ 1.4 Ga and most
340 ages ranging between 1 and 2 Ga.

341

342 *4.2. Rivers of southern Angola (Curoca to Cangala)*

343 The Curoca River marks the northern boundary of the Moçamedes coastal dunefield, and in its
344 terminal tract — as in the terminal tract of the Cunene — mixing with windblown sand is so
345 extensive that composition becomes indistinguishable from littoral sand largely derived ultralong-
346 distance from the Orange River (Fig. 2E). A fossil dune collected upstream of the mouth contains
347 more K-feldspar and less heavy minerals including mainly staurolite and tourmaline, and no
348 pyroxene. Zircon-age spectra in both Curoca mouth and fossil-dune samples display a polymodal
349 distribution undistinguishable from that of Moçamedes desert sand, with the Damaran (Cambrian-
350 late Neoproterozoic) and Namaqua (late Mesoproterozoic) clusters typical of Orange sand
351 associated with subordinate Eburnean (late Paleoproterozoic) ages.

352 All rivers to the north drain the Angola Block, consisting of Eburnean granitoid rocks and
353 subordinate metasediments with an overlying volcano-sedimentary sequence and younger bimodal
354 magmatic products exposed in the "coastal polyorogenic belt" (De Carvalho et al., 2000; Pereira et
355 al., 2011; Ernst et al., 2013). Sand composition ranges from quartzo-feldspathic to feldspar-rich
356 feldspatho-quartzose with plagioclase > orthoclase > microcline (Fig. 2B). Granitoid rock fragments
357 and polycrystalline quartz are very common, associated with a few metamorphic (quartz-mica,
358 quartz-epidote, low-grade metabasite, rare schist, and amphibolite) and rare sedimentary (sparite
359 and quartz siltstone) lithics. Muscovite and biotite occur sporadically. Moderately rich heavy-
360 mineral suites range from amphibole-epidote in Bero and Giraul sands in the south, to epidote-
361 amphibole in the north (Fig. 3). Clinopyroxene and hypersthene in Bero sand may be supplied by
362 the Cunene intrusive complex drained in the headwaters. Provenance from chiefly granitoid rocks is
363 reflected by very strong depletion in Mg and enrichment in K, Rb, and especially Ba relative to the
364 UCC standard ($\alpha^{Al}Mg$ 8 ± 2 , $\alpha^{Al}K$ 0.6 ± 0.1 , $\alpha^{Al}Ba$ 0.4 ± 0.0 for Giraul, Bentiaba, Inamangando,
365 and Cangala sands, Bero sand being slightly more mafic) (Fig. 4E). The Eu anomaly is only slightly

366 negative ($Eu^* 0.88 \pm 0.06$). Bero sand yielded few Mesoproterozoic zircon ages, whereas spectra in
367 Giraul and Inamangando sands are unimodal with peak centered around 2 Ga (Fig. 5).

368

369 *4.3. Rivers of southern central Angola (Coporolo to Catumbela)*

370 Composition ranges from quartzo-feldspathic (Cavaco sand) to feldspatho-quartzose (Catumbela
371 sand) with plagioclase \geq microcline \geq orthoclase (Fig. 2B). Common granitoid and metaigneous
372 rock fragments occur. Biotite prevails over muscovite. Heavy-mineral suites are moderately poor
373 and epidote-dominated with amphibole in Coporolo and Cavaco sands. A richer epidote-amphibole-
374 clinopyroxene suite with minor zircon and olivine characterizes Catumbela sand. Mafic igneous
375 detritus in the Catumbela catchment is supplied by the lowermost-course Capilongo tributary
376 draining gabbroic rocks of possible Mesoproterozoic age intruded in the "coastal polyorogenic belt"
377 (Ernst et al., 2013) (Fig. 3). Geochemical signatures are similar to river sands of southern Angola
378 but somewhat depleted in alkali and alkaline-earth metals. Zircon-age spectra are unimodal with
379 sharp peak centered around 2 Ga (Fig. 5).

380

381 *4.4. Rivers of northern central Angola (Culango to Longa)*

382 The quartz/feldspar (Q/F) ratio increases progressively northward, composition ranging from
383 feldspar-rich feldspatho-quartzose in the south to quartz-rich feldspato-quartzose in Longa sand
384 (Fig. 2B,D). K-feldspar mainly prevails over plagioclase. Granitoid and a few metamorphic rock
385 fragments occur. Limestone grains with marine fossils (echinoderm plates), and a few felsic
386 volcanic grains are present in Longa sand, where monocrystalline quartz locally showing rounded
387 outline or abraded overgrowths is dominant. Quicombo sand includes a few carbonate grains.
388 Biotite prevails over muscovite. Poor to moderately rich, epidote-dominated to epidote \approx amphibole
389 suites include zircon. Cambongo sand contains a few mafic volcanic and clinopyroxene grains.
390 Silica concentration is highest in Longa sand, with consequent dilution of most other elements.
391 Zircon-age spectra are unimodal with sharp peak centered around 2 Ga (Fig. 5). Late Neoproterozoic to
392 early Paleoproterozoic ages are represented in Queve sand. A few zircon grains with Early

393 Cretaceous age occur in Quicombo sand, indicating additional supply from local syn-rift volcanic
394 rocks.

395

396 *4.5. Rivers of northern Angola (Cuanza to Luculu)*

397 Composition is mainly quartz-rich feldspatho-quartzose with orthoclase + microcline \geq plagioclase.
398 Lifune, Mebridege, and Luculu sands are quartzose (Fig. 2B,D). Monocrystalline quartz is
399 dominant in most samples, with well rounded grains most common in Cuanza, Bengo, and Luculu
400 sands. Granitoid and metasedimentary rock fragments occur. Micas (muscovite > biotite) and
401 medium-rank schist and quartz-mica lithics are most common in Dande sand. Quartz-siltstone to
402 metasilstone/metasandstone rock fragments are associated with quartz displaying abraded
403 overgrowths or subrounded outline in Mebridege sand. Micrite to sparite grains are most common
404 in Bengo sand and occur in Cuanza and Dande sand, where they may contain Cenozoic planktonic
405 foraminifera. Very poor to rich heavy-mineral suites include mainly amphibole, epidote, garnet, and
406 kyanite (ZTR 6 ± 5). River sands are progressively enriched northward in SiO₂, with stronger and
407 stronger dilution of most other elements. Zircon-age spectra display a peak around 2 Ga and ages
408 ranging mostly between ~ 0.5 and ~ 3 Ga. Loge sand yielded a sharper unimodal Eburnean peak.
409 Luculu sand yielded $\sim 10\%$ of zircon grains with mid-Permian to mid-Triassic ages, which occur
410 rarely in Cuanza and Bengo sands as well (Fig. 5).

411

412 *4.6. Rivers of the Congo*

413 Sands are dominated by subrounded to well rounded monocrystalline quartz; SiO₂ reaches above
414 98%. The Congo River carries to the mouth a few K-feldspar (orthoclase > microcline) and
415 quartzose siltstone grains (Fig. 2D). The very poor heavy-mineral suite includes zircon, tourmaline,
416 epidote, staurolite, kyanite, and minor rutile, amphibole, garnet, sillimanite, and andalusite (ZTR 51
417 ± 16). Sands in the Republic of Congo to the north also contain a few quartzose siltstone grains and
418 very poor suites dominated by zircon and tourmaline, with rutile, staurolite, kyanite, epidote, and
419 garnet (ZTR 74 ± 4).

420

421 **5. Detrital signatures of coastal and offshore sediments**

422

423 *5.1. The Orange littoral cell*

424 Dune and beach sands between the Cunene mouth and south of Namibe are feldspatho-quartzose
425 with lathwork volcanic rock fragments and moderately poor to extremely rich garnet-
426 clinopyroxene-epidote-amphibole suites including hypersthene and staurolite. Higher Fe, Mg, Ca,
427 Ti, P, Mn, Sc, Y, HREE, V, Cr, Co, and Ni than most other Angolan sands reflect the presence of
428 common volcanic rocks fragments, pyroxene, and garnet. Zircon-age spectra show broad
429 Cambrian/late Neoproterozoic, late Mesoproterozoic, and late Paleoproterozoic peaks, with a few
430 younger (as young as Early Cretaceous) and older ages (as old as late Neoproterozoic; [Fig. 5](#)). Such
431 petrographic, mineralogical, geochemical, and geochronological fingerprints resemble closely those
432 of the Skeleton Coast Erg, pointing to ultralong-distance littoral transport from the Orange mouth
433 with subordinate contributions from the Cunene River and ephemeral rivers draining the Damara
434 orogen in Namibia ([Garzanti et al., 2014c, 2017](#)).

435 Sediments collected on the Walvis Ridge and offshore of the Cunene mouth are dominated by
436 planktonic foraminifers and green glaucony grains, respectively ([Fig. 6](#)). The sandy-silt to silty-sand
437 feldspatho-quartzose terrigenous fraction contains mafic volcanic rock fragments and a rich,
438 clinopyroxene-dominated heavy-mineral suite with subordinate epidote, amphibole and garnet,
439 indicating provenance mostly from the Orange mouth. Deep-sea silty clays offshore of Baia dos
440 Tigres contain diatom tests and siliceous sponge spicules with minor radiolaria, silicoflagellates,
441 glaucony, plant debris, and organic matter. The subordinate feldspar-rich feldspatho-quartzose
442 terrigenous fraction yielded either the same clinopyroxene-dominated suite as Walvis Ridge and
443 offshore Cunene samples, or a clinopyroxene-epidote-amphibole-garnet suite comparing closely
444 with beaches and dunes of the southernmost Moçamedes Desert, thus revealing additional
445 contribution from the Cunene River.

446

447 5.2. From Namibe beach to Coporolo mouth

448 Beaches in the oblique Namibe rifted-margin segment bear the same feldspar-rich quartzo-
449 feldspathic composition and epidote-amphibole suites of local river sand, reflecting dominant
450 supply from largely granitoid rocks exposed in adjacent Angolan highlands. Chemical composition
451 is also similar, with higher Fe, Mg, Ti, P, Mn, Sc, U, V, Cr, Co, Ni, and Cu in Namibe beach
452 reflecting the abundance of heavy minerals including pyroxenes fed from the Bero River. Zircon-
453 age spectra are mostly unimodal with peak centered around 2 Ga (Fig. 5). Mesoproterozoic ages are
454 significant in Namibe beach. In Equimina beach, 22% of zircon grains yielded ages between 84 and
455 104 Ma, indicating provenance from mid-Cretaceous trachytic lavas exposed locally (Neto, 1960).
456 Supply from Meso-Cenozoic strata of the Namibe basin is revealed in Chapeu Armado beach by
457 microcline \approx plagioclase, occurrence of a few carbonate and mafic volcanic lithics, and poor suites
458 with common durable heavy minerals (ZTR 24). Carbonate, quartz-siltstone, and basaltic grains
459 occur also at Mariquita and Bentiaba beaches. Beach sand is derived dominantly from basement
460 rocks in the north where the Namibe basin tapers out, as documented by abundant feldspar,
461 metabasite rock fragments and up to very rich epidote-amphibole suites, although a few carbonate
462 rock fragments also occur. Presence of carbonate grains is reflected in locally high CaO and LOI
463 (e.g., Baia Binga beach). Local differences in sand composition along this coastal stretch, where the
464 shelf width is reduced to only some kilometers thus limiting the possibility of longshore sand drift,
465 indicate that pocket beaches are largely independent sediment systems fed from the associated
466 stream and coastal cliffs.

467

468 5.3. The Coporolo and Catumbela littoral cells

469 The perennial Coporolo and Catumbela Rivers represent the major source of beach sand along the
470 Benguela transform rifted-margin sement (Dinis et al., 2016). Beaches north of the Coporolo mouth
471 display the same feldspar-rich feldspatho-quartzose signature with moderately poor to moderately
472 rich epidote-dominated suites with amphibole as Coporolo river sand. A few carbonate and quartz-
473 bearing siltstone and sandstone grains are however added along the coast, and quartz increases in

474 the Baia Farta sand spit indicating additional supply from local streams and coastal cliffs. The shelf
475 is very narrow offshore of the Baia Farta promontory, where a submarine canyon may capture
476 longshore-drifting sand thus hampering northward elongation of the spit. The silty clay collected in
477 the deep sea is quartzo-feldspathic with the same epidote-dominated suite as beach sand in the
478 Coporolo cell. North of Benguela the continental shelf widens again to ~ 25 km, and northward
479 littoral drift is renewed. Beach sand between the Cavaco and Catumbela mouths is nearly as
480 feldspar-rich as Cavaco river sand. Sand with the same feldspatho-quartzose signature as Catumbela
481 river sand with a rich epidote-dominated suite including amphibole and clinopyroxene is carried
482 northward of the strongly asymmetric Catumbela delta to build prograding beach ridges culminating
483 in the 5 km-long Lobito spit. Zircon-age spectra are unimodal with sharp Eburnean peak. Beach
484 sand north of Coporolo mouth yielded a few zircon grains with younger ages, one of which mid-
485 Cretaceous (Fig. 5).

486

487 *5.4. From Lobito to Cabo Ledo*

488 Beach sand along the oblique-rifted Cuanza margin segment is invariably feldspatho-quartzose, and
489 thus richer in quartz than local river sands suggesting either additional supply from coastal cliffs or
490 northward longshore drift from the Catumbela mouth. The moderately poor to rich epidote-
491 amphibole-clinopyroxene heavy-mineral suite may include rare olivine. Clinopyroxene, notably
492 more abundant than in river sands, is derived from volcanic rocks best exposed in the Sumbe area
493 (Marzoli et al., 1999) and/or longshore from the Catumbela mouth. Local enrichment in heavy
494 minerals is revealed by higher Fe, Mg, Ti, Sc, REE, Zr, Hf, V, Nb, Ta, and Cr than in river sands.
495 Beach sand at Cabo Ledo, containing carbonate grains as reflected by locally high CaO and LOI,
496 yielded a unimodal zircon-age spectrum with sharp Eburnean peak as beach sand at Porto Amboim.
497 A few mid-Cretaceous ages confirm additional detritus from syn-rift igneous rocks along the coast.
498 Silty sands collected offshore of the Tapado mouth include mostly benthic foraminifers, which
499 become dominant beyond the shelfbreak (Fig. 6). The quartzo-feldspathic to feldspar-rich
500 feldspatho-quartzose terrigenous fraction yielded a rich, epidote-dominated heavy-mineral suite

501 with subordinate amphibole. Feldspar is more abundant than in beach sand and clinopyroxene
502 minor, indicating provenance directly from the Balombo and/or Tapado river mouths rather than
503 from longshore drifting sand. The silty clay collected in the deep sea south of the Cuanza seamounts
504 (Fig. 1A), dominated by planktonic and subordinately benthic foraminifers, yielded a similar
505 epidote-amphibole suite with no clinopyroxene.

506

507 *5.5. The Cuanza littoral cell*

508 Beach sand from the Cuanza mouth to Ilha de Luanda is quartz-rich feldspatho-quartzose with
509 plagioclase \approx microcline \approx orthoclase. The poor to moderately rich epidote-amphibole-
510 clinopyroxene heavy-mineral suite includes minor garnet, kyanite, staurolite, andalusite, and rare
511 olivine. Feldspar is more common and ZTR indices lower than in Cuanza river sand. Clinopyroxene
512 increases along the littoral cell where polycrystalline quartz is also more abundant, suggesting
513 mixing with northward drifting sand along the broad littoral zone. Chemical data confirm
514 significant differences between Cuanza river sand and the Mussulo and Ilha de Luanda spits, which
515 have less Si, higher alkali and alkaline earth metals (and hence systematically lower weathering
516 indices), lower Zr and Hf, and much steeper REE patterns with no Eu anomaly.

517

518 *5.6. From Luanda to the Congo*

519 Beach sands along the orthogonal to oblique-rifted Congo-margin segment vary from mainly
520 quartz-rich feldspatho-quartzose in the south to quartzose north of N'Zeto, with a progressive
521 northward decrease in feldspar (orthoclase \geq plagioclase \geq microcline). The heavy-mineral
522 concentration of epidote-amphibole-garnet-kyanite suites tends to decrease from mainly rich in the
523 south to moderately rich and finally poor in the north, where staurolite becomes significant and
524 zircon, tourmaline, and rutile increase (ZTR up to 13). Beach sand in the Republic of Congo
525 consists almost entirely of commonly rounded monocrystalline quartz grains displaying deep etch
526 pits, with only a few microcline grains and a very poor suite with zircon, rutile, tourmaline,

527 staurolite, kyanite, and rare andalusite (ZTR 66). SiO₂ increases progressively northward, reflecting
528 increasing and finally overwhelming quartz dilution.

529 Zircon grains in Ambriz sand yielded a sharp unimodal Eburnean peak along with 5% of Upper
530 Permian ages. The Eburnean peak becomes broader north of the Loge mouth, where Permian ages
531 also occur. The zircon-age spectrum is multimodal in beach sand south of the Congo mouth with
532 Pan-African (0.5-0.7 Ga), Namaqua (~ 1 Ga), Eburnean (~ 2 Ga) and Neoproterozoic (2.5-2.7 Ga)
533 clusters. A similar spectrum characterizes Congo estuary sand (Fig. 5).

534 Fine-grained quartzose sand collected on the inner shelf just south of the Congo mouth includes
535 common brown goethite ooids (Fig. 6), a few plagioclase and K-feldspar grains, and a moderately
536 rich heavy-mineral suite including epidote, garnet, amphibole, zircon, rutile, kyanite, minor
537 andalusite, tourmaline, apatite and staurolite, indicating provenance from northward longshore drift
538 rather than from the Congo River. Very fine-grained shelfal sediment north of the Congo mouth
539 consists of brown pellets and foraminifers (Fig. 6).

540

541 **6. Provenance control on sand composition**

542

543 The geology of Angola includes two main domains, the Paleoproterozoic Angola Block with the
544 prominent Mesoproterozoic Cunene intrusive complex in the south, and the Pan-African West
545 Congo Belt in the north. The dual time structure of source rocks is best reflected by U-Pb ages of
546 detrital zircons, characterized by unimodal spectra with sharp Eburnean peak at ~ 2 Ga from north
547 of Namibe to south of the Cuanza mouth, and by multimodal distributions in northern Angola with
548 Cambrian-Neoproterozoic, latest Mesoproterozoic, and Neoproterozoic clusters best defined in Congo
549 estuary sand. The Mesoproterozoic peak with Eburnean shoulder displayed by detrital zircons in
550 Cunene sand at Epupa Falls reflects contribution from the Cunene intrusive complex, documented
551 as far north as Namibe beach fed by the Bero River. To the south, sharply different detrital modes,
552 heavy-mineral suites, and zircon-age clusters characterize dunes and beaches of the Moçamedes
553 Desert, documenting ultralong-distance provenance mainly from the Orange River.

554 Detrital modes, heavy minerals, and concentration of chemical elements reflect local source-rock
555 distribution as faithfully as zircon age-spectra only in dry southern Angola north of the Orange
556 littoral cell, where abundance of feldspar and granitoid rock fragments among main framework
557 grains and of epidote and amphibole among transparent heavy minerals indicate the Angola Block
558 as the dominant source ("basement uplift subprovenance" of [Dickinson, 1985](#); "dissected
559 continental block subprovenance" of [Garzanti, 2016](#)). Metamorphic rock fragments of mainly
560 medium to high rank ($MI^* 312 \pm 50$), epidote/hornblende ratio mostly > 1 and dominantly
561 blue/green amphiboles (HCI mostly ≤ 10) point to mainly greenschist-facies to lower amphibolite-
562 facies source rocks. Scarcity of garnet, lack of kyanite, andalusite and sillimanite, and only sporadic
563 occurrence of metasedimentary rock fragments reflect predominance of granitoid and metagranitoid
564 source rocks, as confirmed by geochemical signatures. Metabasite rock fragments and higher
565 heavy-mineral concentration indicate more significant supply from mafic greenschist-facies rocks in
566 beach sand from Lucira to the Coporolo mouth. Meso-Cenozoic strata and intercalated lavas of the
567 Namibe basin supply minor amounts of sedimentary and volcanic detritus to pocket beaches locally.
568 Lower Cretaceous mafic rocks contribute a few basaltic fragments and clinopyroxene, whereas
569 more felsic younger products provide zircon grains with mid-Cretaceous ages (e.g., Equimina
570 beach).

571 Provenance is dominantly from the Angola Block also in central Angola, as reflected by the same
572 unimodal Eburnean zircon-age spectra and epidote-amphibole suites. Sand is however characterized
573 by higher quartz/feldspar (Q/F; [Fig. 2B](#)) and lower plagioclase/feldspar (P/F) ratios, chiefly
574 reflecting increasing weathering conditions (as discussed in [section 8](#) below) rather than different
575 provenance. Mafic igneous detritus documented by gabbroic rock fragments and clinopyroxene is
576 most significant in Catumbela sands. In beaches from the Catumbela mouth to Luanda,
577 clinopyroxene is invariably much more common than in local rivers ([Fig. 3](#)), indicating coastal
578 dispersal of mafic igneous detritus derived from either the Catumbela mouth or Cretaceous syn-rift
579 tholeiitic to post-rift alkaline basalts exposed at the landward termination of the Cuanza volcanic
580 line ("anorogenic volcanic provenance" of [Garzanti, 2016](#)). Heavy-mineral concentration tends to

581 be higher in beach sands than in river sands, but differences in heavy-mineral suites cannot be
582 ascribed to hydrodynamic processes, because clinopyroxene is only slightly less dense than epidote
583 and none of our samples show markedly anomalous SRD indices. Offshore samples are pyroxene-
584 poor, which argues against longshore drift from the Catumbela mouth.

585 In northern Angola, the provenance signal is dimmed further by progressively stronger weathering
586 effects. Mainly low-grade to lower medium-grade metasedimentary detritus from the West Congo
587 Belt is documented by medium-rank schist and quartz-mica lithics, muscovite, biotite, and common
588 garnet and kyanite; staurolite increases in relative abundance from northernmost Angola to the
589 Congo. Quartz and durable heavy minerals recycled largely from Meso-Cenozoic coastal basins, but
590 also partly from Neoproterozoic successions, are common in sands of rivers cutting across the
591 Cuanza basin in their terminal tract, including the Longa, Bengo, and especially the Cuanza (as
592 discussed in [subection 7.3](#) below).

593 In the Congo, the provenance signal is largely obliterated by extreme weathering conditions, which
594 coupled with extensive recycling produce pure quartzose sand with high ZTR ([Figs. 2D, 4A,C](#);
595 "craton interior provenance" of [Dickinson, 1985](#); "recycled clastic provenance" of [Garzanti, 2016](#)).

596

597 **7. Physical processes controlling sand composition**

598

599 In this section we discuss the compositional modifications associated with physical processes in the
600 depositional environment, namely hydraulic sorting, mechanical breakdown, and recycling.

601

602 *7.1. Hydraulic sorting*

603 Sorting of detrital grains based on their size, density, and shape occurs in all three stages of the
604 sedimentary cycle, during erosion by selective entrainment, during transport by vertical partitioning
605 of particles carried in suspension, and eventually during deposition according to their settling
606 velocity ([Komar 2007](#)). Erosive processes may lead to extreme modifications of sand composition,
607 as documented by common formation of garnet and magnetite beach placers in southern Angola

608 (Garzanti et al., 2017). These lag deposits are generated during storm events when tectosilicate
609 grains — having smaller pivoting angles and projecting higher above the bed than smaller settling-
610 equivalent dense minerals and thus experiencing greater flow velocities and drag forces (Komar and
611 Li 1988) — are selectively removed offshore. These phenomena are readily revealed by the
612 anomalous concentration of dense and ultradense minerals, as reflected by bulk-sediment grain
613 densities notably higher or lower than the estimated weighted average density of source rocks
614 (ranging from $\sim 2.65 \text{ g/cm}^3$ for quartzarenites, quartzites and granites to $2.8\text{-}3.0 \text{ g/cm}^3$ for mafic
615 igneous and metamorphic rocks; Garzanti et al., 2009). Among our samples, anomalously high
616 grain densities are estimated for the Inamangando outer berm and for the Equimina and Cabo Ledo
617 beaches (SRD between 2.77 and 2.94 g/cm^3), all notably enriched in opaque Fe-Ti-Cr oxides and
618 zircon relative to adjacent beaches derived from similar source rocks. Placer lags rich in Fe-Ti-Cr
619 oxides and zircon were sampled at the Catumbela mouth and sporadically observed in northern
620 Angolan beaches (e.g., garnet-rich Ambriz sample). Such strong selective-entrainment effects are
621 revealed readily by the anomalous concentration of elements hosted preferentially in ultradense
622 minerals (e.g., Ti, REE, Zr, Hf, Nb, Ta, Cr), as well as by highly negative Eu anomalies (Garzanti et
623 al., 2010). Zr and Hf concentrations reach 2474 and 49 ppm in Equimina beach ($\text{Eu}/\text{Eu}^* 0.45$),
624 reflecting zircon abundance. The Eu anomaly is most strongly negative in the Inamangando outer
625 berm ($\text{Eu}/\text{Eu}^* 0.28$). Such hydrodynamic effects were not observed in river sands.

626

627 *7.2. Mechanical breakdown*

628 Mechanical wear is scarcely effective during even long-distance fluvial and longshore transport in
629 aqueous media (Russell, 1937; Shukri, 1950; Breyer and Bart, 1978). Although feldspar is softer
630 than quartz and has good cleavage, no significant decrease in the Q/F ratio or in feldspar size is
631 observed over a distance of eighteen hundred kilometers from the Orange mouth to southern
632 Angola. We thus concluded that mechanical breakdown cannot modify the composition of passive-
633 margin sand significantly (Garzanti et al., 2015).

634 The Q/F ratio, however, decreases in offshore silts. In northern Namibia to southern Angola,
635 tectosilicate modes are very close in fine to medium Skeleton Coast and Moçamedes desert sands
636 (Q 78 ± 4 K 8 ± 2 P 14 ± 3 and Q 74 ± 4 K 10 ± 3 P 16 ± 3) and in very fine shelfal sand offshore of the
637 Cunene mouth (Q 77 ± 1 K 8 ± 0 P 15 ± 1), but do change in deep-sea silt from the Walvis Ridge (Q 57
638 K 12 P 31). Cleavable feldspar thus concentrates in the silt fraction. In central Angola, feldspar
639 increases with grain size from Coporolo medium sand (Q 63 ± 2 K 21 ± 3 P 16 ± 2) to deep-sea silty
640 clay (Q 39 K 25 P 36), but only slightly from fine to medium Balombo and Tapado river sands (Q
641 55 ± 2 K 27 ± 0 P 18 ± 3) to very fine sand offshore (Q 49 ± 6 K 29 ± 0 P 23 ± 5). Fine sand offshore of the
642 Congo mouth is richer in feldspar than beach sand, but chiefly because of longshore transport from
643 the south.

644 North of Namibe, where eolian dunefields are lacking, most silicates are observed to remain mainly
645 subangular in beach sand, indicating that mechanical wear during fluvial and longshore transport
646 has a limited influence on both composition and texture of sand-sized sediment.

647

648 *7.3. Recycling*

649 Recycling represents the most elusive problem by far in provenance analysis of modern sands
650 (Blatt, 1967). The presence of recycled detritus is qualitatively revealed by the occurrence of quartz
651 grains with abraded overgrowths or well-rounded outline and of shale, siltstone or sandstone rock
652 fragments, together with a series of parameters, including high Q/F and ZTR indices and low P/F
653 and heavy-mineral concentration (Fig. 4C). All of these features characterize river and beach sands
654 in northern Angola, and become extreme from northernmost Angola to the Congo (Fig. 7F,G).

655 Any attempt to quantify precisely the amount of recycled versus first-cycle detritus is however
656 unrobust, because mineralogical parameters indicative of recycling are generally affected by
657 weathering as well (Garzanti, 2017). In southern Angola north of the Orange littoral cell, where
658 climate is arid and weathering negligible, detritus recycled locally from Meso-Cenozoic strata of the
659 Namibe basin is estimated by forward compositional modelling (Garzanti et al., 2012b) to be $\geq 15\%$

660 in the Mariquita and Chapeu Armado pocket beaches, and very minor elsewhere. In central Angola,
661 beach sands tend to be richer in quartz than river sands, suggesting that detritus recycled from
662 coastal outcrops of sedimentary strata may reach up to ~ 30%, a percentage estimated tentatively to
663 increase to ~ 40% for Longa sand and to \geq 50% for Cuanza sand. The recycled component,
664 estimated to be ~ 30% for Bengo sand and ~ 15% for Dande sand, is prominent in beaches of
665 northern Angola and overwhelming in the Congo, but an accurate estimate cannot be made because
666 of strong superposed weathering effects. Recycled monocrystalline quartz is dominant in rivers
667 draining Kalahari fossil dunes including middle Cunene tributaries as well as the uppermost
668 Cuanza, the endorheic Okavango, the Cuando, and the upper Zambezi draining into the Indian
669 Ocean. Detritus recycled from quartzose sandstones and metasandstones of the Neoproterozoic
670 West Congo Belt occurs in all major rivers from the Cuanza to the Congo and it is most evident in
671 Mebridgege sand. Quartzose detritus recycled from siliciclastic strata of the Meso-Cenozoic Cuanza
672 basin is most significant in Longa, Cuanza and Bengo sands, whereas Luculu sand and beach sand
673 in northernmost Angola are recycled from Lower Congo basin strata (Fig. 1D).

674 Rivers of southern and central Angola carry very few garnet grains (< 3% of transparent heavy
675 minerals), reflecting virtual lack of garnet in granitoid rocks of the Angola Block. Garnet represents
676 8 %tHM in Chapeau Armado beach, indicating recycling of sandstones in the coastal Namibe basin.
677 In northern Angola, garnet is significant (5-13 %tHM) in rivers between the Cuanza and the Onzo
678 draining the northern part of the Cuanza basin, and it is common in Luculu sand draining the Lower
679 Congo basin (33 %tHM) and in all beach sands between the Dande and Congo mouths (20 \pm 10
680 %tHM). The areal distribution of garnet along the Angolan rifted margin thus reveals that it is
681 mostly recycled from siliciclastic strata of Meso-Cenozoic basins — ultimately derived from
682 metasedimentary rocks of the West Congo Belt in northern Angola and of the Kaoko Belt in
683 southern Angola — rather than first-cycle from basement rocks.

684 Interpretation of chemical data also suffers from the superposed effects of weathering and recycling.
685 In river and beach sands from southern Angola to the Congo, Si increases markedly (corr. coeff. $r =$

686 0.88, sign. lev. 0.1%), whereas most other elements decrease, non-mobile Al most evidently ($r = -$
687 0.92). The sharp northward decrease in aluminum results from the superposed effects of selective
688 feldspar breakdown by paleoweathering during previous sedimentary cycle(s), by intrastratal
689 dissolution between successive cycle(s), and only last by hydrolisis in the present subequatorial
690 climate. Because of such strong multiphase depletion in Al, chemical-weathering proxies including
691 the CIA, PIA, CIX, and α^{Al} may give unexpectedly low values in sand generated in hyper-humid
692 equatorial climates (e.g., Congo sand; [Table 1](#)). The Eu anomaly is more strongly negative in sands
693 containing more recycled quartz as suggested independently by detrital modes and textural features
694 ($Eu/Eu^* 0.45-0.50$ for Cuanza and Longa sands, $Eu/Eu^* 0.50-0.70$ for Mebridege and Congo river
695 sands and all beaches north of $6^{\circ}30'S$). In order to single out the recycling effect we may calculate
696 the ratio between the CIX (a truer indicator of weathering unaffected by the addition of biogenic,
697 pedogenic, or extrabasinal carbonate grains) and the WIP (strongly influenced by quartz dilution;
698 [Garzanti et al., 2013](#)). The CIX/WIP ratio does increase northward ([Fig. 4](#)), but the correlation ($r =$
699 0.49) is barely significant.

700

701 **8. Weathering control and relative durability of detrital minerals**

702

703 The effect of weathering on Angolan continental-margin sediments, associated with both latitudinal
704 climatic zonation and strong inland rainfall gradients, has been discussed in detail in a previous
705 study focusing on clay minerals and sediment geochemistry ([Dinis et al., 2017](#)). Here we focus
706 instead on sand composition, and discuss specifically the effects of subequatorial weathering on
707 detrital modes and surface textures of quartz, feldspar, and heavy minerals ([Fig. 7](#)).

708

709 *8.1. Main framework grains*

710 In order to consider a homogeneous population of sediment samples, we have selected 23 river
711 sands lacking clues of polycyclic origin (e.g., quartz grains with abraded overgrowths or well
712 rounded outline, arenaceous rock fragments) and thus held to be derived dominantly first-cycle

713 from crystalline basement. In this sample set, the quartz/feldspar ratio is ≥ 1 in southern Angola, ~ 2
714 in central Angola, and ~ 7 in northern Angola (Figs. 2B and 4A). The plagioclase/K-feldspar ratio is
715 > 1 in southern Angola and ≤ 1 in central and northern Angola; cross-hatched microcline represents
716 $\sim 30\%$ of K-feldspar grains in southern Angola and $\sim 45\%$ in central and northern Angola (Fig. 4B).
717 Because source rocks are broadly homogeneous in southern and central Angola, these trends are
718 interpreted as chiefly weathering-controlled. Plagioclase disappears in the Congo, where microcline
719 may be the only detrital feldspar present. Quartz is thus inferred to resist chemical dissolution far
720 better than feldspars, K-feldspar somewhat better than plagioclase, and microcline somewhat better
721 than orthoclase, confirming earlier observations (e.g., Blatt, 1967; James et al., 1981; McBride,
722 1985).

723 The same set of selected river sands were photographed, and the degree of weathering evaluated on
724 both thin sections and microphotographs by two different operators, based on the presence and
725 depth of etch pits, and on the degree of replacement of quartz and feldspar lattice by clay plasma
726 (Fig. 7). We subdivided the samples in groups with similar frequency and extent of alteration
727 features, and then ranked them — from the apparently least weathered to the apparently most
728 weathered — in five independent trials. A trend of increasing weathering with decreasing southern
729 latitude was obtained in all trials, with a median of Spearman rank correlation coefficients of $0.76 \pm$
730 0.25 (sign. lev. 0.1%); the rather large standard deviation reflects the considerable degree of
731 subjectivity involved in qualitative judgments.

732 The compositional change associated with the strong W/E climatic gradient was investigated in the
733 Catumbela catchment, where we studied sands carried by three tributaries draining progressively
734 more humid areas farther from the coast. The Q/F ratio is observed to increase rapidly and the P/F
735 ratio to decrease in sands derived from more humid areas of the interior. Highest Q/F and lowest
736 P/F ratios characterize trunk-river sand derived largely from humid headwater highlands (Fig. 2C).
737 These samples, although limited in number, display a much sharper weathering-controlled
738 compositional trend than the one associated with latitudinal climatic zonation.

739

740 *8.2. Heavy minerals*

741 Pyroxene and amphibole grains are mostly corroded, and show a higher degree of weathering
742 relative to all other heavy minerals in most samples. Epidote and staurolite are commonly corroded
743 and occasionally deeply etched or skeletal. Andalusite, garnet, kyanite, titanite, and apatite are also
744 corroded, but etched or deeply etched only locally. Tourmaline is corroded more commonly than
745 rutile, and rutile more commonly than zircon. Etch pits do not appear to increase remarkably in
746 frequency or depth with decreasing latitude. A latitudinal trend is best displayed by clinopyroxene,
747 and specifically by deeply etched to skeletal clinopyroxene grains, which are most common north of
748 Luanda where etched garnet is also recorded. Garnet, widely considered as relatively resistant
749 during diagenesis (Morton and Hallsworth, 2007), resulted to be very unstable in hyperhumid
750 equatorial rift highlands (fig. 9 in Garzanti et al., 2013) and yet it is common in all beach samples
751 between Luanda and the Congo mouth, and most abundant in Ambriz beach and Luculu river sands.
752 Garnet becomes rare only in rivers and beaches of the Congo, indicating stronger weathering
753 conditions than in northern Angola.

754

755 *8.3. Chemical proxies*

756 Among chemical indices of weathering, only the WIP correlates very well with latitude ($r = 0.91$).
757 This is simply because it reflects the combined effects of weathering and recycling, which both
758 increase notably in northern Angola and very markedly in the Congo (Fig. 4). Among indices
759 unaffected by quartz dilution but sensitive to source-rock lithology, only the CIX correlates with
760 latitude relatively well ($r = 0.62$, sign. lev. 0.1%); α^{AlNa} , α^{AlK} and α^{AlRb} correlate poorly ($r = 0.42$ -
761 0.49 ; sign. lev. 1%). Indices considering Ca (CIA, PIA, α^{AlCa}) are uncorrelated with latitude, being
762 low to even very low in Bengo, Cuanza, and Longa river sands of northern Angola containing
763 carbonate rock fragments derived from Meso-Cenozoic basins (Fig. 8). Relative to aluminum and
764 with reference to the UCC standard, depletion in Na, Ca, and K is not particularly strong in northern
765 Angolan sands, suggesting that plagioclase is not much less durable than K-feldspar and that

766 recycling is more responsible than weathering for feldspar depletion (Nesbitt et al., 1997) (Fig. 4).
767 A markedly negative Eu anomaly produced by plagioclase breakdown (McLennan, 1993; Condie et
768 al., 1995) is not observed in river sands of northernmost Angola (e.g., Eu/Eu* 0.89 in Loge sand),
769 where plagioclase is as abundant as K-feldspar. Final plagioclase breakdown is indicated by the
770 sharp increase of the CIA, CIX, and especially of $\alpha^{Al}Na$ in the Congo (Table 1).

771

772 **9. Passive-margin sand and the arkose-quartzarenite problem**

773

774 The traditional view in sedimentary petrology is that passive-margin sands are generated from the
775 adjacent cratonic block — held to consist of gneissic-granitoid basement with cover strata — and
776 thus contain quartz and feldspar with few lithic fragments ("continental block provenance" of
777 Dickinson, 1985). Such sand suites are basically distinguished by their Q/F ratio, which may be ≤ 1
778 ("ideal arkose" of Dickinson, 1985) either where tectonic activity is particularly intense ("tectonic
779 arkose" of Folk, 1980; "basement uplift subprovenance" of Dickinson, 1985) and/or in arid climatic
780 conditions ("climatic arkose" of Folk, 1980). Instead, in case of tectonic quiescence favoring
781 prolonged weathering in low-relief shield areas, quartz would dominate over mainly K-feldspar
782 ("craton interior subprovenance" of Dickinson, 1985).

783 As a first approximation, this case study on sediment generation along the magma-poor Angolan
784 passive margin represents a neat confirmation of such a simplified conceptual framework. First-
785 cycle quartzo-feldspathic to feldspar-rich feldspatho-quartzose sand eroded from mid-crustal
786 granitoid gneisses of the Angola Block exposed in the dynamically uplifted Bié-Huila dome is
787 deposited in arid southern Angola, whereas quartz-rich feldspatho-quartzose to quartzose sand
788 characterizes the lower-relief and more intensely weathered rifted margin of humid northern
789 Angola, where sedimentary and metasedimentary rocks are more common. Pure quartzose, largely
790 recycled sand is generated in the vast, low-lying hyperhumid Congo basin (Giresse, 2005). Sand
791 composition in these three segments of the African passive margin thus reflects northward

792 decreasing dissection of the Congo craton (i.e., dissected, transitional, and undissected stages of
793 continental block provenance; [Garzanti, 2016](#)).

794 Such an apparently simple scenario stems from three distinct processes acting in accord. Stronger
795 tectonic uplift in the arid south, and stronger weathering coupled with extensive recycling in the
796 humid north determine a progressive northward increase of more durable quartz and heavy minerals
797 such as zircon, tourmaline and rutile, relative to less durable minerals such as plagioclase,
798 amphibole and pyroxene. The superposition of such partly interrelated effects makes it difficult to
799 quantify the relative importance of each, which has always represented a major drawback in the
800 paleotectonic and paleoclimatic interpretation of ancient passive-margin sandstones (e.g., [Johnsson,](#)
801 [1993](#)). Particularly hard to infer from detrital modes is the importance of recycling, because
802 chemical dissolution during diagenesis of parent sandstones may produce similar effects as
803 weathering of daughter sands ([Velbel and Saad, 1991](#)). One key, although chiefly qualitative, is
804 provided by the shape of quartz grains, rounded in the case of recycling especially of eolian parent
805 sandstones, deeply etched in the case of subequatorial weathering ([Fig. 7](#)).

806 Angolan sands present a relatively simple case also because this tract of the Atlantic rifted margin is
807 essentially non-volcanic, with limited syn-rift to post-rift magmatic activity confined to a few
808 coastal areas. Basaltic detritus derived from as far as Lesotho highlands represents $4 \pm 2\%$ of dune
809 and beach sands of the Moçamedes Desert, where it provides the most distinctive trace of ultralong-
810 distance littoral transport from the Orange river mouth. North of the Moçamedes Desert, basaltic
811 rocks fragments, clinopyroxene and detrital zircons with mid-Cretaceous U-Pb ages occur in a few
812 river sands (e.g., Cambongo, Quicombo) and in several beaches from southern to central Angola,
813 where volcanic detritus represents $\leq 2\%$ of bulk sand. Anorogenic volcanic provenance is thus
814 minor, which is not the case for many other regions of southern Africa still widely preserving
815 remnants of the widespread Jurassic Karoo and Cretaceous Etendeka flood basalts ([Garzanti et al.,](#)
816 [2014a](#)).

817

818 **10. Conclusion**

819

820 The N/S trending Angolan continental margin, located in the subequatorial belt of the Southern
821 Hemisphere where climate changes progressively with latitude from hyperarid in tropical Namibia
822 to hyperhumid in equatorial Congo, offers an exceptionally suitable natural laboratory in which to
823 study the effects of climatic-induced weathering on sediment composition. Excepting the
824 Moçamedes Desert in the south, where sand is mainly supplied by ultralong littoral transport from
825 the Orange river mouth, Angolan sand suites consist chiefly of quartz and feldspar derived first-
826 cycle from granitoid Precambrian basement but also partly recycled from Meso-Cenozoic basins
827 formed during initial rifting of the South Atlantic. Sand samples were thus formally classified by
828 their quartz/feldspar ratio as quartzo-feldspathic ($Q/F < 1$), feldspar-rich ($1 < Q/F < 2$) to quartz-rich
829 ($4 < Q/F < 9$) feldspatho-quartzose, and quartzose ($Q/F > 9$; pure quartzose if $Q\%QFL > 95$; [Fig.](#)
830 [2A](#)).

831 Quartzo-feldspathic to feldspar-rich feldspatho-quartzose sands characterize the arid south adjacent
832 to dynamically uplifted highlands. Moving toward the humid equator, river and beach sands pass
833 from feldspatho-quartzose in central Angola to mostly quartz-rich feldspatho-quartzose in northern
834 Angola, quartzose in northernmost Angola, and eventually pure quartzose in the Congo. Sharper
835 compositional trends are associated with the steep E/W climatic gradient, reflecting much stronger
836 chemical weathering in humid highlands than along the arid coast. Minor volcanic detritus is
837 supplied locally by Cretaceous lavas exposed close to the coast. Heavy-mineral suites are generally
838 dominated by epidote and amphibole. North of the Moçamedes Desert, clinopyroxene is present in a
839 few pocket beaches of southern Angola and widespread in beaches of central Angola. Garnet
840 associated with kyanite is common in beaches of northern Angola, where it is largely recycled from
841 sandstone strata of the Cuanza and Lower Congo basins and ultimately derived from
842 metasedimentary rocks of the West Congo Belt ([Fig. 3](#)). The dual time structure of Angolan source
843 rocks is best reflected by U-Pb ages of detrital zircons, characterized by unimodal spectra with
844 sharp Eburnean peak at ~ 2 Ga from north of Namibe to south of the Cuanza mouth, and by broader
845 multimodal distributions in northern Angola, with a few Permo-Triassic ages locally and additional

846 clusters at 0.5-0.7 Ga, ~ 1 Ga, and 2.5-2.7 Ga best defined in Congo estuary sand (Fig. 5). Among
847 chemical indices of weathering, the WIP correlates well with latitude because it reflects the
848 superposed effects of weathering and recycling. Other indices correlate poorly because they are
849 largely affected by source-rock lithology as well (Fig. 4). Depletion in Ca, Na, and K is not very
850 strong even in northern Angola, where plagioclase results to be not much less durable than K-
851 feldspar and recycling is more responsible than weathering for feldspar depletion. Feldspar and
852 garnet are weathered out only in the hyperhumid climate of the Congo.

853 This modern study of Angolan sands shows how extracting the climatic signal from sediment
854 composition is not straightforward, because multiple controlling factors operate in nature
855 simultaneously, and because mineralogical and geochemical proxies do not respond to weathering
856 only but are affected by other variables, including source-rock lithology, hydraulic sorting, and
857 recycling as well.

858

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868

869 **SUPPLEMENTARY MATERIAL**

870 Supplementary data associated with this article, to be found in the online version at
871 http://dx.doi._____, include information on sampling sites (Table A1) and river catchments

872 (Table A2), together with the complete datasets on bulk-sand petrography (Table A3), heavy
873 minerals (Table A4), surface textures (Table A5), sand geochemistry (Table A6) and detrital-zircon
874 geochronology (Appendix B). Figure A1 illustrating river profiles, and the Google-Earth™ map of
875 sampling sites *Angolamargin.kmz* are also provided.

876

877 **FIGURE CAPTIONS**

878

879 **Figure 1.** Topography, climate, and geology of Angola. **A)** Google Earth™ image showing the
 880 recently uplifted Bié-Huila plateau of southwestern Angola, lying largely at altitudes above 1500 m
 881 and delimited by very steep scarps cleaved by deep narrow gorges in the northwest (Lopes et al.,
 882 2016). Location of river (green), dune (orange), beach (yellow), and offshore samples (white;
 883 MARUM sites GeoB 1000->1008, 1011->1014, 1019->1022, 1702, 1704->1706, and 4918; DSDP
 884 Leg 40 site 365 and ODP Leg 175 sites 1078->1080 in *italics*) is indicated. **B)** Location and river
 885 map. **C)** Rainfall map (annual precipitation in mm; Hijmans et al., 2005). **D)** Geological sketch map
 886 (redrawn after Schlüter, 2008 and various sources cited in the text).

887

888 **Figure 2.** Petrographic classification and effects of weathering and recycling in subequatorial
 889 passive-margin sands. Data points: intensity of blue color increases with increasing rainfall at
 890 decreasing southern latitude; outline ranges from black to light grey with increasing estimated
 891 proportion of recycled component. **A)** The proposed refined classification of lithic-poor sands and
 892 sandstones. **B)** River sands chiefly derived first-cycle from basement rocks display a latitudinal
 893 trend of quartz enrichment controlled by climatic zonation. Latitude and average rainfall are
 894 indicated for Mebridege River in the north and Bero River in the south. **C)** The W/E inland trend of
 895 weathering-controlled quartz enrichment is well displayed in the Catumbela catchment. The Q/F
 896 ratio increases systematically in sand of tributaries draining progressively farther from the coast,
 897 being highest in trunk-river sand derived largely from humid interior highlands. The opposite trend
 898 is displayed by the P/F ratio, indicating selective dissolution of plagioclase. Longitude of
 899 headwaters and average rainfall are indicated for each catchment. **D)** River sands containing a
 900 significant recycled component as revealed by textural and mineralogical parameters are enriched
 901 further in quartz, reaching pure quartzose compositions in hyperhumid Congo. **E)** Cunene sands are
 902 a mixture of pure quartzose sand recycled from the fossil Kalahari Desert in the middle reaches and
 903 unweathered first-cycle quartzo-feldspathic sand eroded from crystalline basement in the dry lower

904 reaches (Fig 1C). Sand composition in the terminal tract of the Curoca River reflects extensive
 905 mixing with eolian coastal sand derived mainly from the Orange mouth.

906

907 **Figure 3.** Heavy-mineral suites of Angolan passive-margin sands. Moçamedes desert sands are
 908 richer than Skeleton Coast sands in amphibole and epidote contributed by the Cunene River;
 909 offshore sands are hydrodynamically depleted in garnet. Angolan basement chiefly supplies epidote
 910 and amphibole. From the Catumbela mouth to Luanda, beach sand is richer in clinopyroxene than
 911 river sand. From northern Angola to the Congo, beach sand is enriched in garnet, next in kyanite,
 912 and finally in staurolite, all largely recycled from Meso-Cenozoic strata of the northern Cuanza and
 913 Lower Congo basin and ultimately derived from metasedimentary rocks of the West Congo Belt.
 914 Both multivariate observations (points) and variables (rays) are displayed in the compositional biplot
 915 (Gabriel, 1971). The length of each ray is proportional to the variance of the corresponding element
 916 in the data set. If the angle between two rays is close to 0° , 90° , or 180° , then the corresponding
 917 elements are directly correlated, uncorrelated, or inversely correlated, respectively.

918

919 **Figure 4.** Latitudinal variation of key compositional parameters in river sands. Three-point-
 920 centered moving averages of all indices reflect extensive recycling of fossil Kalahari dunes in the
 921 middle Cunene catchment (right end of panels) and the combined effects of weathering and
 922 recycling in the Congo (left end of panels). **A)** The quartz/feldspar (Q/F) ratio shows the same trend
 923 as the CIX/WIP ratio, indicating the overwhelming effect of recycling. **B)** Metamorphic rank of
 924 rock fragments (MI) reflects mainly first-cycle provenance of Angolan sands from low to medium-
 925 grade basements; mixing with detritus from siliciclastic covers is major in rivers draining the
 926 Kalahari (right end), the Cuanza basin (center left), and the Congo basin (left end). Plagioclase (P)
 927 tends to decrease slightly with decreasing latitude relative to K-feldspar (KF; $r = 0.60$), and
 928 microcline (Mic) to increase imperceptibly relative to orthoclase ($r = 0.37$). **C)** The anticorrelation
 929 between the relative concentration of durable minerals (ZTR) and the absolute concentration of
 930 transparent heavy minerals (tHMC) is chiefly a recycling effect. **D)** Among the variants of the CIA

931 unaffected by quartz dilution, the CIX proves to be the most accurate indicator of weathering. The
 932 WIP, markedly affected by quartz dilution, decreases sharply in northern Angola. **E)** Among alfa
 933 indices: $\alpha^{Al}Na$ is the best indicator of weathering, although affected by recycling as well; $\alpha^{Al}Ca$ is
 934 affected by local occurrence of carbonate grains (e.g., Bengo sand); $\alpha^{Al}Mg$ and $\alpha^{Al}K$ show opposite
 935 behaviour, increasing in sands derived from felsic and mafic sources, respectively.

936

937 **Figure 5.** U-Pb age spectra of detrital zircons (age vs. frequencies plotted as Kernel Density
 938 Estimates using the *provenance* package of [Vermeesch et al., 2016](#)). Moçamedes desert sand in the
 939 south displays the same Damara and Namaqua peaks as Skeleton Coast sand mostly derived from
 940 the Orange River. Unimodal distributions with sharp Eburnean peak characterize sands of central
 941 Angola. Zircons from Cretaceous magmatic rocks are most common in Equimina beach. Spectra
 942 become broader and finally markedly multimodal in northern Angola, with prominent Pan-African,
 943 Namaqua, Eburnean, and Neoproterozoic clusters best displayed in Congo estuary sands.

944

945 **Figure 6.** Petrography of offshore sediments sampled during DSDP, ODP and Meteor expeditions.
 946 Shelfal sands to deep-sea sandy silts are variously enriched in planktonic foraminifers (Walvis
 947 Ridge), green glaucony (offshore Cunene mouth), benthic foraminifers (offshore Tapado mouth),
 948 and brown goethite ooids or pellets (offshore Congo mouth). Skeleton Coast dunes include basaltic
 949 rock fragments and clinopyroxene derived from the Orange River; the Congo River carries pure
 950 quartzose sand. Water depths are indicated; white circle for scale = 250 μm .

951

952

953 **Figure 7.** The effects of weathering and recycling on surface textures of sand grains. A,B)
 954 Numerous etch pits and re-entrants commonly filled with clay "plasma" testify to intense
 955 weathering even of quartz and other durable grains in river sands of northern Angola. Note strong
 956 leaching and deep penetration of iron-stained clay in: C) monocrystalline quartz; D) polycrystalline
 957 quartz; E) microcline. F,G) Abundance of rounded quartz grains testifies to extensive recycling of

958 Meso-Cenozoic siliciclastic units in the Cuanza and Congo catchments. All photos with crossed
 959 polars; blue bar for scale = 250 μm .

960

961 **Figure 8.** Inability of chemical indices to effectively single out weathering control from the effects
 962 of provenance, recycling, and hydraulic sorting in Angolan river sands. Both multivariate
 963 observations (points) and variables (rays) are displayed in the compositional biplot (Gabriel, 1971).
 964 Mobile alkali and alkaline earth elements largely reflect their original concentration in source-rock
 965 lithologies, and thus the CIA and its variants correlate poorly with latitude; $\alpha^{\text{Al}}\text{Mg}$, $\alpha^{\text{Al}}\text{Ca}$ and $\alpha^{\text{Al}}\text{Sr}$
 966 (higher for felsic sources) anticorrelate with $\alpha^{\text{Al}}\text{K}$, $\alpha^{\text{Al}}\text{Rb}$ and $\alpha^{\text{Al}}\text{Ba}$ (higher for mafic sources).
 967 Being affected by quartz dilution, the WIP decreases markedly toward the equator reflecting the
 968 combined effect of recycling and weathering. The more reliable indicator of weathering is $\alpha^{\text{Al}}\text{Na}$.
 969 The Eu anomaly tends to become more negative with weathering, but also with recycling and in
 970 sands derived from more felsic source rocks or in placer lags. The hydraulic-sorting effect is
 971 displayed by placer and semiplacer deposits formed at the Catumbela and Inamangando mouths
 972 (symbols with purple outline). Some beaches and a few river sands (e.g., Bengo, Longa; symbols
 973 with yellow outline) contain intrabasinal or extrabasinal carbonate grains and thus plot lower in the
 974 diagram because of their very low $\alpha^{\text{Al}}\text{Ca}$ (and low $\alpha^{\text{Al}}\text{Sr}$).

975

976 **Table 1.** Key petrographic, heavy-mineral, and geochemical parameters of Angolan passive-margin
 977 sands. N° = number of samples; Qz = quartz; KF = K-feldspar; Pl = plagioclase; L = aphanitic lithic
 978 grains (Lvm = volcanic and metavolcanic; Lsm = sedimentary and metasedimentary; Lm = high-
 979 rank metamorphic); MI* = metamorphic index; HM = heavy minerals; tHMC = transparent heavy-
 980 mineral concentration; ZTR = zircon + tourmaline + rutile; Ep = epidote; Grt = garnet; St =
 981 staurolite; Ky = kyanite; Amp = amphibole; Px = pyroxene; &tHM = other transparent heavy
 982 minerals (apatite, titanite, and rare sillimanite, andalusite, olivine, barite, monazite, allanite, or
 983 scheelite). Chemical weathering indices cannot be calculated reliably for offshore sediments

984 containing abundant intrabasinal grains and authigenic minerals, including calcareous and non-
985 calcareous allochems, glaucony, phosphates, sulphates, or iron ooids.

986

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Figure 1
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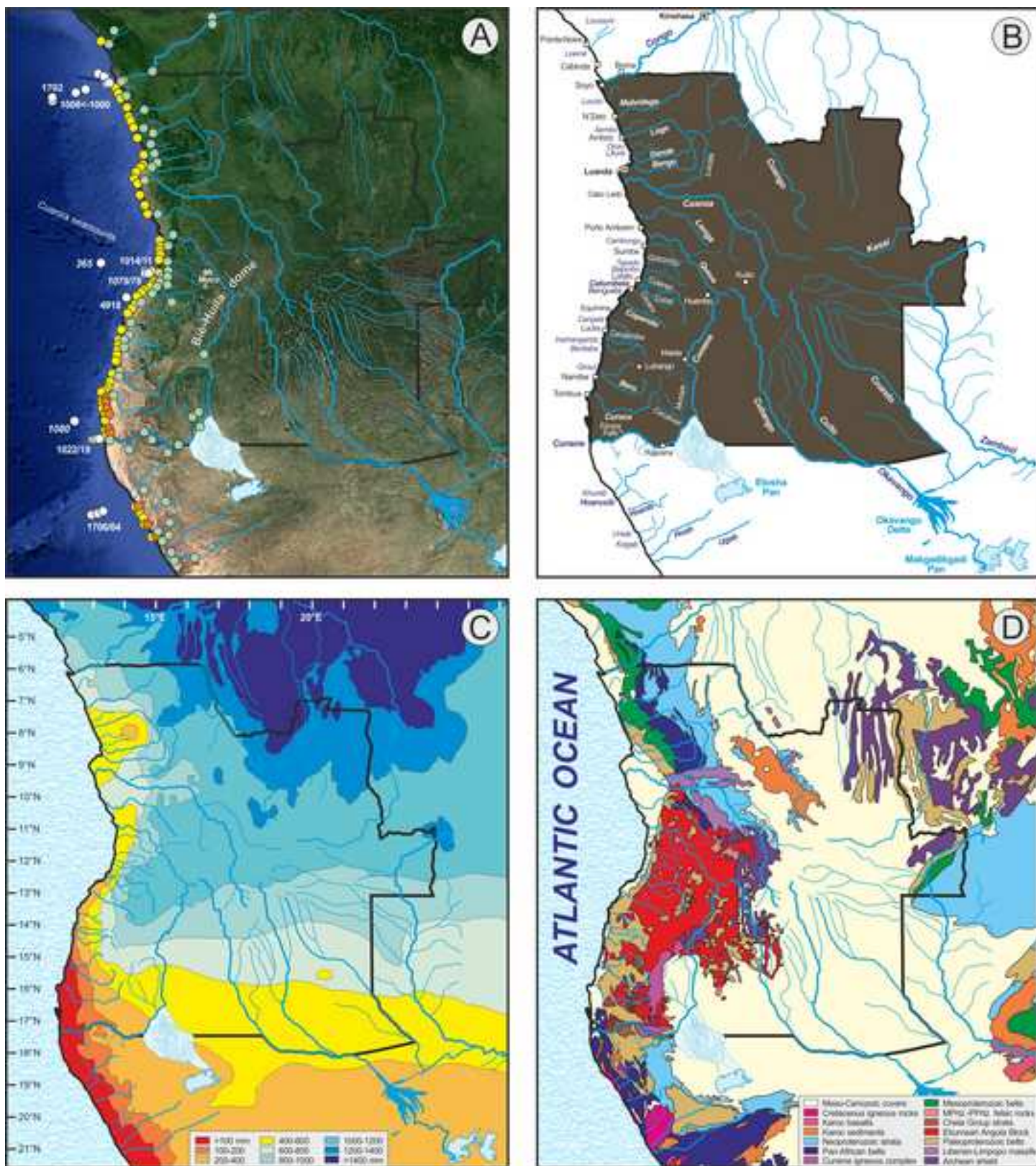


Figure 1 Angola

Figure 2
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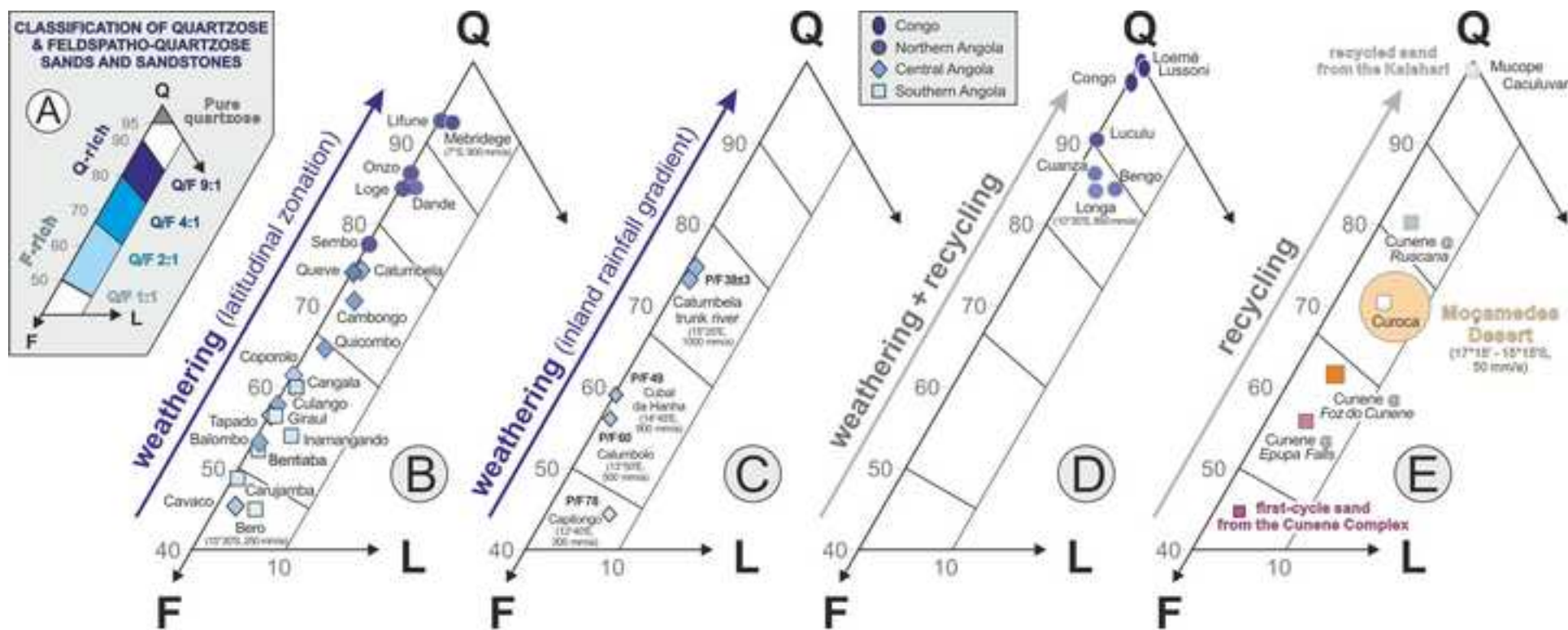


Figure 2 Angola

Figure 3
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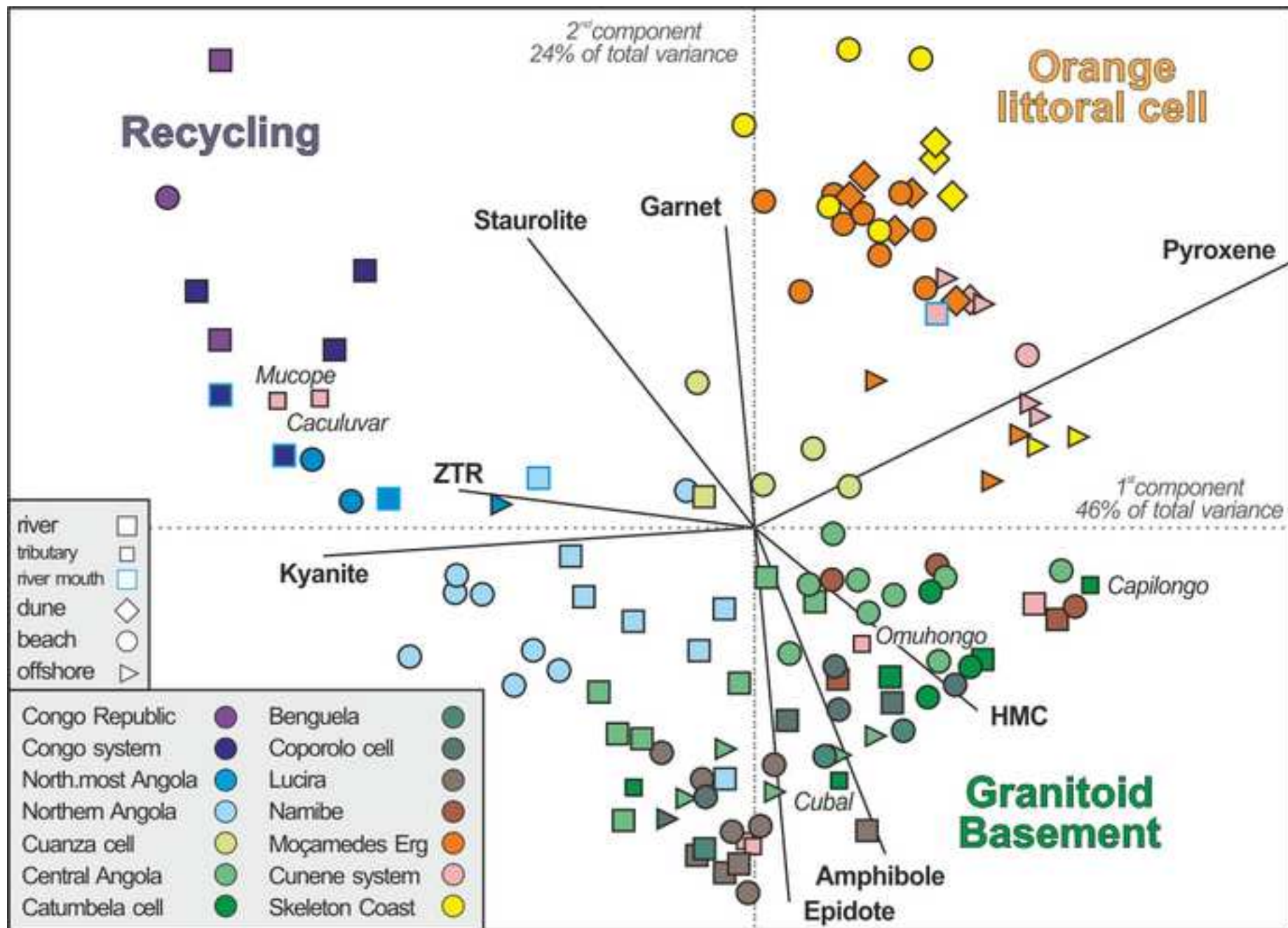


Figure 3 Angola

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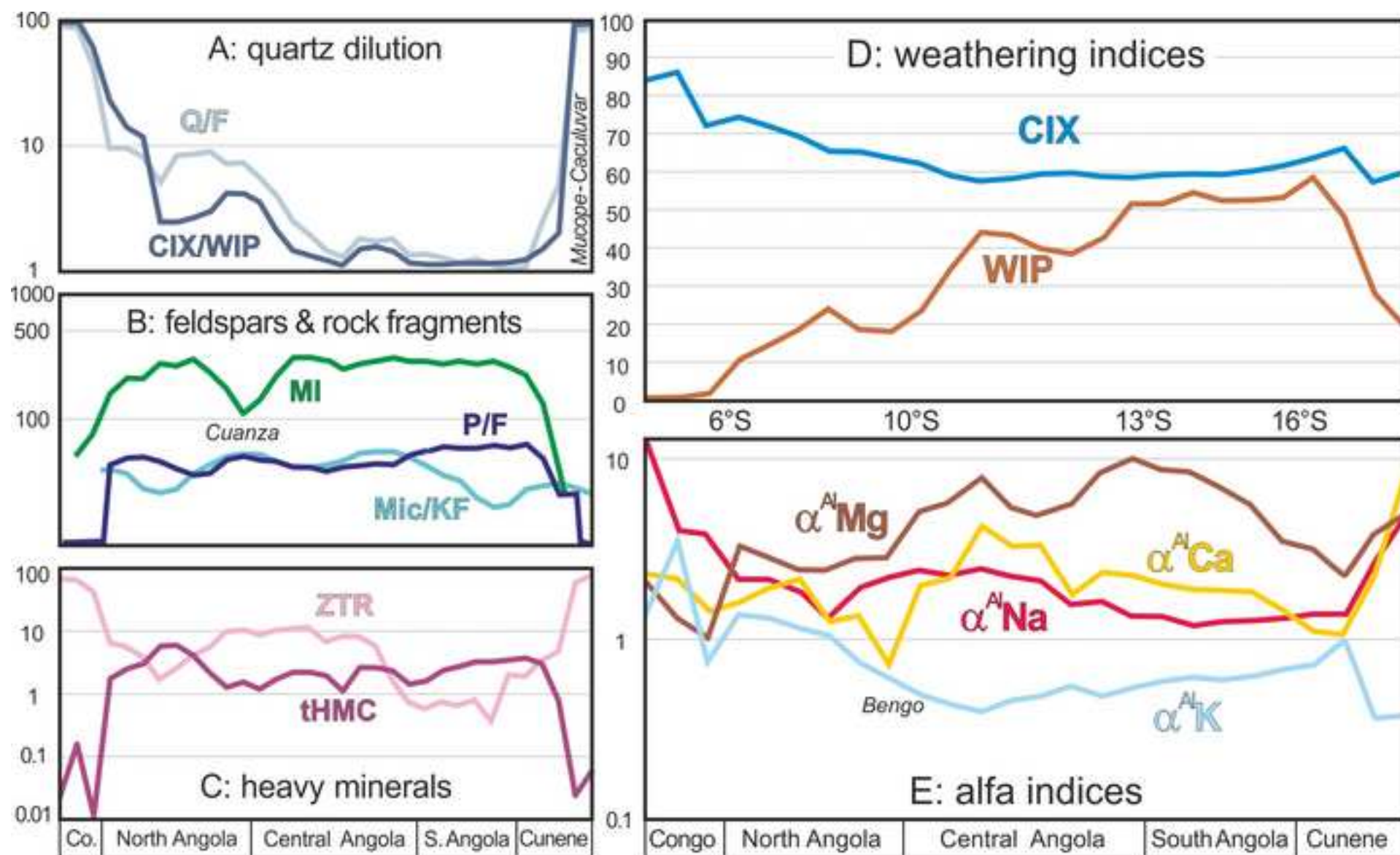


Figure 4 Angola

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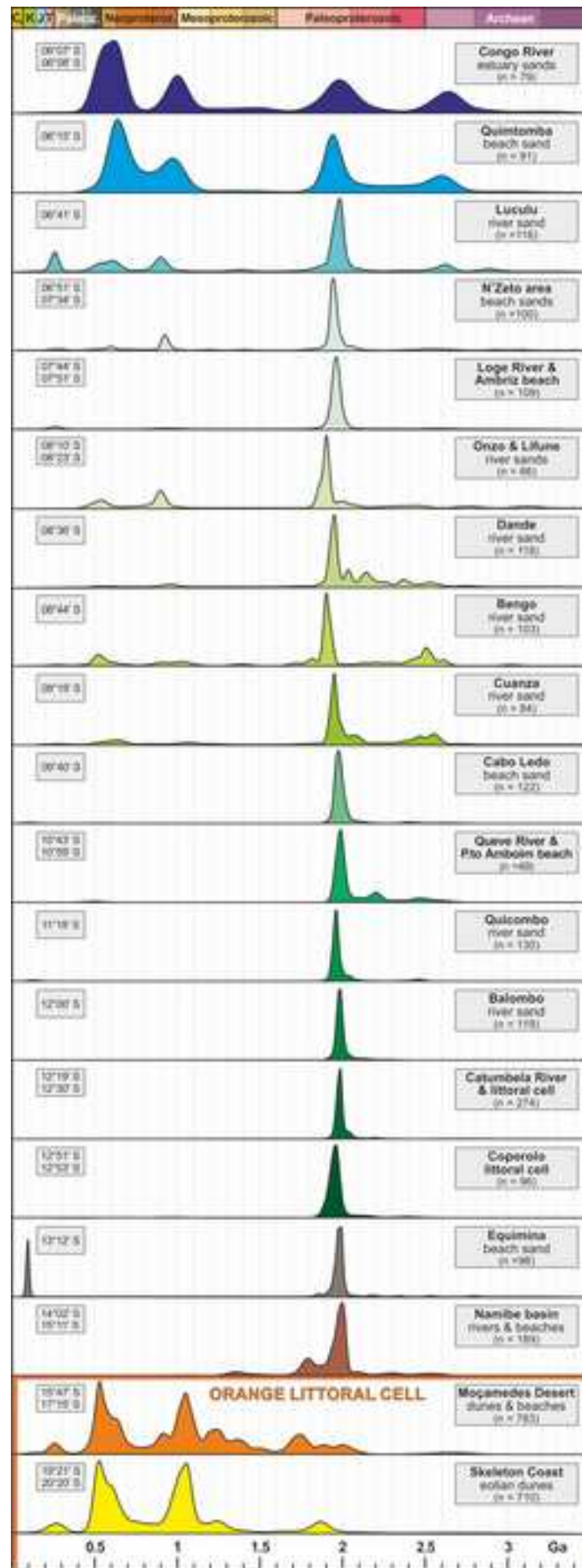


Figure 5 Angola

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Figure 6 Angola

Figure 7

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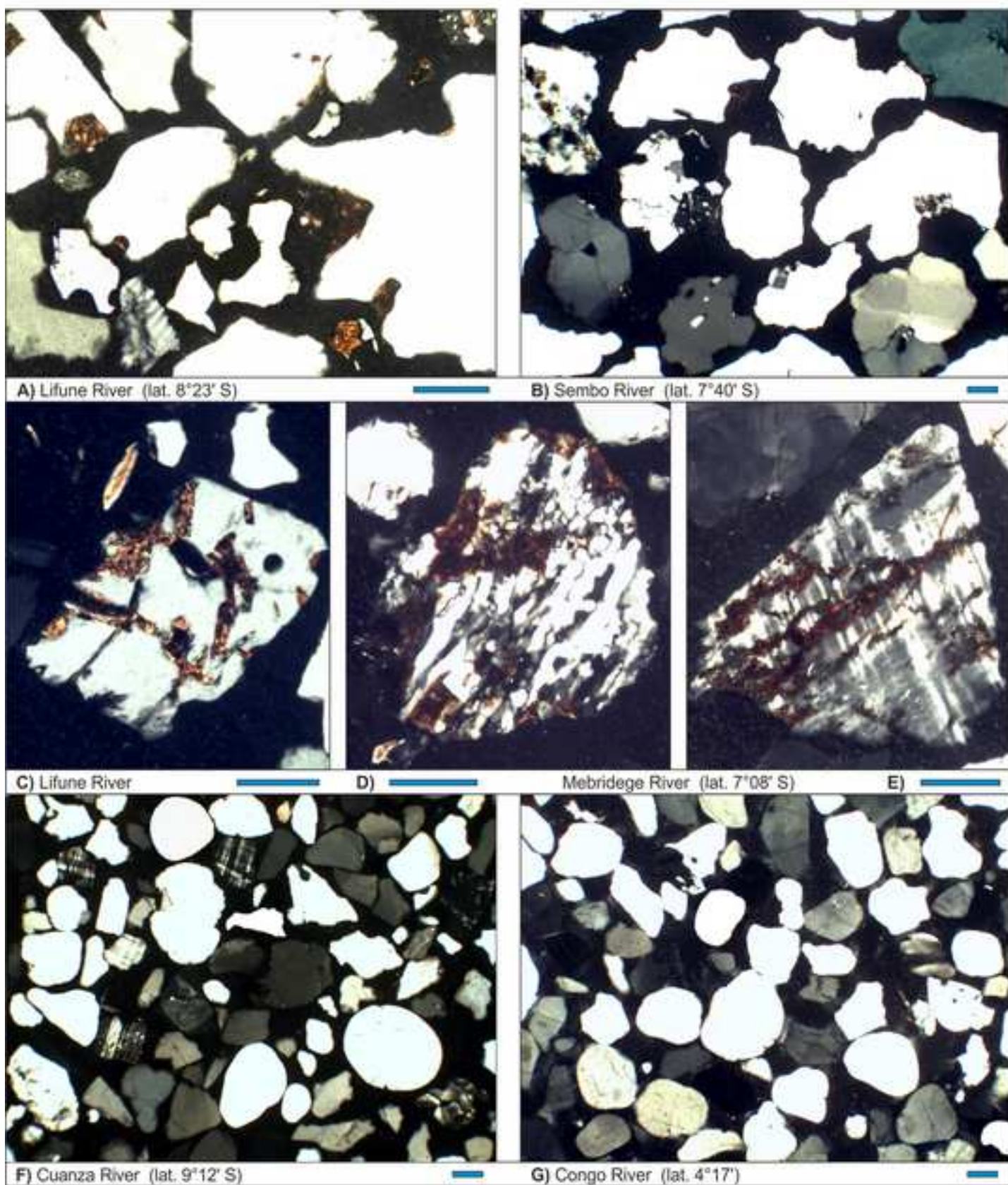


Figure 7 Angola

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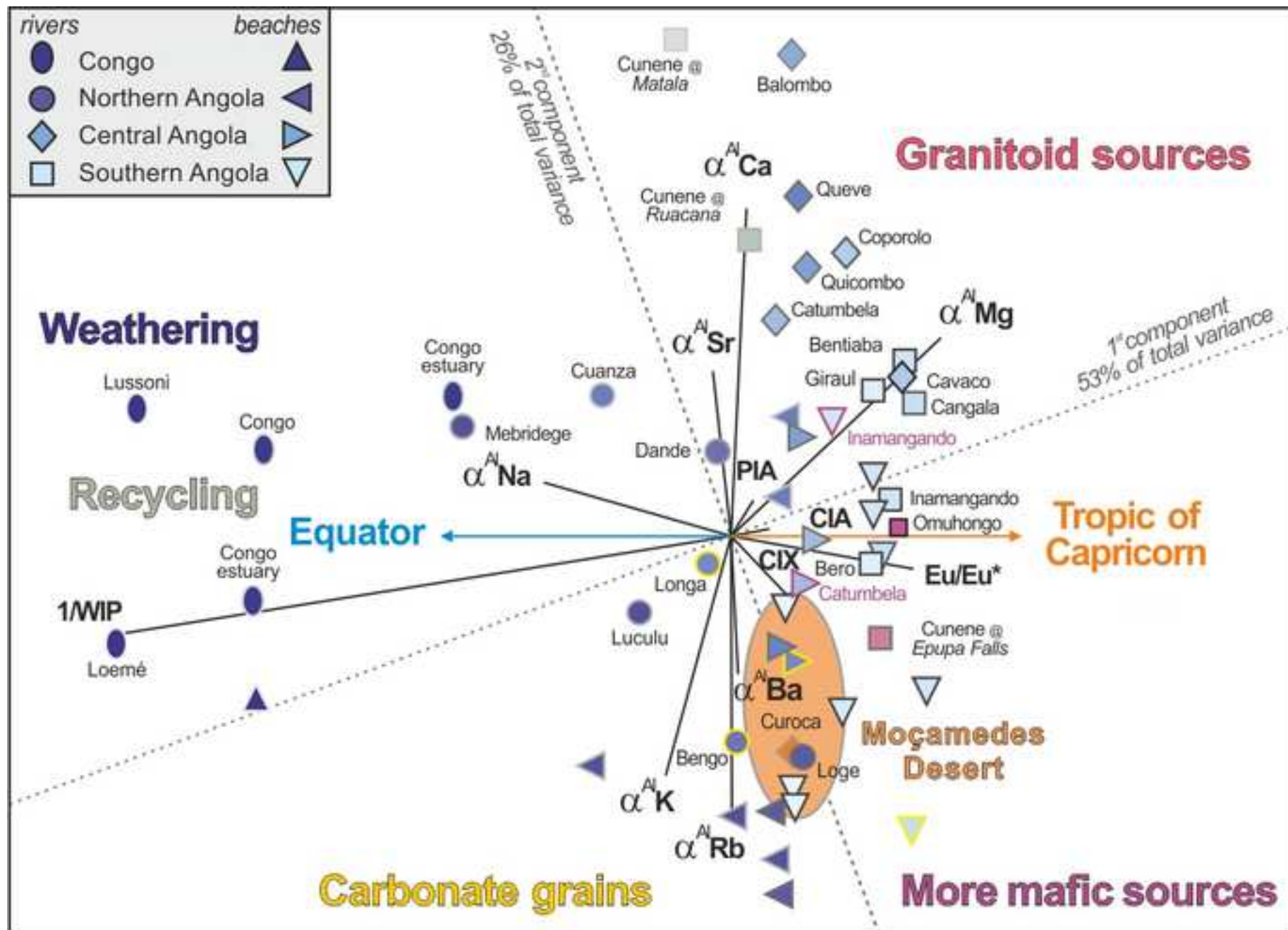


Figure 8 Angola

Table 1. Angola Margin

	N°	Qz	KF	PI	Lvm	Lsm	Lm	mica	HM		MI*	tHMC	ZTR	Ep	Grt	St	Ky	Amp	Px	&tHM	CIA	CIX	WIP	$\alpha^{Al}Na$	
RIVER SANDS																									
Republic of Congo	2	99	0	0	0	0	0	0	1	100.0		0.2	74	5	4	8	6	0	0	3	100.0	87	86	0.3	6.2
		0	0	0	0	0	0	0	0			0.2	4	4	6	3	6	0	0	2		10	6	0.3	0.2
Congo	3	98	1	0	0	0	0	0	1	100.0		0.2	57	13	2	14	7	3	0	3	100.0	63	72	1	3.9
		1	1	0	0	0	0	0	1			0.2	18	7	1	4	4	2	0	2					
Congo Estuary	2	98	1	1	0	0	0	0	0	100.0		0.1	41	14	3	7	20	12	0	4	100.0	81	85	2	7.8
		0	0	0	0	0	0	0	0			0.1	8	8	0	2	18	2	0	3		12	13	2	7.0
Luculu	1	86	4	5	0	0	0	0	5	100.0		3.8	7	30	33	2	15	13	0	1	100.0	62	74	11	2.2
Mebridege	1	93	3	3	0	1	1	0	0	100.0	183	0.02	9	8	1	0	3	71	1	7	100.0	70	75	4	3.2
Sembo	1	77	14	8	0	0	0	0	0	100.0		1.1	2	21	2	0	1	72	1	0	100.0				
Loge	1	76	6	8	0	0	0	0	10	100.0		7.3	1	70	2	0	0	26	0	2	100.0	50	67	27	1.1
Onzo	1	81	7	6	0	0	0	0	6	100.0	325	5.8	1	27	7	0	6	57	2	0	100.0				
Lifune	1	87	5	1	0	0	0	2	5	100.0		3.3	7	35	11	0	11	34	1	0	100.0				
Dande	1	80	8	6	0	0	1	4	2	100.0	275	1.3	6	48	6	0	4	36	1	1	100.0	58	65	24	1.3
Bengo	1	82	6	5	0	4	0	0	4	100.0		2.1	6	30	13	1	10	37	1	1	100.0	33	64	20	1.5
Cuanza	1	87	5	6	0	1	0	0	0	100.0		0.3	18	52	5	0	2	15	7	1	100.0	54	66	10	2.9
Longa	1	82	6	6	1	2	0	0	3	100.0		2.1	8	85	0	0	0	6	0	1	100.0	38	62	23	2.3
Queue	1	73	16	9	0	0	0	1	1	100.0		0.6	1	43	2	0	0	51	0	2	100.0	55	59	37	2.1
Cambongo	1	68	13	14	2	0	0	1	2	100.0	425	1.4	22	51	1	0	0	18	6	1	100.0				
Quicombo	1	60	21	10	0	1	1	2	5	100.0	340	3.2	10	64	3	0	0	15	2	6	100.0	51	57	43	2.3
Tapado	1	56	27	16	0	0	0	1	0	100.0		0.4	3	78	1	0	0	17	0	0	100.0				
Balombo	1	53	26	19	0	0	0	1	0	100.0		0.5	8	89	0	0	0	3	1	0	100.0	55	57	52	3.1
Culango	1	57	23	18	0	0	0	1	1	100.0		0.9	14	71	0	0	0	14	0	1	100.0				
Catumbela	2	69	15	9	0	0	0	0	6	100.0	317	2.2	4	59	0	0	0	18	14	4	100.0	54	59	32	1.9
		5	2	0	0	0	0	0	7			1.8	3	14	0	0	0	4	11	5					
Cavaco	1	43	26	24	0	0	1	1	3	100.0	350	1.9	1	71	0	0	0	24	0	3	100.0	52	59	48	1.3
Coporolo	2	61	20	16	0	0	0	1	2	100.0	300	1.2	2	67	0	0	0	28	3	1	100.0	53	58	48	1.7
		1	4	2	0	0	0	0	1			0.7	1	0	0	0	0	0	3	1					
Cangala	1	59	13	25	0	0	1	0	2	100.0	267	1.8	0	76	0	0	0	22	0	2	100.0	52	58	57	1.1
Carujamba	1	48	21	29	0	0	0	0	2	100.0		1.5	0	63	0	0	0	35	0	1	100.0				
Inamangando	1	52	18	22	1	1	2	0	5	100.0	275	3.1	1	78	0	0	0	18	0	3	100.0	51	61	49	1.2
Bentiaba	1	52	19	28	0	0	1	0	1	100.0	317	2.0	0	62	0	0	0	36	0	2	100.0	52	58	56	1.2
Giraul	1	53	17	23	0	0	1	0	5	100.0	250	3.2	1	42	3	0	0	52	1	2	100.0	52	59	51	1.3
Bero	1	42	16	32	0	1	2	1	5	100.0	350	3.2	0	21	0	0	0	60	16	2	100.0	51	62	51	1.2
Cunene@Matala	1	75	19	6	0	0	0	0	0	100.0		0.1	93	2	0	1	1	1	1	3	100.0	58	60	20	4.7
Cunene@Ruacana	1	79	12	4	1	3	0	0	1	100.0		0.7	5	32	1	0	0	55	0	7	100.0	51	57	28	2.7
Cunene@Epupa F.	1	54	13	26	1	1	1	0	3	100.0	295	3.0	1	27	1	0	0	44	23	6	100.0	51	66	48	1.4
Cunene@Foz	1	57	12	20	3	1	1	0	6	100.0	256	4.2	1	19	10	4	0	33	28	4	100.0				
BEACH SANDS																									
Republic of Congo	1	99	1	0	0	0	0	0	0	100.0		0.1	67	1	0	13	18	0	0	1	100.0	>59	>89	<0.9	>11
North.most Angola	2	96	2	1	0	0	0	0	1	100.0		0.6	9	35	13	7	16	19	0	1	100.0	42	72	7	1.8
		2	1	0	0	0	0	0	1			0.1	5	6	0	3	2	5	0	0					
Northern Angola	8	83	6	3	0	1	0	0	6	100.0	425	4.4	4	43	19	1	6	24	1	2	100.0	39	67	23	1.2
		7	2	1	0	0	0	0	5			3.0	3	12	12	1	3	9	2	1		5	2	4	0.2
Cuanza cell	4	81	11	6	0	0	0	0	1	100.0		0.9	4	51	3	1	1	22	14	4	100.0	39	55	27	1.5
		2	1	2	0	0	0	0	1			0.5	3	4	2	1	1	5	7	1		2	0	3	0.1
Central Angola	9	71	13	11	0	1	0	0	5	100.0	325	3.7	7	54	1	0	0	19	14	4	100.0	38	60	38	1.7
		3	2	2	0	1	0	0	3			2.3	5	8	1	0	1	4	9	4		9	2	5	0.1
Catumbela cell	3	61	14	13	0	0	1	1	10	100.0	374	6.2	4	63	1	0	0	19	13	1	100.0	47	61	43	1
		13	3	4	0	0	1	1	6		66	4.4	2	6	1	0	0	6	5	0					
Benguela	1	64	11	18	0	2	1	1	4	100.0	260	4.0	2	71	0	0	0	18	4	6	100.0				
Coporolo cell	4	65	19	12	0	0	0	0	3	100.0		2.1	1	72	0	0	0	22	2	2	100.0				
		4	6	3	0	0	0	0	3			1.7	1	7	0	0	0	7	1	0					
Southern Angola	10	50	16	23	0	1	1	0	9	100.0	319	5.4	5	55	2	0	0	32	3	3	100.0	47	62	52	1.1
		6	5	7	0	1	1	0	8		55	4.0	8	16	2	0	0	16	6	3		4	2	6	0.2
Moçamedes Desert	13	67	9	13	3	0	0	0	7	100.0	282	4.7	4	14	35	4	0	15	26	2	100.0	49	66	28	1.3
		6	3	2	1	1	0	0	7		87	4.1	2	3	21	2	0	7							

APPENDIX

"Dynamic uplift, recycling, and climate control on the petrology of passive-margin sand (Angola)"

by Garzanti E., Dinis P., Vermeesch P., Andò S., Hahn A., Huvi J., Limonta M., Padoan, M., Resentini, A., Rittner M., Vezzoli, G.

APPENDIX A

Figure A1. Channel profiles of Angolan rivers draining into the Atlantic Ocean (same horizontal and vertical scale for all profiles). The alternation of subhorizontal and very steep tracts along the course of major rivers reflects the presence of stepped planation surfaces separated by escarpments, a characteristic feature of the dynamically uplifted landscapes of southwestern Angola. Fluvial network delineated in TecDEM (software shell implemented in MATLAB; [Shahzad and Gloaguen, 2011](#)) from a 30 m resolution digital elevation model provided by ASTER GDEM (<http://www.gdem.aster.ersdac.or.jp>). Channel concavity θ and steepness k_s (referenced to a fixed concavity 0.45 to compare gradients in channels with different drainage areas; [Korup and Schlunegger, 2009](#)) are defined by a power-law relationship between the local channel slope S and the contributing drainage area A used as a proxy for discharge ($S = k_s A^{-\theta}$; [Flint 1974](#)).

Table A1. Sample information. Location of the studied river, beach, shelfal and deep-sea sediment samples with year of sampling (see also the Google Earth file [Angolamargin.kmz](#)).

Table A2. Geomorphology and hydrology of Angolan river systems (data after [National Directorate of Water, 2005](#)). Sediment loads of Congo, Cuanza and Cunene rivers after [Hay \(1998\)](#), [Holisticos \(2012\)](#), and [Bremner and Willis \(1993\)](#), respectively.

Table A3. Sand petrography. GSZ= grain size. Q= quartz (Qp= polycrystalline); F= feldspars (KF= K-feldspar; P= plagioclase; Mic= cross-hatched microcline); L= aphanitic lithic grains (Lv= volcanic and subvolcanic; Ls= sedimentary; Lc= carbonate; Lh= chert; Lp= shale/siltstone; Lm= metamorphic; Lms= low-rank metasedimentary; Lmv= low-rank metavolcanic; Lmf= high-rank

metapelite/metapsammite/metafelsite; Lmb= high-rank metabasite; Lu= ultramafic). HM= heavy minerals. Rock fragments: V= volcanic; Vm= intermediate and mafic volcanic; M= metamorphic; Mb= mafic metamorphic; n.d. = not determined. The Metamorphic Indices MI and MI* express the average metamorphic rank of rock fragments in each sample. MI varies from 0 (detritus shed by exclusively sedimentary and volcanic cover rocks) to 500 (very-high-rank detritus shed by exclusively high-grade basement rocks). MI* considers only metamorphic rock fragments, and thus varies from 100 (very-low-rank detritus shed by exclusively very low-grade metamorphic rocks) to 500 (Garzanti and Vezzoli, 2003).

Table A4. Heavy minerals. GSZ= grain size. HM= heavy minerals; tHM= transparent heavy minerals; HMC and tHMC = total and transparent-heavy-mineral concentration indices (Garzanti and Andò, 2007); RF= rock fragments; n.d. = not determined. The ZTR index (sum of zircon, tourmaline and rutile over total transparent heavy minerals) evaluates the “chemical durability” of the detrital assemblage (Hubert 1962). The HCI (Hornblende Colour Index) and MMI (Metasedimentary Minerals Index) vary from 0 in detritus from greenschist-facies to lowermost amphibolite-facies rocks yielding exclusively blue/green amphibole and chloritoid, to 100 in detritus from granulite-facies rocks yielding exclusively brown hornblende and sillimanite, and are used to estimate the average metamorphic grade of metaigneous and metasedimentary source rocks, respectively (Andò et al. 2014).

Table A5. Surface textures of heavy minerals in Angolan sediment samples. Determination of corrosion features on transparent heavy-mineral grains by three operators, following the classification of Andó et al. (2012). Q= quartz; F= feldspar; L= aphanitic lithic grains; n.d.= not determined.

Table A6. Geochemistry of Angolan sands (analyses made at ACME Laboratories, Vancouver). Following a lithium metaborate/tetraborate fusion and nitric acid digestion, major oxides and several minor elements were determined by inductively coupled plasma emission spectroscopy, and trace elements by ICP-MS. Discrepancies in replicate analyses are $\leq 1\%$ for major elements and $\leq 5\%$ for most trace elements (for further information on adopted procedures, geostandards used, and precision for various elements of group 4A-4B and code LF200 see <http://acmelab.com>). A separate split was digested in aqua regia and analyzed by ICP-MS for Mo, Ni, Cu, Ag, Au, Zn, Cd, Hg, Tl, Pb, As, Sb, Bi, and Se. Elements analysed by aqua regia digestion (a.r.d.) are commonly underestimated because of only partial leaching of refractory minerals. Chemical weathering indices are defined in Nesbitt and Young (1982; CIA), Harnois (1988; CIW), Fedo et al. (1995; PIA),

Parker (1970; WIP), and Garzanti et al. (2014; CIX). In order to avoid bias caused by hydraulic concentration of heavy minerals hosting Ti, REE and Th, α^{Al} values were normalized to non-mobile Al (Garzanti et al., 2014). The Eu anomaly is the measured chondrite-normalized Eu value over the value that Eu would have in a linear extrapolation between chondrite-normalized values of Sm and Gd. The Ce anomaly, indicative of redox state, is the measured PAAS-normalized Ce value over the value that Ce would have in a linear extrapolation between PAAS-normalized values of La and Pr. MREE is the average of Eu, Gd, Tb and Dy normalized to PAAS, MREE* the average of LREE (La, Ce, Pr, Nd) and HREE (Er, Tm, Yb, Lu) values (Haley et al., 2004). The chondrite-normalized La_N/Yb_N , La_N/Sm_N , Gd_N/Ho_N , and Ho_N/Yb_N ratios are also given. GSZ= grain size; D.L. = detection limit; n.d.= not determined.

APPENDIX B

U-Pb detrital zircon geochronology of modern sands from northern Namibia and southern Angola (analyses made at the London Geochronology Centre, University College London). We used $^{206}Pb/^{238}U$ and $^{207}Pb/^{206}Pb$ ages for zircons younger and older than 1100 Ma, respectively; grains with >10% age discordance were discarded. No common Pb correction was applied. Grains with +5/-15% age discordance were discarded.

CITED REFERENCES

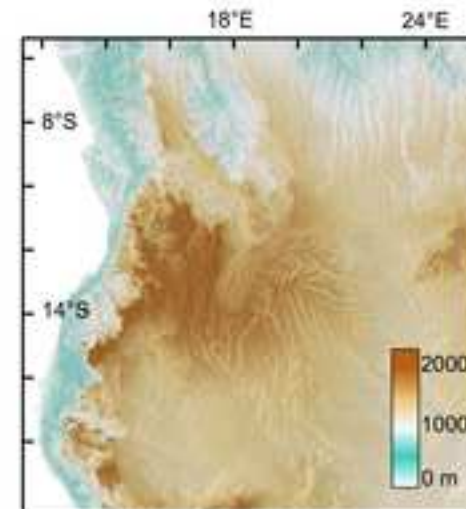
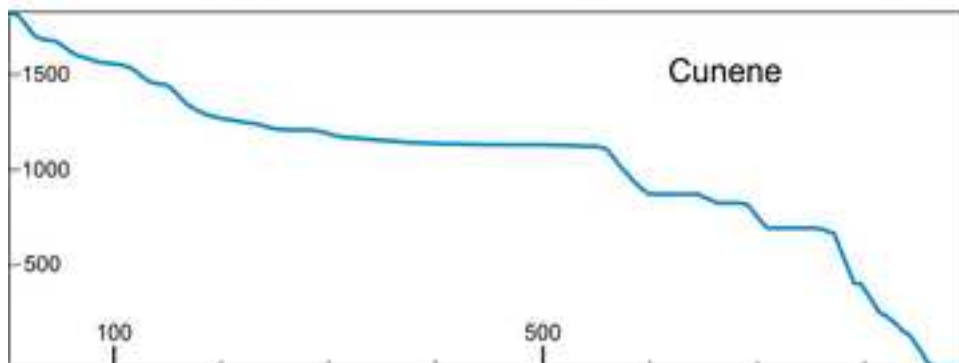
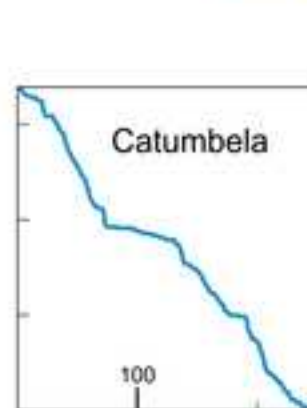
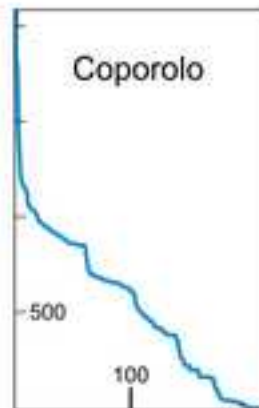
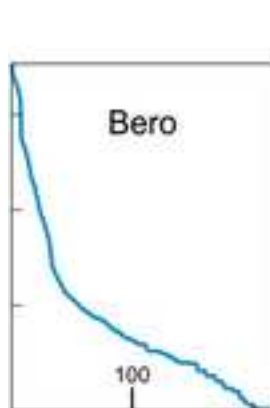
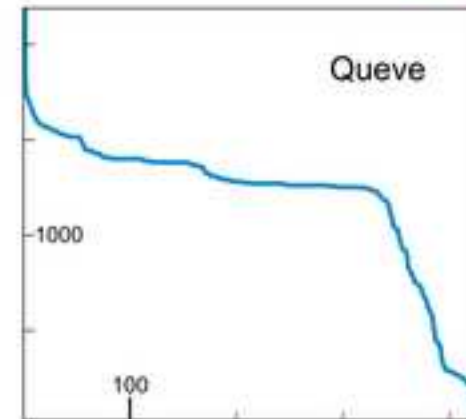
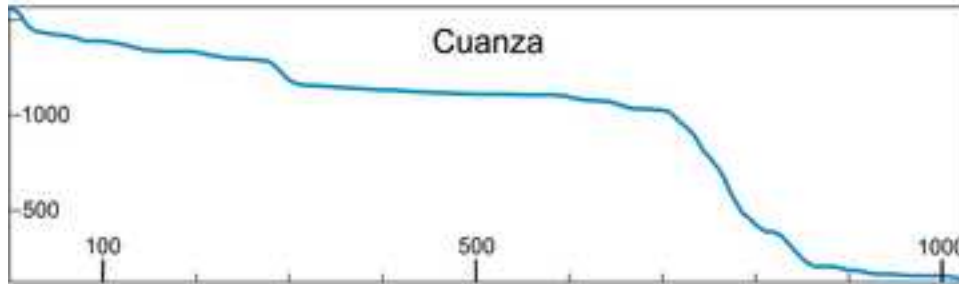
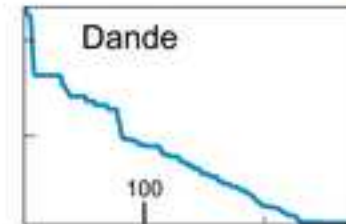
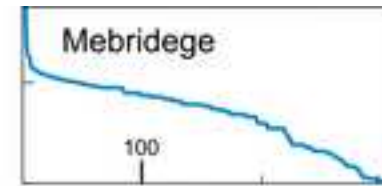
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Appendix Figure A1

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River	Area km ²	Length km	Elevation mean m	k_{in}	θ
Mebridege	19071	290	589	121	0.13
Dande	11446	285	621	141	0.29
Longa	23031	310	888	106	0.09
Queve	22815	300	1360	114	0.11
Catumbela	16533	240	1321	107	-0.06
Coporolo	15239	220	908	94	0.21
Bero	10476	210	718	121	0.31
Cunene	113835	1050	1286	66	-0.02



Appendix Table A1

[Click here to download Table: Table A1 Angola Samples.pdf](#)

Sample	River / Desert	Site	Facies	Country	Collected by	Year	Latitude	Longitude
S4053	Loussoni	St. Paul	river sand	Congo	M.Orlando	2009	S 04 28	E 12 12
S0162		Pointe Noire	beach sand	Congo	D. Dell'Era	1997	S 04 48	E 11 50
S0161	Loémé	Djeno	river sand	Congo	D. Dell'Era	1997	S 04 54	E 11 56
S3533	Congo	Brazzaville	river sand	Congo	F.Bolognesi	2005	S 4 17	E 15 17
S5114	Congo	Kinshasa	river sand	Congo	P.Vermesch	2016	S 04 19 32	E 15 13 42
S5115	Congo	Boma	river sand	Congo	P.Vermesch	2016	S 05 51 58	E 13 00 14
S4899	Congo	Soyo	estuary sand	Angola	E.Garzanti	2015	S 06 07 52	E 12 22 25
S4898	Congo	Soyo estuary	estuary sand	Angola	E.Garzanti	2015	S 06 07 09	E 12 21 32
S4901		Quimtomba	beach sand	Angola	E.Garzanti	2015	S 06 15 00	E 12 20 46
S4902		Quifuma	beach sand	Angola	E.Garzanti	2015	S 06 23 30	E 12 25 37
S4904	Luculu	Quivanda	estuary sand	Angola	E.Garzanti	2015	S 06 41 15	E 12 34 52
S4903		Quivanda	beach sand	Angola	E.Garzanti	2015	S 06 41 17	E 12 34 49
S4905		Mbua-Moyo	beach sand	Angola	E.Garzanti	2015	S 06 51 42	E 12 44 32
S4897	Mebridege	Cassa de Telha	river sand	Angola	E.Garzanti	2015	S 07 08 24	E 12 58 58
S4906		N'Zeto	beach sand	Angola	E.Garzanti	2015	S 07 13 42	E 12 51 14
S4907		Pidi mouth	beach sand	Angola	E.Garzanti	2015	S 07 18 02	E 12 53 03
S4908		Musserra	beach sand	Angola	E.Garzanti	2015	S 07 34 51	E 13 00 16
S4895	Sembo	Musserra	river sand	Angola	E.Garzanti	2015	S 07 39 35	E 13 11 26
S4894	Loge	Porta Freitas Morna	river sand	Angola	E.Garzanti	2015	S 07 43 56	E 13 19 56
S4909		Ambriz	beach sand	Angola	E.Garzanti	2015	S 07 50 58	E 13 06 17
S4893	Onzo	Tabi	river sand	Angola	E.Garzanti	2015	S 08 09 47	E 13 25 31
S4892	Lifune	Conde Loca	river sand	Angola	E.Garzanti	2015	S 08 23 18	E 13 28 16
S4910		Barra do Dande	beach sand	Angola	E.Garzanti	2015	S 08 28 04	E 13 22 58
S4891	Dande	Caxito	river sand	Angola	E.Garzanti	2015	S 08 36 23	E 13 33 33
S4911	Bengo	Quifangondo	estuary sand	Angola	E.Garzanti	2015	S 08 43 48	E 13 26 17
S4912		Cacuaco	beach sand	Angola	E.Garzanti	2015	S 08 46 24	E 13 22 15
S4652		Ilha de Luanda	beach sand	Angola	R.Giardini	2013	S 8 47	E 13 14E
S3764		Mussulo	beach sand	Angola	M.Orlando	2008	S 8 59	E 13 02
S4913		Palmeirinhas	beach sand	Angola	E.Garzanti	2015	S 09 07 16	E 13 00 57
S4914		Barra do Cuanza	estuary sand	Angola	E.Garzanti	2015	S 09 20 24	E 13 09 02
S4915	Cuanza	Barra do Cuanza	river sand	Angola	E.Garzanti	2015	S 09 19 22	E 13 09 45
S3765		Cabo Ledo	beach sand	Angola	M.Orlando	2008	S 9 40	E 13 13
S4916		Sao Braz	beach sand	Angola	E.Garzanti	2015	S 09 58 27	E 13 19 31
S4917	Longa	Calamba	river sand	Angola	E.Garzanti	2015	S 10 11 50	E 13 31 16
S4918		Porto Amboim	beach sand	Angola	E.Garzanti	2015	S 10 42 36	E 13 46 30
S4919		Sumbe	beach sand	Angola	E.Garzanti	2015	S 11 09 47	E 13 50 19
S4920		Quicombo	beach sand	Angola	P.Dinis	2015	S 11 19 05	E 13 48 55
S4921		Candunga	beach sand	Angola	P.Dinis	2015	S 11 44 24	E 13 47 39
S4922	Queue	Cachoeira	river sand	Angola	E.Garzanti	2015	S 10 59 19	E 14 05 45
S4923		Carimba	beach sand	Angola	E.Garzanti	2015	S 11 04 35	E 13 51 20
S4924	Cambongo	Sumbe	river sand	Angola	E.Garzanti	2015	S 11 11 56	E 13 50 47
S4925	Quicombo	Quicombo	river sand	Angola	E.Garzanti	2015	S 11 19 18	E 13 50 25
S4926	Tapado	Chitonde	river sand	Angola	E.Garzanti	2015	S 11 48 38	E 13 56 47
S4927	Balombo	Canjala	river sand	Angola	E.Garzanti	2015	S 11 59 45	E 13 59 45
S4928		Egito Praia	beach sand	Angola	E.Garzanti	2015	S 11 57 37	E 13 45 42
S4929	Culango	Culango	river sand	Angola	E.Garzanti	2015	S 12 18 04	E 13 49 54
S4930		Praia Sousa	beach sand	Angola	P.Dinis	2015	S 11 36 36	E 13 47 06
S4934		Lobito spit	beach sand	Angola	E.Garzanti	2015	S 12 18 58	E 13 34 52
S4865		Lobito spit	beach sand	Angola	P.Dinis	2014	S 12 19 14	E 13 34 45
S4863		Lobito spit	beach sand	Angola	P.Dinis	2014	S 12 20 19	E 13 33 04
S4864		Lobito spit	beach sand	Angola	P.Dinis	2014	S 12 20 22	E 13 33 06
S4862		Catumbela N	beach sand	Angola	P.Dinis	2014	S 12 24 51	E 13 29 42
PB.AC		Catumbela mouth	beach sand	Angola	P.Dinis	2014	S 12 26 14	E 13 28 58
S4861		Catumbela mouth	beach sand	Angola	P.Dinis	2014	S 12 27 02	E 13 28 33
S4655	Catumbela	Praia Bebe	beach placer	Angola	A. Pereira	2013	S 12 27 00	E 13 28 30
S5118	Cubal da Hanha	Vista Alegre	river sand	Angola	J.Huvi	2017	S 12 49 44	E 13 59 14
CAV	Catumbela	Fazenda Santo Antonio	river sand	Angola	P.Dinis	2014	S 12 30 17	E 13 46 54
S5116	Calumbolo	Supua	river sand	Angola	J.Huvi	2017	S 12 27 00	E 13 46 06
S5117	Capilongo	Capilongo	river sand	Angola	J.Huvi	2017	S 12 27 50	E 13 36 22
S4933	Catumbela	Catumbela	river sand	Angola	E.Garzanti	2015	S 12 26 27	E 13 33 03
S4860	Catumbela	Catumbela	river sand	Angola	P.Dinis	2014	S 12 27 00	E 13 34 32
S4859		Catumbela airport	beach sand	Angola	P.Dinis	2014	S 12 30 17	E 13 28 56
S4932	Cavaco	Benguela	river sand	Angola	E.Garzanti	2015	S 12 34 00	E 13 25 04
S4931		Benguela	beach sand	Angola	E.Garzanti	2015	S 12 34 45	E 13 23 40
S4935		Baia Farta	beach sand	Angola	E.Garzanti	2015	S 12 35 51	E 13 12 15
S4858		Baia Farta	beach sand	Angola	P.Dinis	2014	S 12 36 57	E 13 11 26
S4857		Canucua	beach sand	Angola	P.Dinis	2014	S 12 47 53	E 12 59 14
S4856		Saco NE	beach sand	Angola	P.Dinis	2014	S 12 51 25	E 12 57 03

S4855		Saco N	beach sand	Angola	P.Dinis	2014	S 12 51 42	E 12 56 18
S4854		Saco SE	beach sand	Angola	P.Dinis	2014	S 12 52 30	E 12 57 16
S4853		Saco S	beach sand	Angola	P.Dinis	2014	S 12 52 46	E 12 56 39
S4852	Coporolo	Santa Tereza	beach sand	Angola	P.Dinis	2014	S 12 54 01	E 13 05 20
S4936	Coporolo	Dombe Grande	river sand	Angola	E.Garzanti	2015	S 12 55 01	E 13 06 20
S4937		Cuio	beach sand	Angola	E.Garzanti	2015	S 12 58 58	E 12 58 43
S4938		Equimina	beach sand	Angola	E.Garzanti	2015	S 13 11 49	E 12 46 54
S4939		Baia Binga	beach sand	Angola	E.Garzanti	2015	S 13 20 19	E 12 39 06
S4940	Cangala	Santa Maria	river sand	Angola	E.Garzanti	2015	S 13 33 33	E 12 38 28
S4941		Lucira	beach sand	Angola	E.Garzanti	2015	S 13 51 58	E 12 31 14
S4942	Carujamba	Lucira	river sand	Angola	E.Garzanti	2015	S 13 59 19	E 12 31 01
S4943	(low outer berm)	Inamangando	beach sand	Angola	E.Garzanti	2015	S 14 02 33	E 12 23 09
S4944	(high inner berm)	Inamangando	beach sand	Angola	E.Garzanti	2015	S 14 02 34	E 12 23 10
S4945	Inamangando	Inamangando	river sand	Angola	E.Garzanti	2015	S 14 03 04	E 12 25 38
S4946		Baia das Salinas	beach sand	Angola	E.Garzanti	2015	S 14 11 18	E 12 20 38
S4947	Bentiaba	Bentiaba	river sand	Angola	E.Garzanti	2015	S 14 16 05	E 12 22 46
S4948		Bentiaba	beach sand	Angola	E.Garzanti	2015	S 14 17 23	E 12 22 11
P4948		Bentiaba	beach placer	Angola	E.Garzanti	2015	S 14 17 23	E 12 22 11
S4949		Chapeu Armado	beach sand	Angola	E.Garzanti	2015	S 14 26 55	E 12 20 38
S4950		Mariquita	beach sand	Angola	E.Garzanti	2015	S 14 45 41	E 12 17 04
S4951	Giraul	Giraul	river sand	Angola	E.Garzanti	2015	S 15 04 30	E 12 09 18
S4952	Bero	Namibe	river sand	Angola	E.Garzanti	2015	S 15 09 54	E 12 10 05
S4953		Namibe	beach sand	Angola	E.Garzanti	2015	S 15 11 32	E 12 08 50

Moçamedes Desert

S4954		Subida Grande	beach sand	Angola	E.Garzanti	2015	S 15 25 21	E 12 01 54
S4802	(fossil dune)	Nonguai	fossil dune	Angola	P.Dinis	2014	S 15 45 57	E 12 04 40
S4955	Curoca	Curoca mouth	river sand	Angola	E.Garzanti	2015	S 15 43 53	E 11 55 24
S4956		Curoca mouth	beach sand	Angola	E.Garzanti	2015	S 15 43 56	E 11 54 38
S4804		Tombua	beach sand	Angola	P.Dinis	2014	S 15 47 54	E 11 51 18
S4805	(outer spit)	Tombua	beach sand	Angola	P.Dinis	2014	S 15 47 20	E 11 49 08
S4774		Tombua	eolian dune	Angola	E.Baptista	2014	S 15 47 56	E 11 51 54
S4957		Vanesa	beach sand	Angola	E.Garzanti	2015	S 15 57 09	E 11 46 06
S4961		Cova dos Medos	eolian dune	Angola	E.Garzanti	2015	S 16 01 24	E 11 48 47
S4958		Vanesinha	beach sand	Angola	E.Garzanti	2015	S 16 09 25	E 11 47 36
S4959		Praia do Navio	beach sand	Angola	E.Garzanti	2015	S 16 16 23	E 11 48 35
S4960		Praia do Navio	eolian dune	Angola	P.Vermeesch	2015	S 16 16 25	E 11 48 44
S5058		Riscos	beach sand	Angola	A.Sampaio	2016	S 16 30 00	E 11 49 21
S5059		Riscos	eolian dune	Angola	A.Sampaio	2016	S 16 30 00	E 11 49 21
S5057		Saco dos Tigres	beach sand	Angola	A.Sampaio	2016	S 16 48 27	E 11 48 13
S5055		Praia dos Esponjas	beach sand	Angola	A.Sampaio	2016	S 17 05	E 11 44 40
S5056		Praia dos Esponjas	eolian dune	Angola	A.Sampaio	2016	S 17 05	E 11 44 40
S5054		Foz do Cunene	eolian dune	Angola	A.Sampaio	2016	S 17 15 24	E 11 45 18
S5053		Foz do Cunene	beach sand	Angola	A.Sampaio	2016	S 17 15 24	E 11 45 18

Cunene catchment

S5050	Caculuar	Techango	river sand	Angola	A.Trindade	2016	S 16 38 15	E 14 54 16
S5049	Mucupe	Techiulo	river sand	Angola	A.Trindade	2016	S 16 31 46	E 14 52 22
S3931	Cunene	Ruacana	river sand	Namibia	L.Ciceri	2008	S 17 24 30	E 14 13
S3934	Omuhongo	Oryeheke	river sand	Namibia	L.Ciceri	2008	S 16 59 20	E 13 22 10
S4775	Cunene	Epupa Falls	river sand	Namibia	F.Vermeesch	2014	S 17 00	E 13 15
S5052	Cunene	Foz do Cunene	estuary sand	Angola	A.Sampaio	2016	S 17 15 24	E 11 45 18

OFFSHORE CORE TOP SAMPLES

			Depth (m)		Corer			
1004	Meteor M6/6	GeoB1004-3	-31	Congo mouth	Giant box corer	1988	S 06 05 58	E 12 07 39
1001	Meteor M6/6	GeoB1001-1	-45	Congo mouth	Giant box corer	1988	S 05 51 58	E 11 58 20
1011	Meteor M6/6	GeoB1011_2	-73	Tabado mouth	Giant box corer	1988	S 11 48 14	E 13 39 41
1012/1	Meteor M6/6	GeoB1012_1	-99	Tabado mouth	Giant box corer	1988	S 11 48 33	E 13 35 27
1012/2	Meteor M6/6	GeoB1012_2	-99	Tabado mouth	Giant box corer	1988	S 11 48 33	E 13 35 27
1013	Meteor M6/6	GeoB1013_2	-250	Tabado mouth	Giant box corer	1988	S 11 47 55	E 13 26 47
2/2 5-7	ODP Leg 175	1078 A	-438	Balombo mouth	Piston corer	1997	S 11 55 13	E 13 24 08
3/1 15-17	ODP Leg 175	1078 A	-438	Balombo mouth	Piston corer	1997	S 11 55 13	E 13 24 08
1/2 105-106	DSDP Leg 40	365	-3040	Cuanza volcanic ridge	Rotary core barrel	1975	S 11 39 06	E 11 53 43
1/6 65-67	DSDP Leg 40	365	-3040	Cuanza volcanic ridge	Rotary core barrel	1975	S 11 39 06	E 11 53 43
4918	Meteor M41/1	GeoB4918_5	-1339	Coporolo mouth	Gravity corer	1998	S 12 50 24	E 12 41 48
1/1 89-91	ODP Leg 175	1080 A	-2766	Baia dos Tigres	Piston corer	1997	S 16 33 35	E 10 49 12
2/4 31-33	ODP Leg 175	1080 A	-2766	Baia dos Tigres	Piston corer	1997	S 16 33 35	E 10 49 12
1/6 65-67	ODP Leg 175	1080 B	-2768	Baia dos Tigres	Piston corer	1997	S 16 33 36	E 10 49 12
1019	Meteor M6/6	GeoB1019-3	-75	Cunene mouth	Giant box corer	1988	S 17 10 29	E 11 38 50
1020	Meteor M6/6	GeoB1020-1	-110	Cunene mouth	Giant box corer	1988	S 17 10 07	E 11 32 53
1021	Meteor M6/6	GeoB1021-3	-173	Cunene mouth	Giant box corer	1988	S 17 10 34	E 11 24 00
1022	Meteor M6/6	GeoB1022-2	-551	Cunene mouth	Giant box corer	1988	S 17 10 24	E 11 17 53
1704	Meteor 20/2	GeoB1704-1	-399	Walvis Ridge	Giant box corer	1992	S 19 24 24	E 11 36 42
1705	Meteor 20/2	GeoB1705-1	-642	Walvis Ridge	Gravity corer	1992	S 19 30 18	E 11 23 54

Appendix Table A2

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Table A2

River	Area <i>km²</i>	Length <i>km</i>	Perimeter <i>km</i>	Elevation		Annual water discharge				Precipitation <i>mm/a</i>	Sdm.load <i>ton/a</i>	Sdm.yield <i>ton/km2 a</i>	Erosion <i>mm/a</i>
				mean <i>m a.s.l.</i>	max <i>m a.s.l.</i>	mean <i>m³/s</i>	mean spec. <i>l/s km²</i>	max spec. <i>l/s km²</i>	min spec. <i>q</i>				
Congo	4014500	4700	11895	955	1548	41000	8.7	11.6	5.5	1375	60000000	15	0.006
Luculu	1449	95	185	151	315	8	5.6	6.0	4.7	836			
Mebridege	19071	290	890	589	1295	124	6.5	11.4	3.5	974			
Sembo	2093	125	250	284	734	7	3.3	4.4	2.6	550			
Loge	12819	230	581	531	1281	53	4.1	10.4	1.9	612			
Onzo	2942	130	324	359	913	7	2.4	2.8	2.0	415			
Lifune	3019	150	334	376	1133	9	2.9	4.5	2.4	513			
Dande	11446	285	649	621	1474	59	5.2	10.0	2.4	832			
Bengo	11089	300	663	483	1530	44	3.9	7.4	1.5	883			
Cuanza	150446	965	2702	1200	1964	1064	7.1	14.2	1.6	1188	618500	4	0.0016
Longa	23031	310	791	888	2099	138	6.0	9.1	3.3	991			
Queve	22815	300	899	1360	2575	213	9.4	16.3	2.6	1131			
Cambongo	2309	310	286	867	2260	14	5.9	10.9	2.6	763			
Quicombo	5512	50	434	1081	2545	40	7.2	15.6	0.9	965			
Tapado	1617	40	190	582	1583	7	4.1	7.4	2.3	745			
Balombo	4414	200	451	1186	2609	42	9.6	18.4	2.8	1100			
Culango	2881	120	266	949	2142	13	4.3	8.4	2.3	883			
Catumbela	16533	240	748	1321	2570	149	9.0	17.3	1.8	1182			
Cavaco	4398	120	313	738	1570	19	4.4	6.8	1.6	751			
Coporolo	15239	220	667	908	2406	70	4.6	9.2	1.1	846			
Cangala	363	60	115	547	827	0.3	0.8	1.1	0.4	312			
Carujamba	2931	160	305	664	1230	5	1.8	3.8	0.2	560			
Inamangando	1859	145	263	623	1477	2	1.1	3.0	0.1	479			
Bentiaba	6935	210	472	873	2325	12	1.8	3.7	0.2	648			
Giraul	4709	190	393	615	2322	4	0.8	2.5	0.1	409			
Bero	10476	210	588	718	2094	5	0.5	1.5	0.1	364			
Curoca	19338	324	849	762	1864	4	0.2	0.9	0.0	238			
Cunene	113835	1050	2390	1286	2484	290	2.5	13.7	0.0	704	8700000	76	0.029

Appendix Table A3

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River	Site	Sample	Operator	GSZ (mm)	Q	KF	P	Lv	Lc	Lh	Lp	Lms	Lmv	Lmf	Lmb	Lu	mica	H/M	total	Q	F	L	MP	MI	Lm	Lv	Ls	OpQ	P/F	MicF	VmvV	Mb/M	
RIVERS & BEACHES IN THE CONGO																																	
Loussoni	St. Paul	S4053	G.Vezzoli	325	99	0	0	0	0	0	1	0	0	0	0	0	0	0	100.0	99	0	1	n.d.	100	n.d.	n.d.	n.d.	8	n.d.	n.d.	n.d.	n.d.	
	Pointe Noire	S0162	G.Vezzoli	510	99	1	0	0	0	0	0	0	0	0	0	0	0	0	100.0	99	0	0	n.d.	n.d.	n.d.	n.d.	n.d.	4	n.d.	100	n.d.	n.d.	
Loémé	Djeno	S0161	G.Vezzoli	220	99	0	0	0	0	0	0	0	0	0	0	0	0	1	100.0	100	0	0	n.d.	n.d.	n.d.	n.d.	n.d.	2	n.d.	n.d.	n.d.	n.d.	
	Brazzaville	S3533	G.Vezzoli	240	97	2	0	0	0	0	0	0	0	0	0	0	0	0	100.0	97	2	0	n.d.	n.d.	n.d.	n.d.	n.d.	2	0	38	n.d.	n.d.	
Congo	Kinshasa	S5114	G.Vezzoli	160	98	1	0	0	0	0	0	0	0	0	0	0	0	1	100.0	99	1	0	n.d.	n.d.	n.d.	n.d.	n.d.	2	n.d.	n.d.	n.d.	n.d.	
	Boma	S5115	G.Vezzoli	125	97	1	1	0	0	0	0	0	0	0	0	0	0	1	100.0	99	1	0	n.d.	n.d.	n.d.	n.d.	n.d.	1	40	75	n.d.	n.d.	
Congo	Sojo	S4899	G.Vezzoli	175	98	1	1	0	0	0	0	0	0	0	0	0	0	0	100.0	98	2	0	n.d.	n.d.	n.d.	n.d.	n.d.	1	57	0	n.d.	n.d.	
	Sojo estuary	S4898	A.Resentini	180	98	1	1	0	0	0	0	0	0	0	0	0	0	0	100.0	99	1	0	n.d.	n.d.	n.d.	n.d.	n.d.	3	40	40	n.d.	n.d.	
BEACHES & SHELF IN NORTHERN ANGOLA																																	
-31 m inner shelf offshore of Congo mouth																																	
	Quimtoomba	S4901	G.Vezzoli	630	97	1	1	0	0	0	0	0	0	0	0	0	0	0	100.0	98	2	0	n.d.	n.d.	n.d.	n.d.	n.d.	4	43	57	n.d.	n.d.	
	Quiluma	S4902	A.Resentini	345	95	3	1	0	0	0	0	0	0	0	0	0	0	1	100.0	96	4	0	n.d.	n.d.	n.d.	n.d.	n.d.	4	31	17	n.d.	n.d.	
	Quiwanda	S4903	A.Resentini	350	88	6	3	0	0	0	0	0	0	0	0	0	0	3	100.0	91	9	0	n.d.	n.d.	n.d.	n.d.	n.d.	10	35	20	n.d.	n.d.	
	Mbua-Moyo	S4905	A.Resentini	380	90	5	4	0	0	0	0	0	0	0	0	0	0	2	100.0	91	9	0	n.d.	n.d.	n.d.	n.d.	n.d.	9	43	18	n.d.	n.d.	
	Pidi mouth	S4907	A.Resentini	245	77	9	3	0	0	0	0	0	0	0	0	0	0	10	100.0	86	14	0	n.d.	n.d.	n.d.	n.d.	n.d.	8	27	8	n.d.	n.d.	
	Musserra	S4908	G.Vezzoli	325	72	6	6	1	0	0	0	0	0	0	0	0	0	15	100.0	85	14	2	n.d.	n.d.	13	63	25	5	49	13	n.d.	n.d.	
	Ambriz	S4909	A.Resentini	330	82	5	3	0	1	0	0	0	0	0	0	0	0	9	100.0	90	9	1	n.d.	n.d.	n.d.	n.d.	n.d.	3	39	23	n.d.	n.d.	
	Barra do Dande	S4910	G.Vezzoli	460	82	9	3	0	0	0	0	0	0	1	0	0	0	4	100.0	86	13	1	425	309	50	0	50	7	36	60	n.d.	0	
	Cacuaco	S4912	A.Resentini	120	92	4	1	0	1	0	0	0	0	0	0	0	1	1	100.0	94	4	2	n.d.	n.d.	17	0	83	4	20	36	n.d.	n.d.	
RIVERS IN NORTHERN ANGOLA																																	
Luculu	Quiwanda	S4904	A.Resentini	370	86	4	5	0	0	0	0	0	0	0	0	0	0	5	100.0	91	9	0	n.d.	n.d.	n.d.	n.d.	n.d.	4	53	34	n.d.	n.d.	
	Cassa de Telha	S4897	A.Resentini	345	93	3	3	0	0	0	0	1	0	0	0	0	0	0	100.0	93	6	1	183	122	60	20	20	55	6	n.d.	0		
Mebridege	Sembo	S4895	A.Resentini	750	77	14	8	0	0	0	0	0	0	0	0	0	0	0	100.0	78	22	0	n.d.	n.d.	n.d.	n.d.	n.d.	35	38	19	n.d.	n.d.	
	Loge	S4894	A.Resentini	305	76	6	8	0	0	0	0	0	0	0	0	0	0	10	100.0	84	15	0	n.d.	n.d.	n.d.	n.d.	n.d.	12	55	9	n.d.	n.d.	
Onzo	Tabi	S4893	A.Resentini	480	81	7	6	0	0	0	0	0	0	0	0	0	0	6	100.0	86	14	0	325	325	n.d.	n.d.	n.d.	4	43	19	n.d.	13	
	Lifune	S4892	A.Resentini	345	87	5	1	0	0	0	0	0	0	0	0	0	2	5	100.0	93	7	0	n.d.	n.d.	n.d.	n.d.	n.d.	7	21	27	n.d.	n.d.	
Dande	Caxito	S4891	A.Resentini	150	80	8	6	0	0	0	0	0	0	1	0	0	4	2	100.0	85	14	2	275	275	100	0	0	4	42	27	n.d.	0	
	Bengo	S4891	A.Resentini	135	82	6	5	0	4	0	0	0	0	0	0	0	0	4	100.0	85	11	4	n.d.	n.d.	0	0	100	6	45	27	n.d.	0	
CUANZA & CUANZA LITTORAL CELL																																	
Cuanza	Ilha de Luanda	S4652	G.Vezzoli	570	81	10	7	0	0	0	0	0	0	0	0	0	0	0	100.0	81	18	1	n.d.	143	0	25	75	17	42	35	n.d.	n.d.	
	Mussulu	S3764	G.Vezzoli	550	83	13	4	0	0	0	0	0	0	0	0	0	0	0	100.0	83	17	0	n.d.	n.d.	n.d.	n.d.	n.d.	23	24	44	n.d.	n.d.	
Cuanza	Palmeirinhas	S4913	A.Resentini	560	82	10	7	0	0	0	0	0	0	0	0	0	0	1	100.0	83	17	0	n.d.	n.d.	n.d.	n.d.	n.d.	19	40	20	n.d.	n.d.	
	Barra do Cuanza	S4914	A.Resentini	530	79	12	8	0	1	0	0	0	0	0	0	0	0	1	100.0	79	20	1	n.d.	n.d.	13	38	50	22	41	21	n.d.	n.d.	
	Barra do Cuanza	S4915	A.Resentini	260	87	5	6	0	1	0	0	0	0	0	0	0	0	0	100.0	87	12	1	n.d.	n.d.	0	0	100	7	54	25	n.d.	n.d.	
BEACHES & SHELF IN CENTRAL ANGOLA																																	
Cabo Ledo	Sao Braz	S3765	G.Vezzoli	360	69	14	8	1	1	0	0	0	0	0	0	0	0	7	100.0	74	24	2	n.d.	140	0	60	40	20	37	26	n.d.	n.d.	
	Porto Amboim	S4916	A.Resentini	205	67	14	13	0	2	0	0	0	0	0	0	0	0	3	100.0	70	28	3	n.d.	n.d.	6	0	94	14	48	29	n.d.	n.d.	
Carimba	Praia Sousa	S4918	A.Resentini	430	72	11	11	0	0	0	0	0	0	0	0	0	0	6	100.0	76	23	1	n.d.	n.d.	n.d.	n.d.	n.d.	26	52	17	n.d.	n.d.	
	Egito Praia	S4923	A.Resentini	350	71	8	10	0	0	0	0	0	0	0	1	0	0	9	100.0	78	21	1	325	260	60	20	20	22	55	20	n.d.	100	
-73 m shelf offshore of Tapado mouth	1011	G.Vezzoli	90	40	26	24	0	1	0	0	0	0	0	0	0	0	2	6	100.0	44	54	2	n.d.	n.d.	20	0	80	2	48	21	n.d.	n.d.	
	-99 m shelf offshore of Tapado mouth	1012/2	G.Vezzoli	115	47	26	17	0	2	0	0	0	0	0	0	0	0	7	100.0	51	46	3	n.d.	n.d.	13	0	88	9	40	19	n.d.	n.d.	
RIVERS IN CENTRAL ANGOLA																																	
Longa	Calamba	S4917	A.Resentini	130	82	6	6	1	2	0	0	0	0	0	0	0	0	3	100.0	84	13	3	n.d.	75	9	27	64	6	51	27	n.d.	n.d.	
	Queve	S4922	A.Resentini	350	73	16	9	0	0	0	0	0	0	0	0	0	1	1	100.0	74	26	0	n.d.	n.d.	n.d.	n.d.	n.d.	25	34	34	n.d.	n.d.	
Cambongo	Sunbe	S4924	A.Resentini	335	68	13	14	2	0	0	0	0	0	0	0	0	1	2	100.0	71	27	2	425	283	14	86	0	21	52	18	n.d.	25	
	Quicombo	S4925	A.Resentini	260	60	21	10	0	1	0	0	0	1	1	0	0	2	5	100.0	65	33	2	340	340	67	0	33	26	33	28	n.d.	40	
Tapado	Chitonde	S4926	A.Resentini	500	56	27	16	0	0	0	0	0	0	0	0	0	1	0	100.0	57	43	0	n.d.	n.d.	n.d.	n.d.	n.d.	37	37	22	n.d.	n.d.	
	Balombo	S4927	A.Resentini	190	53	26	19	0	0	0	0	0	0	0	0	0	1	0	100.0	53	46	1	n.d.	n.d.	n.d.	n.d.	n.d.	22	42	31	n.d.	n.d.	
Culango	Culango	S4929	A.Resentini	355	57	23	18	0	0	0	0	0	0	0	0	0	1	1	100.0	58	42	0	n.d.	n.d.	n.d.	n.d.	n.d.	13	43	31	n.d.	n.d.	
	Lobito spit	S4863	A.Resentini	280	52	16	16	0	0	0	0	0	0	0	0	1	0	2	14	100.0	61	38	1	420	420	n.d.	n.d.	n.d.	11	45	29	n.d.	100
Cubal da Hanha	Praia Bebe	S4862	A.Resentini	310	70	12	10	0	0	0	0	0	0	0	0	1	0	1	6	100.0	75	23	2	327	327	93	7	0	11	49	22	n.d.	82
	Vista Alegre	S4655	G.Vezzoli	690	50	10	17	0	0	0	0	0	0	0	0	0	1	22	100.0	64	35	1	363	322	n.d.	n.d.	n.d.						

Table A6

Sample	River	Site	Analyse class µm	Sieving	wt% finer	Analyse class wt%	M.D.L. wt% coarser	0.01 SiO ₂ wt%	0.01 Al ₂ O ₃ wt%	0.04 Fe ₂ O ₃ wt%	0.01 MgO wt%	0.01 CaO wt%	0.01 Na ₂ O wt%	0.01 K ₂ O wt%	0.01 TiO ₂ wt%	0.01 P ₂ O ₅ wt%	0.01 MnO wt%	0.1 LOI wt%	0.01 Total wt%	0.1 Rb ppm	0.1 Cs ppm	1 Be ppm	0.5 Sr ppm	1 Ba ppm	1 Sc ppm	0.1 Y ppm	0.1 La ppm	0.1 Ce ppm	0.02 Pr ppm	0.3 Nd ppm	0.05 Sm ppm	0.02 Eu ppm	0.05 Gd ppm	0.01 Tb ppm	0.05 Dy ppm	0.02 Ho ppm	0.03 Er ppm	0.01 Tm ppm	0.05 Yb ppm	0.01 Lu ppm	0.2 Th ppm	0.1 U ppm
RIVERS & BEACHES OF THE CONGO																																										
S4053	Loussoni	St. Paul	63-2000	dry	0%	98%	2%	98.3	0.3	0.7	0.0	0.0	<0.01	0.0	0.1	0.02	<0.01	0.5	99.9	2	0	<1	4	19	<1	2	5	6	1	3	1	0.1	1	0.1	0	0.1	0	0.1	0	0.0	1	0
S 162		Point Noire	63-2000	wet	0%	99%	1%	98.2	0.5	0.2	0.0	0.2	<0.01	0.0	0.0	0.04	<0.01	0.8	99.9	1	<0.1	<1	13	27	<1	2	2	4	0	2	0	0.1	0	0.0	0	0.1	0	0.0	0	0.0	1	0
S0161	Loémé	Djeno	63-2000	dry	0%	99%	1%	99.0	0.2	0.1	0.0	0.0	0.0	0.0	0.2	0.02	<0.01	0.4	99.9	0	0	<1	2	6	<1	2	1	1	0	1	0	0.0	0	0.1	0	0.1	0	0.1	0	0.1	1	0
S 3533	Congo	Brazzaville	63-2000	dry	0%	100%	0%	98.5	0.3	0.2	0.1	0.1	0.0	0.1	0.1	0.02	<0.01	0.7	100.0	3	0	<1	7	38	<1	2	3	6	1	2	0	0.1	0	0.1	0	0.1	0	0.0	0	0.0	1	0
S 4898		Soyo estuary	63-2000	wet	0%	100%	0%	96.2	1.4	0.5	0.1	0.1	0.1	0.2	0.1	0.02	0.01	1.4	100.0	4	0	1	19	94	<1	2	3	5	1	2	0	0.1	0	0.1	0	0.1	0	0.0	0	0.0	1	0
S 4899		Soyo	63-2000	wet	0%	100%	0%	96.6	1.1	0.4	0.0	0.1	0.0	0.0	0.2	0.02	<0.01	1.7	100.0	1	<0.1	<1	6	15	<1	2	2	4	0	2	0	0.1	0	0.0	0	0.0	0	0.0	1	0		
RIVERS OF NORTHERN ANGOLA																																										
S 4904	Luculu	Quivanda	63-2000	wet	1%	99%	0%	91.9	3.8	1.3	0.2	0.6	0.4	0.6	0.3	0.05	0.04	0.7	99.9	9	0	<1	86	242	4	10	9	19	2	8	2	0.3	1	0.3	2	0.4	1	0.2	1	0.2	3	1
S 4897	Mebridege	Cassa de Telha	63-2000	wet	0%	96%	4%	95.2	1.4	0.9	0.1	0.1	0.1	0.3	0.1	0.03	<0.01	1.9	100.0	6	<0.1	<1	14	68	<1	3	4	7	1	4	1	0.2	1	0.1	1	0.1	0	0.1	0	0.0	1	0
S 4894	Loge	Porta Freitas Morna	63-2000	wet	0%	100%	0%	84.7	7.0	1.7	0.4	1.9	1.5	0.9	0.2	0.02	0.02	1.6	99.9	15	<0.1	<1	307	501	7	7	12	24	3	12	2	0.6	2	0.3	1	0.3	1	0.1	1	0.1	2	0
S 4891	Dande	Caxito	63-2000	wet	2%	98%	0%	87.0	5.9	1.6	0.4	0.6	1.1	1.4	0.3	0.03	0.02	1.2	99.6	33	0	3	69	435	3	7	11	22	2	9	2	0.3	1	0.2	1	0.3	1	0.1	1	0.1	3	1
S 4911	Bengo	Quifangondo	63-2000	wet	3%	96%	1%	88.8	3.3	0.8	0.2	2.7	0.5	0.9	0.1	0.05	0.01	2.5	100.0	16	0	1	139	319	1	4	5	11	1	5	1	0.2	1	0.1	1	0.1	0	0.1	0	0.1	1	1
BEACHES OF NORTHERN ANGOLA																																										
S 4902		Quifuma	63-2000	wet	0%	100%	0%	94.9	1.6	0.8	0.1	0.9	0.2	0.3	0.1	0.07	<0.01	1.0	99.9	4	<0.1	<1	64	140	1	3	4	8	1	3	1	0.1	1	0.1	0	0.1	0	0.1	0	0.1	1	0
S 4905		Mbua-Moyo	63-2000	wet	0%	100%	0%	89.0	3.9	1.1	0.3	2.3	0.7	0.7	0.2	0.18	0.02	1.5	99.9	10	<0.1	<1	202	329	4	8	9	17	2	8	2	0.4	2	0.2	1	0.3	1	0.1	1	0.1	2	1
S 4906		N'Zeto	63-2000	wet	0%	100%	0%	85.5	4.8	1.1	0.4	3.2	1.1	0.8	0.1	0.31	0.02	2.4	99.9	11	<0.1	<1	271	482	4	8	8	15	2	7	1	0.4	1	0.2	1	0.3	1	0.1	1	0.1	1	2
S 4908		Musserra	63-2000	wet	0%	100%	0%	86.4	5.3	1.6	0.3	2.4	1.1	0.8	0.2	0.04	0.04	1.6	99.8	11	<0.1	<1	292	468	7	12	13	28	3	12	2	0.6	2	0.3	2	0.4	1	0.2	1	0.2	3	1
S 4909		Ambriz	63-2000	wet	0%	100%	0%	84.3	4.8	1.1	0.4	4.4	0.8	0.8	0.2	0.05	0.05	3.1	99.9	11	<0.1	<1	299	504	6	13	9	18	2	9	2	0.5	2	0.3	2	0.5	1	0.2	2	0.2	1	1
CUANZA & CUANZA LITTORAL CELL																																										
S4652		Ilha de Luanda	63-2000	dry	0%	100%	0%	90.1	3.6	0.8	0.1	1.6	0.6	1.8	0.1	0.05	<0.01	1.2	99.9	31	0	<1	116	642	<1	2	4	6	1	3	0	0.2	0	0.1	0	0.1	0	0.0	0	0.1	1	0
S3764		Mussulo	63-2000	dry	0%	100%	0%	88.9	4.3	0.9	0.1	1.6	0.7	2.3	0.0	0.06	<0.01	1.2	99.9	44	1	<1	116	776	<1	2	5	8	1	3	1	0.2	0	0.1	0	0.1	0	0.0	0	0.0	1	0
S 4915	Cuanza	Barra do Cuanza	63-2000	wet	0%	100%	0%	93.5	2.3	0.5	0.1	0.5	0.2	0.8	0.1	0.03	<0.01	1.8	100.0	16	<0.1	<1	41	283	<1	3	4	7	1	4	1	0.1	1	0.1	1	0.1	0	0.1	0	0.1	2	0
RIVERS OF CENTRAL ANGOLA																																										
S 4917	Longa	Calamba	63-2000	wet	4%	87%	8%	88.1	3.8	0.8	0.3	2.1	0.4	1.6	0.2	0.03	0.02	2.5	99.9	28	0	<1	92	482	3	7	12	24	3	10	2	0.3	1	0.2	1	0.2	1	0.1	1	0.1	5	1
S 4922	Queve	Cachoeira	63-2000	wet	0%	100%	0%	85.4	7.1	0.8	0.1	0.5	0.8	3.3	0.2	0.02	0.02	1.8	99.9	69	0	1	104	865	1	3	6	13	1	5	1	0.2	1	0.1	0	0.1	0	0.1	0	0.1	3	1
S 4925	Quicombo	Quicombo	63-2000	wet	1%	98%	1%	83.3	7.2	1.1	0.2	0.9	0.7	3.9	0.2	0.07	0.02	2.1	99.9	75	1	<1	103	1132	2	4	11	22	2	9	1	0.3	1	0.1	1	0.1	0	0.1	0	0.1	7	1
S 4927	Balombo	Canjala	63-2000	wet	0%	94%	5%	80.7	9.1	1.2	0.2	0.3	0.7	5.3	0.4	0.03	0.03	1.9	99.8	109	1	3	87	1329	3	4	12	27	3	9	1	0.3	1	0.2	1	0.2	1	0.1	1	0.1	9	2
BEACHES OF CENTRAL ANGOLA																																										
S 3765		Cabo Ledo	63-2000	wet	0%	100%	0%	76.9	5.7	2.4	0.6	6.0	0.8	2.2	0.5	0.18	0.05	4.3	99.8	42	0	<1	272	2811	8	12	15	34	4	15	3	0.6	2	0.3	2	0.4	1	0.2	1	0.2	13	2
S 4918		Port Amboim	63-2000	wet	0%	100%	0%	84.2	5.6	1.9	0.6	2.8	0.8	2.0	0.4	0.24	0.04	1.1	99.9	38	0	<1	192	658	7	11	17	39	5	18	3	0.8	3	0.4	2	0.4	1	0.2	1	0.2	11	2
S 4919		Sumbe	63-2000	wet	0%	100%	0%	84.3	6.3	1.5	0.3	1.9	0.8	3.1	0.3	0.12	0.03	1.2	99.8	62	0	<1	162	911	5	7	12	26	3	11	2	0.4	2	0.2	1	0.3	1	0.1	1	0.2	14	2
COPOROLO, CATUMBELA & LITTORAL CELLS																																										
S 4934		Lobito spit	63-2000	wet	0%	100%	0%	79.7	8.1	3.4	0.6	2.1	1.4	2.8	0.9	0.07	0.06	0.6	99.8	57	0	<1	243	870	10	14	35	74	8	30	5	0.8	4	0.5	3	0.5	2	0.2	2	0.3	15	2
S 4655	Catumbela	Praia Bebe	500-2000	dry	0%	30%	70%	80.5	6.7	5.2	0.7	1.8	1.2	2.1	1.0	0.05	0.06	0.5	99.8	42	0	1	179	771	7	6	12	22	3	10	2	0.5	2	0.2	1	0.3	1	0.1	1	0.1	3	1
S 4933	Catumbela	Catumbela	63-2000	wet	0%	98%	1%	87.2	6.1	0.9	0.2	0.7	0.8	2.7	0.2	0.03	0.02	1.2	99.9	56	1	<1	106	798	2	3	6	14	1	5	1	0.2	1	0.1	0	0.1	0	0.0	0	0.1	3	1
S 4932	Cavaco	Benguela	63-2000	wet	0%	90%	10%	82.3	9.3	0.7	0.1	1.2	1.7	3.4	0.1	0.04	0.01	0.9	99.8	63	0	<1	250	1159	1	3	10	18	2	6	1	0.5	1	0.1	0	0.1	0	0.0	0	0.1	2	0
S 4936	Coporolo	Dombe Grande	63-2000	wet	1%	99%	0%	82.5	8.8	0.8	0.1	0.8	1.2	4.0	0.2	0.02	0.02	1.4	99.9	87	1	<1	175	1059	2	4	8	17	2	6	1	0.2	1	0.2	1	0.2	0	0.1	0	0.1	3	1
RIVERS OF SOUTHERN ANGOLA																																										
S 4940	Cangala	Santa Maria	63-2000	wet	0%	100%	0%	79.0	11.1	1.0	0.2	1.4	2.4	3.7	0.2	0.04	0.02	0.7	99.9	74	1	<1	320	1168	3	6	16	27	3	11	2	0.5	1	0.2	1	0.2	1	0.1	1	0.1	3	1
S 4945	Inamangando	Inamangando	63-2000	wet	1%	96%	4%	79.6	10.2	1.4	0.3	1.9	2.0	3.0	0.2	0.06	0.02	1.2	99.8	59	1	<1	354	987	6	6	18	33	4	14	2	0.6	2	0.3	1	0.3	1	0.1	1	0.1	5	1
S 4947	Bentiaba	Bentiaba	63-2000	wet	0%	96%	3%	79.8	10.7	0.9	0.2	1.2	2.1	4.0	0.1	0.02	0.02	0.9	99.8	81	1	<1	278	1110	2	4	11	21	2	9	1	0.4	1	0.1	1	0.1	0	0.1	0	0.1	3	0
S 4951	Giraul	Giraul	63-2000	wet	1%	97%	2%	80.4	9.7	1.4	0.3	1.3	1.7	3.7	0.2	0.03	0.03	1.2	99.9	97	1	<1	255	878	3	6	14	25	3	12	2	0.5	2	0.								

Appendix B

Click here to download Table: Appendix B Angola Zircon.xls

Sample S4899

Congo estuary @ Soyó

06°07'52" S 12°22' 25" E 96 grain analysed

70 concordant ages

grain	concentrations			isotopic ratios				ages					discordance		preferred age				
	U [ppm]	Pb [ppm]	Th/U	Pb207/Pb206	2σ 76	Pb207/U235	2σ 75	Pb206/U238	2σ 68	age 206/238	2σ age 68	age 207/235	2σ age 75	age 207/206	2σ age 76	Δ 68-75 [%]	Δ 68-76 [%]	age	2σ age
X4899_G001	66.8	7.2	0.4933	0.06015	0.0026	0.83537	0.03496	0.10076	0.00256	618.9	15	616.6	23.2	609	93.4	0.4	1.6	618.9	15
X4899_G002	44.3	20.1	0.9887	0.12093	0.0036	6.03036	0.17956	0.00688	0.00688	1990.7	41	1980.2	33.4	1970	53	0.5	1.1	1970	53
X4899_G003	190.5	78.9	0.7495	0.12353	0.0033	5.9634	0.1596	0.00786	0.00786	1935.8	37.6	1970.5	30.4	2007.8	47.4	-1.8	-3.6	2007.8	47.4
X4899_G004	151.4	28.8	0.6185	0.07311	0.0022	1.71571	0.05118	0.17028	0.00339	595.9	21.4	1014.4	24.2	1017	61	-0.1	-0.3	1013.5	21.4
X4899_G005	188.9	103.9	0.5412	0.23891	0.00608	16.31767	0.41926	0.49555	0.01096	2594.6	47.2	2895.6	32.2	3112.7	40.6	-10.4	-16.6		
X4899_G006	279	23.4	0.2793	0.05939	0.00184	0.68404	0.02094	0.08357	0.0019	517.4	11.4	529.2	15.8	581.4	67.2	-2.2	-1.1	517.4	11.4
X4899_G007				0.82317	0.0387	450.67239	48.01494	3.97218	0.4299										
X4899_G008	217.4	24.8	0.6058	0.06881	0.0021	0.97938	0.02936	0.10326	0.00234	633.5	13.6	693.3	19.2	893	63	-8.6	-29.1	633.5	13.6
X4899_G009	106.3	63	0.7073	0.17606	0.00468	11.93279	0.31998	0.49174	0.0112	2578.2	48.4	2599.1	32.8	2616.1	44.2	-0.8	-1.4	2616.1	44.2
X4899_G010	1148.8	102.1	0.8998	0.54374	0.01378	8.36508	0.21118	0.11162	0.00244	682.1	14.2	2271.4	30.4	4364.4	37.2	-7.0	-8.4		
X4899_G011	260.8	139.5	0.6076	0.1814	0.0074	11.60226	0.33051	0.46108	0.01036	2457.4	45.6	2572.9	32.2	2665.6	43.2	-4.5	-7.8		
X4899_G012	135.6	15.3	1.033	0.05771	0.00204	0.72619	0.02514	0.0913	0.00214	563.2	12.6	554.3	18.2	518.8	77.6	1.6	8.6	563.2	12.6
X4899_G013	131.6	72	1.3843	0.12932	0.00356	7.01963	0.19418	0.39383	0.00904	2140.6	41.8	2113.9	31.8	2088.7	48.4	1.3	2.5	2088.7	48.4
X4899_G014	94.8	9.5	0.6306	0.05868	0.00232	0.73287	0.02816	0.09061	0.00222	559.1	13.2	558.2	20	555.2	86.2	0.2	0.7	559.1	13.2
X4899_G015	16.6	8	0.7369	0.13699	0.00484	7.65947	0.27108	0.40566	0.01124	2195	51.6	2191.9	40.4	2189.5	61.4	0.1	0.3	2189.5	61.4
X4899_G016	79.4	8.8	0.4513	0.06046	0.00252	0.87318	0.03534	0.10478	0.00264	642.4	15.4	637.3	23	620.1	90	0.8	3.6	642.4	15.4
X4899_G017	43.1	4.6	0.3457	0.06011	0.00306	0.85741	0.0421	0.10349	0.00284	634.8	16.6	628.7	27.2	607.5	110	1	4.5	634.8	16.6
X4899_G018	260.8	139.5	0.6076	0.1814	0.0074	11.60226	0.33051	0.46108	0.01036	2457.4	45.6	2572.9	32.2	2665.6	43.2	-4.5	-7.8		
X4899_G019	273.1	154.9	0.801	0.1767	0.00452	11.35759	0.2925	0.46634	0.01022	2467.5	45	2552.9	31.4	2622.1	42.6	-3.3	-5.9	2622.1	42.6
X4899_G020	226.1	13.1	2.4728	0.3015	0.01016	1.66518	0.05134	0.04007	0.00106	253.3	6.6	995.4	27.2	3477.9	52.2	-74.6	-92.7		
X4899_G021	720.8	156.4	1.54	0.28769	0.0074	7.33331	0.18838	0.18494	0.00406	1093.9	22	2152.9	30.2	3405.2	40	-49.2	-67.9		
X4899_G022	307.9	37.2	1.1472	0.16054	0.00436	2.57032	0.06902	0.11616	0.0026	708.4	15	1292.2	25.8	2461.4	45.8	-45.2	-71.2		
X4899_G023	75.5	47.8	0.8969	0.18146	0.00488	12.51686	0.33934	0.50047	0.01146	2615.8	49.2	2644	33.2	2666.3	44.6	-1.1	-1.9	2666.3	44.6
X4899_G024	398.4	44.8	1.5353	0.05624	0.00172	0.62255	0.01888	0.08031	0.0018	498	10.8	491.4	14.8	461.9	67.8	1.3	7.8	498	10.8
X4899_G025	129.6	125.3	0.4934	0.21106	0.0036	4.86304	0.1296	0.16717	0.00364	2061.9	42.2	2096.6	39.6	2195.9	48	-1.5	-0.7	2195.9	48
X4899_G026	176.7	20.3	0.8925	0.08684	0.00214	0.84988	0.02788	0.09658	0.00224	594.3	13.2	624.6	19	735.3	71	-4	-19.3	594.3	13.2
X4899_G027	148.2	88.2	0.2682	0.06621	0.0069	21.10511	0.55156	0.57519	0.01282	2929.1	52.4	3143.4	33.2	3283.8	40.8	-6.8	-10.8	3283.8	40.8
X4899_G028	108.7	15.3	0.1153	0.0922	0.00298	1.85494	0.05886	0.14597	0.00346	878.3	19.4	1065.2	26.4	1471.5	61.4	-17.5	-40.3		
X4899_G029	214.6	46.6	0.3716	0.09458	0.00262	2.70771	0.0746	0.20771	0.00464	1216.6	24.8	1330.6	26.4	1471.5	61.4	-8.6	-19.9		
X4899_G030	55.7	14.7	0.7123	0.08946	0.00306	2.81427	0.09446	0.22823	0.00562	1325.2	29.4	1359.3	31.6	1414	65.4	-2.5	-6.3	1414	65.4
X4899_G031	241.1	22.4	0.1923	0.05673	0.00182	0.74133	0.02344	0.09481	0.00216	583.9	12.8	563.2	17	481.1	70.8	3.7	21.4	583.9	12.8
X4899_G032	167.1	49.9	0.4934	0.21106	0.0036	4.86304	0.1296	0.16717	0.00364	2061.9	42.2	2096.6	39.6	2195.9	48	-1.5	-0.7		
X4899_G033	885.6	95.9	0.226	0.27058	0.0071	3.98233	0.10388	0.10388	0.00236	854	13.8	1630.6	27.8	3309.4	41.2	-58.9	-80.2		
X4899_G034	130	39.6	0.2333	0.1118	0.00306	4.86187	0.1331	0.3155	0.00708	1767.7	34.6	1795.7	29.8	1828.9	49.6	-1.6	-3.3	1828.9	49.6
X4899_G035	339.5	77.3	0.5411	0.07956	0.00218	2.26739	0.06618	0.20677	0.00458	1211.6	24.4	1202.2	24.8	1186.1	54.2	0.8	2.1	1186.1	54.2
X4899_G036	106.7	19.6	0.3745	0.0802	0.0025	1.92781	0.05908	0.1744	0.00404	1036.3	22.2	1090.8	26.8	1201.9	61.4	-5	-13.8	1036.3	22.2
X4899_G037	231.6	11.1	0.3655	0.04996	0.00194	0.32031	0.01214	0.04652	0.0011	293.1	6.8	282.1	11.2	193.1	90.2	0.9	5.8	293.1	6.8
X4899_G038	45.4	8.8	0.1742	0.07036	0.00286	1.49426	0.05894	0.15408	0.00396	923.8	22.2	928.1	29.2	938.9	83.4	-3.5	-1.6	923.8	22.2
X4899_G039	132	54.3	0.7044	0.12279	0.00334	5.95212	0.16168	0.35168	0.00788	1942.6	37.6	1968.9	30.6	1997.2	48.4	-1.3	-2.7	1997.2	48.4
X4899_G040	71.5	51.2	0.7458	0.19446	0.00314	1.51752	0.04308	0.15444	0.01286	2895	45.4	2847.6	33.6	2821.8	43.8	-1.3	2.2	2821.8	43.8
X4899_G041	1026.7	74.8	0.7556	0.31737	0.00846	2.89234	0.07612	0.06126	0.00146	412.7	8.8	1379.9	26.2	3557.1	41	-70.1	-81.7		
X4899_G042	35.2	4.4	0.4933	0.06564	0.0032	1.06511	0.05	0.11732	0.00322	715.1	18.6	734.6	29.2	794.9	102.2	-2.6	-1.0	715.1	18.6
X4899_G043	90.9	55	1.1224	0.2685	0.00716	17.51734	0.469	0.47334	0.00322	2498.2	47	2963.6	33.6	3297.3	41.8	-15.7	-24.2		
X4899_G044				0.74766	0.0846				321.57934										
X4899_G045	30.8	21.7	0.9743	0.19701	0.00596	14.73197	0.45236	0.54251	0.01139	2793.9	58	2798.1	37.6	2801.6	49.4	-0.1	-0.3	2801.6	49.4
X4899_G046	50.3	12.5	0.5125	0.07391	0.0028	1.67535	0.06176	0.16446	0.0041	981.5	22.6	999.2	28.8	1039	76.4	-1.8	-5.5	981.5	22.6
X4899_G047	81	8.3	0.3745	0.05961	0.00274	0.81776	0.03644	0.09953	0.0026	611.7	15.2	606.8	24.2	589.4	96.0	0.8	3.8	611.7	15.2
X4899_G048	200.4	38.6	0.4363	0.09616	0.0028	2.36781	0.06824	0.17864	0.00406	1059.5	22.2	1232.9	26.4	1509.9	54.6	-14.1	-31.7	1059.5	22.2
X4899_G049	95.6	54	0.4761	0.16836	0.00468	11.45273	0.32008	0.49351	0.01142	2585.8	49.2	2585.8	33.8	2541.4	46.6	1	1.7	2541.4	46.6
X4899_G050	142.3	24.9	0.3189	0.07278	0.00232	1.71027	0.05354	0.1705	0.00396	1014.9	21.8	1012.4	25.2	1007.8	64.6	0.2	0.7	1014.9	21.8
X4899_G051	82.6	9.3	0.4314	0.06167	0.00266	0.90548	0.03786	0.10652	0.00272	652.5	15.8	654.7	24.2	662.7	92.4	-0.3	-1.5	652.5	15.8
X4899_G052	100.4	43	0.715	0.14454	0.00426	7.19678	0.21138	0.36122	0.00856	1988	40.6	2136.1	33.8	2282.3	50.8	-6.9	-12.9	2282.3	50.8
X4899_G053	66.8	69.8	0.6496	0.03512	0.00182	0.84988	0.02788	0.09658	0.00224	594.3	13.2	624.6	19	735.3	71	-4	-19.3		
X4899_G054	128	46.2	0.7034	0.1126	0.00332	4.81866	0.1408	0.31047	0.0072	1743	35.4	1788.2	31.6	1841.8	53.4	-2.5	-5.4	1841.8	53.4
X4899_G055	39.1	26.8	1.2681	0.18244	0.00554	12.62945	0.38666	0.50223	0.01264	2623.4	54.2	2652.4	37.2	2675.2	50.2	-1.1	-1.9	2675.2	50.2
X4899_G056	56.5	28.1	1.5875	0.11918	0.0037	5.7333	0.177	0.34901	0.00844	1929.9	40.4	1936.4	34	1944	55.6	-0.3	-0.7	1944	55.6
X4899_G057	103.9	16.8	0.531	0.09831	0.00324	1.99817	0.06416	0.14745	0.00352	886.7	19.8	1114.9	27.6	1592.3	61.6	-20.5	-44.3		
X4899_G058	20.6	4.5	0.6126</																