

Carbonate drowning successions of the Bird's Head, Indonesia

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Abstract Drowning unconformities and their related strata are important records of key tectonic and environmental events throughout Earth's history. In the eastern Bird's Head region of West Papua, Indonesia, Middle Miocene strata record a drowning unconformity present over much of western New Guinea, including several offshore basins. This study records platform carbonate strata overlain by mixed shallow- and deep-water units containing benthic and planktonic foraminiferal assemblages in several outcrop locations across the eastern Bird's Head region. These heterolithic beds are interpreted as drowning successions that are terminated by a drowning unconformity. We define a succession exposed along the Anggrisi River in the eastern Bird's Head as a stratotype for carbonate platform drowning in the Bird's Head, analogous to similar faunal turnovers identified in its offshore basins. Detailed facies analyses, biostratigraphic dating, and paleoenvironmental interpretations using larger benthic and planktonic foraminifera collected from the Anggrisi River succession help to constrain the drowning event recorded onshore as beginning in the Burdigalian and ending in the Serravallian. The cause of platform drowning in the Bird's Head is attributed to a reduction in the rates of carbonate accumulation due to the presence of excess nutrients in the depositional

environment. Already foundering carbonate platforms due to environmental deterioration were left vulnerable to submergence and eventually succumbed to drowning. Low rates of carbonate production were outpaced by the rate of relative sea-level rise caused by high-amplitude oscillations in global glacio-eustatic sea-level change and/or regional tectonic subsidence. The duration of the drowning event across the entire Bird's Head region is interpreted to have lasted a duration of approximately 9.5 My, between 18.0 and 8.58 Ma. This has implications when interpreting timings of sedimentary basin fill across western New Guinea and in other basins where carbonate platform drowning is recorded.

Keywords Platform drowning · Drowning unconformity · Biostratigraphy · Carbonates · Foraminifera · Indonesia

Introduction

The concept of drowning unconformities is well established and there are many examples of carbonate platform drowning strata documented in the rock record worldwide (e.g., Schlager 1981, 1989; Hallock and Schlager 1986; Longman et al. 1987; Simone and Carannante 1988; Erlich et al. 1990, 1993; Föllmi et al. 1994; Drzewiecki and Simo 1997; Weissert et al. 1998; Blomeier and Reijmer 1999; Wortmann and Weissert 2000; Wissler et al. 2003; Ruiz-Ortiz et al. 2004; Mutti et al. 2005; Föllmi and Gainon 2008; Sattler et al. 2009; Najarro et al. 2011; Marino and Santantonio 2010; Brandano et al. 2016; Sulli and Interbartolo 2016). Carbonate platforms are defined as flat-topped accumulations of carbonate sediments developed at or very near sea level (Hallock and Schlager 1986). These platforms are complex natural

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systems that are sensitive to a number of environmental factors that may contribute to drowning. Environmental stresses that may lead to carbonate platform demise and the formation of drowning unconformities include oceanic anoxic events, tectonism and flexural loading, hyper- and hypo-salinity, change in carbon dioxide content of the oceans or atmosphere, change in nutrient supply, death of zooxanthellae, temperature, hydrodynamics, turbidity, light penetration, and evolutionary modification (Schlager 1991; Thorne 1992; Hallock and Schlager 1986; Erlich et al. 1990, 1993; Mutti et al. 2005). In all cases, however, platform drowning occurs because the rate of platform top production and accumulation is exceeded by the rate of relative sea-level rise, either because the rate of relative sea-level rise has increased, or because the rate of carbonate production has decreased. The process of carbonate platform drowning presents somewhat of a paradox as rates of carbonate production and accumulation normally exceed rates of long-term sea-level rise and subsidence; glacioeustatic sea-level rises, on the other hand, can exceed carbonate production (Schlager 1981; Hallock and Schlager 1986; Föllmi et al. 1994; Wissler et al. 2003; Ruiz-Ortiz et al. 2004; Mutti et al. 2005; Brandano et al. 2016). Kim et al. (2012) demonstrated that even if the rate of relative-sea level rise is lower than maximum carbonate potential platform drowning can still occur if accumulation rates at the initiation of drowning are lower than the rate of relative sea-level rise. This can be induced by slow response times of carbonate-producing marine benthic organisms to re-colonize the platform top after flooding. The time required for biotic colonization controls the increase in water depth which, in turn, determines carbonate production in a keep-up mode that is directly followed by drowning (Kim et al. 2012). These slow response times cause negative feedback in the rates of carbonate accumulation, which decrease exponentially with increasing water depth. This scenario serves as a possible explanation for the ‘drowning paradox’, where there is no evidence for significant external forcing, such as tectonic activity or glacio-eustatic changes (Kim et al. 2012). A drowning unconformity may be produced above a drowned platform through sediment starvation leading to a condensed section or progressive onlap by aggrading strata (Bosence et al. 1998; Ruiz-Ortiz et al. 2004) or progradation of siliciclastic strata, abruptly terminating carbonate production (Schlager 1989; Erlich et al. 1990, 1993).

In this paper, we present outcrop evidence analogous to interpreted drowning successions occurring in sedimentary basins from on- and offshore New Guinea. We provide ages and a causal interpretation for the drowning event of New Guinea by comparing sedimentological and biostratigraphic data from outcrop samples and review of well data.

Physical characteristics of drowning strata

Features of successions deposited during relative sea-level driven drowning may include a gradational lower and upper contacts, chemical sedimentation (e.g., glauconite, phosphate, ferruginous crusts), open-marine shelf sediments, a variable hiatus at the upper boundary, late-stage shelf-margin growth and sub-horizontal basinal marine onlap (Erlich et al. 1990, 1993). Drowning events can appear in outcrop to occur abruptly due to a sudden shutdown of carbonate platforms through environmental deterioration. However, they may also occur gradually, as younger platforms back-step conformably to continue accumulating in areas with more favorable environmental conditions (Erlich et al. 1990, 1993). In most scenarios, the drowning sequence appears gradational and typically includes a change in lithology from shallow-water carbonates to deeper-water siliciclastics or pelagic sediments of variable composition. These strata define a deepening-upwards sequence. During the drowning process, as water depths increase, shallow and deep-water lithologies may mix, producing short ‘drowning successions’ of heterolithic beds (Schlager 1989). These interbedded sequences are intermediate between platform carbonates and fully deep-water sediments, which contain a mix of benthic and planktonic carbonate factory products, marking the inception and completion of the drowning process (Marino and Santantonio 2010).

Seismic expression of drowning strata

In seismic reflection data, drowning unconformities have a similar seismic character to subaerial unconformities (type 1 sequence boundaries). These features include high-amplitude reflections produced by the contrast in acoustic impedance between the deep- and shallow-water facies above and below the boundary, or presence of condensed sections, and both display marine onlap at the base of the platform margin (Schlager 1989). In traditional sequence stratigraphic interpretations, drowning unconformities are usually considered equivalent to marine flooding surfaces and may occupy opposite positions in the relative sea-level cycle to type 1 sequence boundaries (Van Wagoner et al. 1988; Erlich et al. 1990, 1993; Catuneanu 2006; Masse and Fenerci-Masse 2011). Schlager (1999) sought to define drowning unconformities as a third major boundary type, a ‘type 3 sequence boundary’, prone to forming long marine hiatuses and exhibiting high reflectivity in seismic sections. Drowning unconformities can be distinguished from other sequence boundaries by the retrogradational stacking pattern of carbonate platform margins that typically signify back-stepping (Burgess et al. 2013), in contrast to the traditional incised valleys, coastal onlap, and forced regressive surfaces of type 1 sequence boundaries. Since carbonate strata

backstep during drowning, the drowning unconformity can form diachronously across the platform top and will show a variable hiatus, depending on when and where subsequent deposition of deeper-water strata begins (e.g., Föllmi et al. 1994).

Geological history

The Pacific island of New Guinea is often described as having the appearance of a bird in map view. Papua New Guinea comprises the ‘Bird’s Tail’ in the east, and the Indonesian provinces of West Papua and Papua comprise the ‘Bird’s Head’, ‘Neck’ and ‘Shoulder’ to the west. The ‘Bird’s Neck’ forms the Cenderawasih Bay embayment, to the north of which are located the islands of Yapen, Biak, and Supiori (Fig. 1). Several sedimentary basins are found

within the Bird’s Head, namely the Bintuni, North Irian, and Salawati Basins (Fig. 1). The Salawati and Bintuni Basins, to the west of the Bird’s Head, are known to produce hydrocarbons. The Bird’s Head region is a region of active tectonic deformation that is the result of interaction between the Eurasian, Pacific, and Australian plates, as well as several smaller microplates (Visser and Hermes 1962; Dow and Sukanto 1984; Pieters et al. 1983; Cloos et al. 2005), dissected by several major, mainly strike-slip, fault zones (Fig. 1).

The majority of New Guinea was part of the northern promontory of the Australian continent termed the Sula Spur during the Paleogene (Hall 2002, 2012). In eastern New Guinea, collision occurred between intra-Pacific island arcs formed above a northward-dipping subduction zone, and the Sula Spur during the latest Chattian–Aquitania (Dewey and Bird 1970; Dow 1977; Pigram and

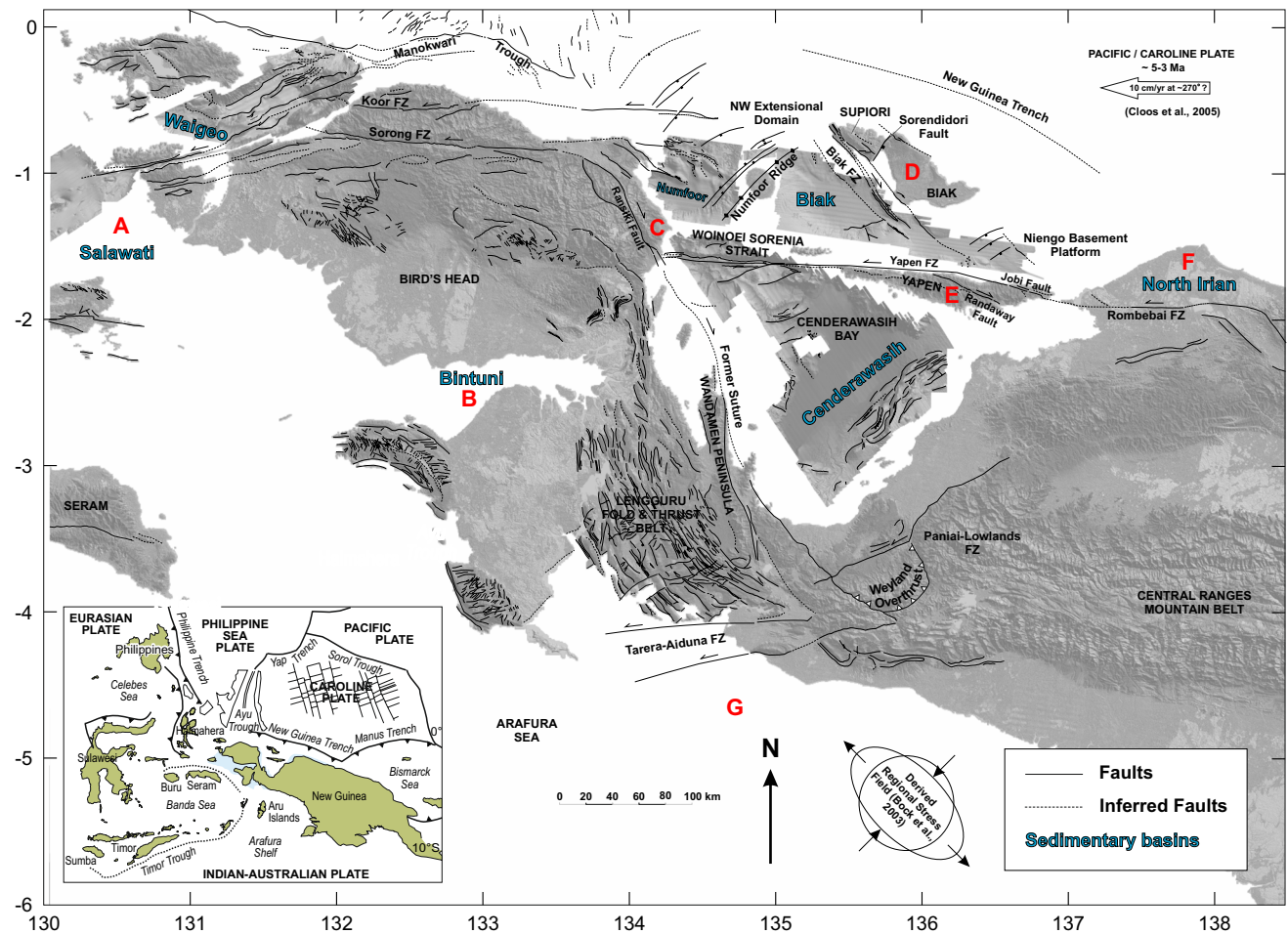


Fig. 1 Structural map of western New Guinea. Faults were drawn based on features identified from ASTER digital elevation data, bathymetric multibeam, and seismic data of the Biak and Cenderawasih Bay Basins provided by TGS, and those encountered in the field. The offshore Manokwari Trough was drawn from GLORIA

sonar imagery (after Milsom et al. 1992). Derived regional stresses are implied after Bock et al. (2003), and vector of Pacific–Caroline plate motion plotted after Cloos et al. (2005). Letters A–G correspond to stratigraphic columns of Fig. 2

Davies 1987; Milsom et al. 1992; Hall 2002, 2012; Hill and Hall 2003; Wallace et al. 2004). Evidence of this arc-continent collision can be found in the eastern Bird's Head where a proposed suture between accreted arcs and the Australian continental margin is presently represented by the Ransiki Fault, interpreted to extend into Cenderawasih Bay, east of the Wandamen Peninsula, and link with the Weyland Overthrust in the New Guinea Central Ranges (Pigram et al. 1982; Dow and Hartono 1982; Milsom 1991, 1992; Fig. 1). The island arc rocks are represented by the Arfak Volcanics east of the Ransiki Fault in the eastern Bird's Head, and the Yapen Volcanics and Auwewa Formation found on Biak and Supiori (Fig. 2). These arc rocks now form the basement to much of the northern New Guinea margin including the eastern Bird's Head, Cenderawasih Bay, and its islands (Fig. 2). Similar sequences of these rock types have been interpreted as part of an ophiolite suite (Dow and Hartono 1982; Pieters et al. 1983; Hall 2002; Hill and Hall 2003; Cloos et al. 2005), abducted on to the New Guinea Central Ranges along the Weyland Overthrust (Fig. 1).

During this time, the Sorong Fault Zone (Fig. 1) was initiated when the Philippine Sea Plate began its clockwise rotation, at approximately 20 Ma, causing its southern boundary to change from one of subduction to sinistral strike-slip (Ali and Hall 1995; Hall 2002). The Sorong Fault forms the northern margin of the Salawati Basin (Redmond and Koesoemadinata 1976; Phoa and Samuel 1986), which is dominated by east–west folds and a complex interplay of strike-slip and extensional faults (Gibson-Robinson and Soedirdja 1986; Wilson 2006).

Arc-continent collision produced a subaerial, erosional, and angular Early Miocene unconformity (Fig. 2) and several compressional structures across northern New Guinea. In the Biak Basin (Fig. 1), this unconformity is marked by a subaerial erosion surface in which folded and thrust island arc material deformed during the collision was uplifted and exposed leading to the erosional truncation of ramp anticlines with up to 820 m of section missing (Gold et al. 2014). The reduction of island arc material by erosion created platforms on which coral reefs flourished and platform carbonates developed (Wilson 2002). These

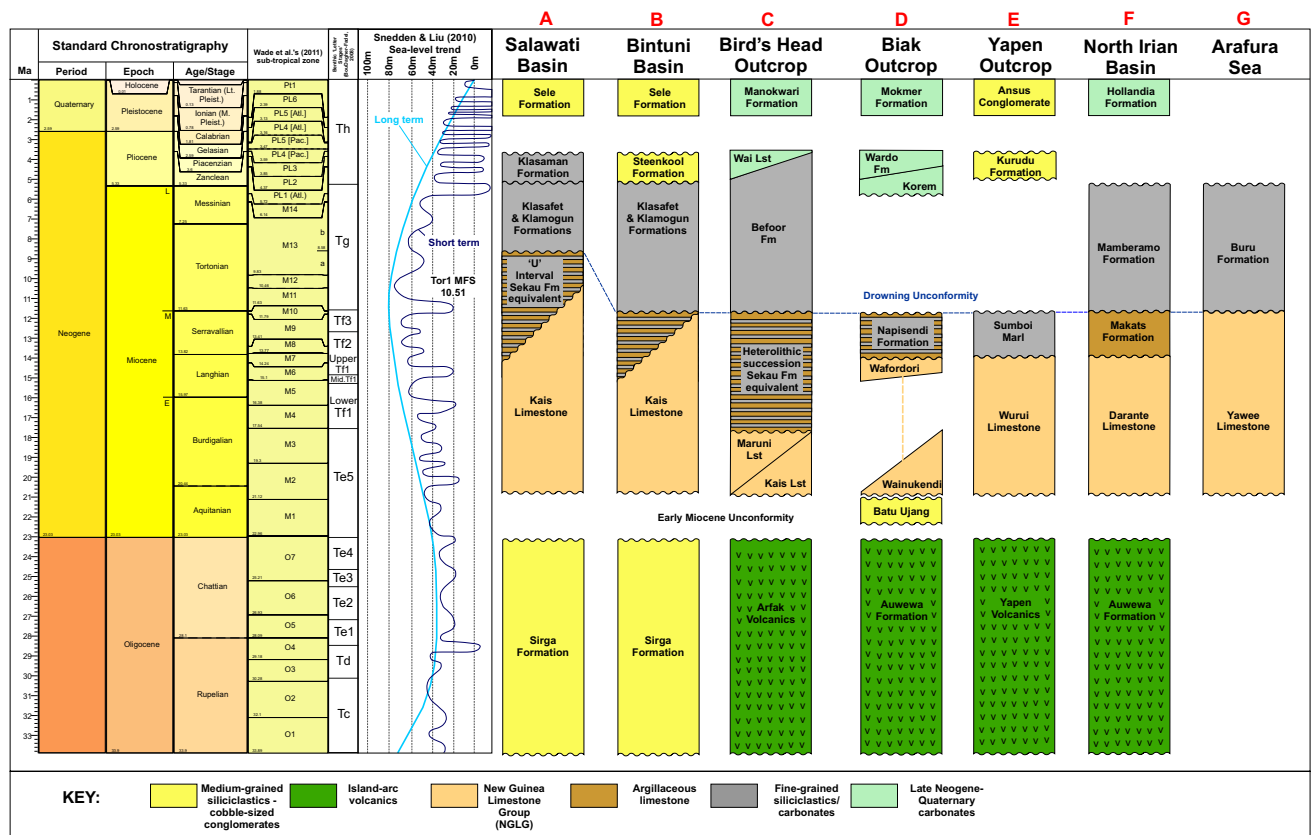


Fig. 2 Stratigraphy of the Bird's Head study area. Correlation of the stratigraphy indicates younging of the drowning unconformity from east to west. Stratigraphy derived from field and biostratigraphic data of this study and composites from listed authors (after Gibson-Robinson and Soedirdja 1986; Chevallier and Bordenave 1986; Waschmuth

and Kunst 1986; Atmawinata et al. 1989). Sea-level curve of Snedden and Liu (2010), Indo–Pacific letter stages from BouDagher-Fadel (2008), stratigraphic columns and planktonic foraminiferal zones from Wade et al. (2011) and Timescale Creator (version 6.1.2, <http://www.tscreator.org>) following the timescale of Gradstein et al. (2012)

carbonate platforms developed across the region during early to middle Miocene slow relative sea-level rise, depositing the ‘New Guinea Limestone Group’ (NGLG) that includes several contemporaneous carbonate formations (Visser and Hermes 1962; Pieters et al. 1983; Brash et al. 1991; Fig. 2). The Middle Miocene saw a significant increase in the rate of relative sea-level rise, driven by regional subsidence (Visser and Hermes 1962; Satyana 2003; Wilson 2006). During this time, carbonate platforms failed to keep pace with the rate of relative sea-level rise, attributed to the onset of extensional tectonics (Visser and Hermes 1962), and drowned.

Extension was driven by the continued convergence of the Pacific and Australian plates whereby the Pacific plate converges WSW at $\sim 248^\circ$ relative to the Australian plate (Wallace et al. 2004; Cloos et al. 2005; Fig. 1). Paleomagnetic measurements from the Bird’s Head show little evidence for rotation away from this azimuth with respect to Australia throughout the Neogene (Klootwijk et al. 1986; McCaffrey and Abers 1991). Convergence about this azimuth during the early Neogene directed maximum compressive stress towards the SW and minimum compressive stresses directed NW–SE. Similar NW–SE-directed extension is described in Seram (Fig. 1) as occurring at the same time as the Bird’s Head at approximately 15 Ma (Pownall et al. 2013, 2014). In the Bird’s Head, minimum compressive stress propagating NW–SE caused the formation of several NE–SW-striking extensional basins. On the northern New Guinea margin, these basins developed in the Niengo Basement Platform (Fig. 1), which formed as a series of horsts and grabens that developed due to the oblique motion between the Australian and Pacific plates (Williams and Amiruddin 1983; McAdoo and Haebig 1999).

The tilted half-grabens and other extensional structures in the Niengo Basement Platform area strike NE–SW (Fig. 1). Examples of these include the NE–SW-trending Sorendidori Fault and Numfoor Ridge (Fig. 1). The Sorendidori Fault is an oblique normal fault that separates and downthrows younger Neogene sediments of the island of Biak to the SE from Early to Middle Miocene carbonates of the island of Supiori to the NW (Fig. 1). Extension is interpreted to have ceased by the end of the Miocene where Pliocene–Pleistocene carbonates are draped undisturbed across extensional tilted blocks.

By the end of the Miocene, the Sorong Fault Zone ceased to be a major feature. Pliocene strata that unconformably cover the eastern end of the Sorong Fault Zone in the northern Bird’s Head are not broken by any active strike-slip faults (Hamilton 1979) and it is now considered to be inactive (Watkinson and Hall 2016). However, during the Pliocene, convergence between the Australian and Pacific plates was accommodated by other strike-slip faults

developed in the Bird’s Head region. The Ransiki Fault, originally formed as a collisional suture, was reactivated as a zone of dextral strike-slip with a recent extensional component (e.g., Watkinson and Hall 2016). The Yapen and Biak Fault Zones (Fig. 1) also became major features. The Biak Fault Zone is a series of segmented, parallel, NW–SE-trending strike-slip faults that form the linear coastline of Supiori’s peninsula, linking with the linear coast of southwest Biak (Fig. 1). These faults also form clearly expressed lineaments on the seafloor that can be observed in multibeam bathymetric data (Fig. 1). The Biak Fault Zone is shown to be a young feature by incising the youngest Pliocene sequences and recent sedimentation within the Biak Basin is controlled by activity along the faults (Bertoni and Garcia 2012; Memmo et al. 2013). Recent seismicity along the Biak Fault Zone exhibits a distinct dextral slip component along the fault zone.

The Yapen Fault forms the E–W-trending linear coastline of the north coast of Yapen Island and a submarine ridge extending westward from Yapen to the eastern Bird’s Head (Fig. 1). It separates the Woinoei Sorenia Strait in the north from a part of Cenderawasih Bay over 2000 m deep in the south. The Yapen Fault is a well-documented left-lateral structure. The fault strikes SE onshore to the east of the island of Yapen where it is termed the Jobi Fault system. The Jobi Fault in eastern Yapen (Fig. 1) has been shown to display left-lateral offset (Atmawinata and Ratman 1982; Dow and Hartono 1982; Atmawinata et al. 1989). The associated Randaway Fault system that cuts obliquely SE through the center of the island (Fig. 1) is proposed to have been active since the Pliocene or later, having incised the Ansum Conglomerate (Fig. 2), dated by foraminiferal assemblages to be of Pliocene age (Charlton 2010). This sub-recent movement of the Yapen Fault is supported by the sharp bathymetric expression of the submarine ridge on the seafloor (Fig. 1) and the occurrence of numerous sinistral strike-slip earthquakes along its course.

In the Bird’s Neck, the Lengguru Fold and Thrust Belt (LFTB) forms the westward extension of the mountains of the Central Ranges (Fig. 1). The LFTB is stated to have formed in the last 7–8 Ma from the Tortonian (Bailly et al. 2011) based on ages of metamorphism from internal rock units within the LFTB. Their results suggested that a high-pressure metamorphic event took place during the latest Tortonian to Messinian (7–8 Ma) followed by rapid exhumation associated with migmatization only 1 or 2 Ma after their burial. This rapid exhumation is able to produce high-pressure metamorphic rocks in less than 5 Ma, validating fission-track ages of metamorphic units in New Guinea that suggest rapid uplift occurred in New Guinea since 10 Ma and in many areas since 5 Ma (Hill and Gleadow 1989). This uplift continued into the Pliocene, culminating with the formation of the regional ‘intra-Pliocene unconformity’

known from within the Salawati Basin to occur at approximately 4 Ma (Decker et al. 2009).

Previously described drowning successions in the Bird's Head

Strata that may represent drowning successions have been described previously in the Bird's Head, but were not recognized as such (e.g., Visser and Hermes 1962; Pieters et al. 1983; Gibson-Robinson and Soedirdja 1986; Brash et al. 1991; McAdoo and Haebig 1999). Examples include successions of limestones and/or siliciclastic sediments with alternating benthic and planktonic fossil assemblages recorded along the northern coast of New Guinea (e.g., Visser and Hermes 1962; Gibson-Robinson and Soedirdja 1986; McAdoo and Haebig 1999).

North Irian Basin

The North Irian Basin (also termed the 'Mamberamo' or 'North Coast' Basin) is located in the Mamberamo region of the Bird's Shoulder along the northern New Guinea margin (Fig. 1). This basin is described as a single large basin with interconnected, separate depocentres sharing a correlative fill (McAdoo and Haebig 1999). The North Irian Basin formed between structural highs of the Niengo Basement Platform, comprising ocean-floor or island arc volcanic and ultramafic material, since the Paleocene (McAdoo and Haebig 1999).

McAdoo and Haebig (1999) described continuous Cenozoic deposition within the North Irian Basin with the exception of a regional depositional hiatus in the Late Miocene. This hiatus is marked by a major unconformity, found at the base of the Mamberamo Formation and the top of the Makats Formation (Fig. 2, Column F), and traceable over large areas of the basin (Visser and Hermes 1962; McAdoo and Haebig 1999). We suggest that this may be a drowning unconformity as it separates shallow-water platform carbonate from overlying deep-water marl and pelagic carbonate. This facies change is contemporaneous with other such changes across New Guinea interpreted to record regional platform drowning and can be traced into the eastern Bird's Head (Fig. 2, Column C) and islands of Biak and Supiori (Fig. 2, Column D).

Salawati and Bintuni Basins

Many authors agree that the Salawati and Bintuni Basins (Fig. 1) are built on continental crust of either the northern promontory (Sula Spur) of the Australian plate (Redmond and Koesoemadinata 1976; Chevallier

and Bordenave 1986; Gibson-Robinson and Soedirdja 1986; Phoa and Samuel 1986; Charlton 1996), or a small microcontinent comprising the Bird's Head (Pigram et al. 1982; Dow and Hartono 1982; Pigram and Panggabean 1984; Pigram and Davies 1987; Struckmeyer et al. 1993; Puntodewo et al. 1994; Hall 2002; Wilson 2006; Tikku et al. 2006).

The Bintuni Basin overlies a basement of NGLG that ceased deposition in the Middle Miocene (Pigram et al. 1982). It is interpreted to have formed as a fore-deep in the late Cenozoic (Chevallier and Bordenave 1986) when the former shallow-marine NGLG platform was rapidly deepened as a result of collision between the Bird's Head microcontinent with the Australian continent along a short-lived east-facing subduction zone (Pigram et al. 1982).

In the Salawati and Bintuni Basins, and in outcrop in the western Bird's Head, the Kais Limestone of the NGLG is overlain by marl and micritic limestone of the Klasafet and Klamogun Formations, both interpreted as pelagic strata (Visser and Hermes 1962; Fig. 2, Columns A and B). On depositionally topographic highs, the upper part of the Kais Limestone is overlain by the Sekau Formation (Visser and Hermes 1962; Koesoemadinata 1976; Pieters et al. 1983). This unit is described as comprising limestones with the same faunal composition as the Kais Limestone, interbedded with marl, clay, and argillaceous limestones occurring on topographically high coral reefs at the time of deposition (Visser and Hermes 1962). It is interpreted here that the interbedded sequences between the Kais Limestone and fully pelagic strata of the Klasafet and Klamogun Formations represent drowning successions. The drowning unconformity is therefore situated at the base of the Klasafet and Klamogun formations (Fig. 2).

In the Salawati Basin, the 'U' Interval overlies the Kais Limestone as a succession of interbedded limestone and siliciclastic strata. This heterolithic succession represents another gradual facies transition of shallow-water platform carbonates of the Kais Limestone to fully pelagic strata of the Klasafet and Klasaman Formations. The boundary between this heterolithic interval and wholly siliciclastic strata of the Upper Klasafet and Klasaman Formations is termed the 'U' Marker (Gibson-Robinson and Soedirdja 1986; Satyana 2003). The 'U' Marker is an easily recognized horizon in seismic and well logs (Valenta 1979; Gibson-Robinson and Soedirdja 1986) marking a sharp transition from mixed limestone and calcareous clays below to mainly clays above (Gibson-Robinson and Soedirdja 1986). We interpret the 'U' Interval as a drowning succession and the 'U' Marker as a drowning unconformity within the Salawati Basin (e.g., Gold et al. 2014).

Materials and methods

Fieldwork was undertaken in the eastern Bird's Head region in 2011 and 2013 during which heterolithic successions of carbonate strata were observed at several localities (Fig. 3). Heterolithic successions were observed at the top of the Maruni Limestone along the eastern coast of the Bird's Head, the Kais Limestone near the village of Anjai in the central Bird's Head (Fig. 3), and Napisendi Formation along the northern coasts of the islands of Biak and Supiori (Fig. 3). However, the best exposure occurs above the Maruni Limestone along the Anggrisi River (Sungai Anggrisi), south of the city of Manokwari (Fig. 3). Vertical facies variations within the heterolithic sections were measured and the sedimentological and paleontological features were recorded at each bed.

At each locality, a clear change in facies together with a faunal turnover from shallow-water benthic to pelagic planktonic foraminiferal assemblages was observed. A simple scheme, based on lithology and foraminiferal assemblage, was devised to record the spatio-temporal distribution of facies across the study area (Fig. 4). The facies scheme contained four categories: (1) shallow-water platform, (2) argillaceous platform, (3) periplatform, and (4) pelagic (Fig. 4). This scheme is based on carbonate rock fabrics, textures, and paleoenvironmental interpretations of foraminiferal assemblages from Bé and Tolderlund (1971), Bé (1977), Hallock and Glenn (1986), and BouDagher-Fadel (2008).

The Sungai Anggrisi section was determined to be representative of heterolithic successions in the eastern Bird's Head and was chosen for more detailed analysis. Fourteen localities along Sungai Anggrisi were visited (Fig. 5) and stratigraphic sections were measured, and sedimentological and paleontological features were recorded. Systematic sampling occurred at each facies change for detailed microfacies and biostratigraphic analyses. Thin-sections were prepared of the samples and stained with a mixture of Alizarin Red S and potassium ferricyanide (Dickson 1966) to aid in the identification of ferroan and non-ferroan phases of calcite. Thin-sections were then classified according to microfabrics and microfossil composition. During these detailed analyses, six distinct foraminiferal assemblages were identified within the facies scheme (Fig. 6). Data from all 14 localities along Sungai Anggrisi were collated to produce a composite log of the eastern Bird's Head heterolithic succession (Fig. 6).

Biostratigraphic data from outcrop samples collected from localities with a close field-relationship to the heterolithic successions depicted in Fig. 3 were used to display the full spatial distribution of facies belts across the field area. Gross depositional environment maps were produced of the field area during the early, middle, and late

Miocene, comprising 105 data points (Fig. 7). Biostratigraphic data from 15 public domain wells (Table 1) were also reviewed and identified a similar turnover of faunal assemblages and facies in areas to the west of the field area. Ages were assigned to well data and outcrop samples based on assemblages of planktonic and larger benthic foraminifera. Ranges of taxa were taken from Adams (1970), Iaccarino et al. (2007), and BouDagher-Fadel (2008) and were calibrated to the Indo-Pacific 'letter stages' (Adams 1970; BouDagher-Fadel 2008; Lunt 2013) and Wade et al. (2011) sub-tropical planktonic foraminiferal zones. Biozonation schemes used in this study are the Indo-Pacific 'letter stages' of Adams (1970), BouDagher-Fadel (2008), and Lunt (2013) for larger benthic foraminifera, and Wade et al. (2011) for sub-tropical planktonic foraminiferal zones. Sub-tropical planktonic foraminiferal zones were calibrated to Timescale Creator (version 6.1.2, <http://www.tscreator.org>) following the geological timescale of Gradstein et al. (2012).

New outcrop evidence for platform drowning in the Bird's Head

We define the section along Sungai Anggrisi (Figs. 3b–c, 5, 6) as a stratotype for carbonate platform drowning in the Bird's Head. This stratigraphic section serves as an analogue to similar successions observed in the subsurface of offshore basins to the west of New Guinea. The heterolithic succession at Sungai Anggrisi is approximately 20 m thick and displays all four facies defined in Fig. 4. Six unique foraminiferal assemblages are identified within the succession that reflect the faunal turnover from shallow-water carbonate platform deposition to wholly pelagic strata through time (Fig. 6).

Shallow-water platform facies

The basal unit of the Sungai Anggrisi succession comprises reefal platform carbonate of the Maruni Limestone (Figs. 3b, 6). This unit is well bedded, with planar beds dipping 250/10°NW, comprising white-colored and well-cemented fossiliferous limestone (Figs. 3b, 6). The limestone is dominated by larger benthic foraminifera; however, bivalve fragments, corals, and the rhodophyte *Lithophyllum* spp. are also present. Through petrographic analysis, it is classified as a mud-lean packstone comprising ~70% bioclasts, ~20% non-ferroan spar cement, and ~10% matrix. Larger benthic foraminifera within this limestone include *Eulepidina* spp. (including *E. badjirraensis*), *Lepidocyclina* (*Nephrolepidina*) *brouweri*, *Lepidocyclina* (*N.*) *oneatensis*, *Lepidocyclina* (*N.*) *sumatrensis*, *Miogypsina kotoi*, *Miogypsina tani*, *Miogypsinodella primitiva*, *Miogypsinoides* spp. (including *M.*

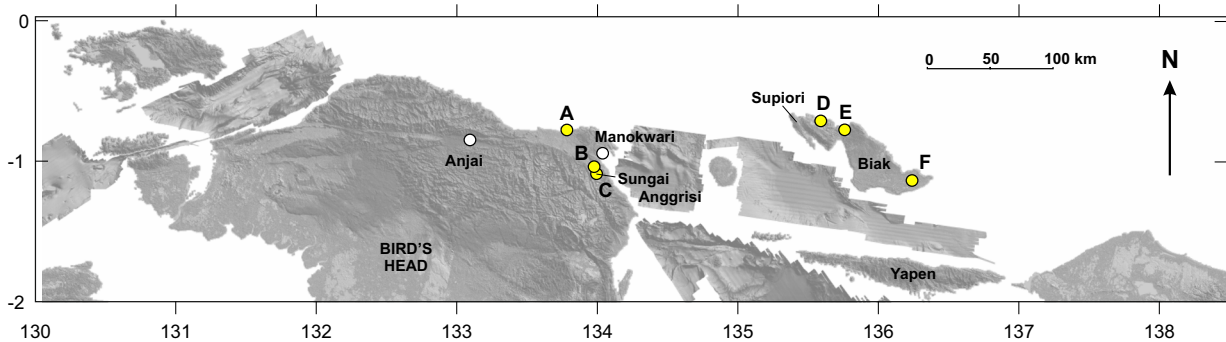


Fig. 3 Field photos of drowning successions and drowning unconformities across the northern Bird's Head. The drowning succession is marked by a sharp change in facies from platform carbonates beneath and dark shales or heterolithic grey marl and brown packstone above. **a** Befoor Formation unconformably overlies the Maruni Limestone in the north-eastern corner of the Bird's Head, north of Manokwari. **b** Drowning succession at the top of the Maruni Limestone along the river Sungai Anggrisi, south of Manokwari. **c** Heterolithic drowning succession beds observed farther along Sungai Anggrisi. **d** Deep-water facies rocks containing abundant planktonic foraminifera within the Napisendi Formation on the northern coast of Supiori. **e** Heterolithic grey marl and better indurated brown packstone within the Napisendi Formation on the island of Biak. **f** Underlying platform limestone of the Napisendi Formation beneath dark grey shales in the south-east of the island of Biak

dehaarti and *M. bantamensis*), and *Spiroclypeus tidoenganensis*. These taxa represent foraminiferal Assemblage 1 and suggest an Early Miocene, late Aquitanian (middle Te5 letter stage, equivalent to sub-tropical zone M2, 21.12–19.3 Ma) to middle Burdigalian age (upper Te5 letter stage, equivalent to sub-tropical zone M3, 19.3–17.54 Ma) for the deposition of this limestone (Adams 1970; BouDagher-Fadel 2008). The microfossil assemblage, containing reef-loving organisms such as lepidocyclinid and miogypsinid foraminifera, and red algae that favor the seaward side of reef crests, together with the relatively large amount of early diagenetic marine spar and fragmentation, indicate deposition within a relatively high-energy reef-front environment (Hallock and Glenn 1986). The shallow-water platform facies was observed at three localities along Sungai Anggrisi (Figs. 5, 6).

Argillaceous platform facies

The planar bedded white limestone unit is paraconformably overlain by alternating irregular beds of grey marl and brown packstone in a succession approximately 20 m thick (Figs. 3b, 6). This heterolithic succession is equivalent to the Sekau Formation of the western Bird's Head (Fig. 2). The white limestone has a sharp planar contact with the overlying heterolithic succession parallel to bedding. Carbonate interbeds within the succession comprise brown, well-indurated, argillaceous mud-lean packstone and grainstone that contain abundant larger benthic and planktonic foraminifera, bivalve fragments, rhodophyte red algae, and ferroan cement (Fig. 4). *Trypanites* ichnofacies vertical burrows are observed along bed boundaries in contact with the overlying grey marl. Biostratigraphic analyses of the lower brown packstone interbeds contain a larger benthic foraminiferal assemblage including *Lepidocyclina* (*Nephrolepidina*) *brouweri*, *L. (N.) nephrolepidinoides*, *L. (N.) stratifera*, *L. (N.) verbeeki*, *Miogypsina cushmani*, *M. intermedia* and *Miogypsinoides indica*. This assemblage suggests an Early to Middle Miocene, Burdigalian to early Langhian (lower Tf1–upper Te5 letter stages, equivalent to sub-tropical zones

M3–M5, 19.3–15.1 Ma) age for these units (Adams 1970; BouDagher-Fadel 2008). The foraminiferal assemblage in these interbeds represent Assemblage 3 and indicate an inner neritic depositional setting. The argillaceous platform facies was observed at seven localities (Figs. 5, 6). Fossils identified from within this facies are interpreted to have been deposited in situ as bioclasts display no evidence for transportation (e.g., through abrasion, fragmentation, or as intraclasts). The presence of *Trypanites* ichnofacies borings also indicate that these carbonates were deposited in situ and the platform top may have formed hard grounds.

Periplatform facies

The uppermost packstone beds contain an increasing abundance of large, flat, rotaliina foraminifera, which include *Cycloclypeus indopacificus*, *Katacycloclypeus annulatus*, *Lepidocyclina* (*Nephrolepidina*) *ferreroi*, *L. (N.) inflata*, *L. (N.) martini*, *Miogypsina antillea*, *M. kotoi*, *M. regularia*, *Planostegina* spp. and *Operculina* spp. These foraminifera represent Assemblage 5 and suggest a Middle Miocene, middle Langhian to latest Serravallian (upper Tf1–Tf3 letter stages, equivalent to sub-tropical zones M6–10, 15.1–11.63 Ma) age (Adams 1970; BouDagher-Fadel 2008). The abundance of large, platy benthic foraminifera indicates deposition in relatively moderate water depths between 20 and 50 m in platform slope or periplatform environments (Hallock and Glenn 1986). The latest Serravallian age assigned to the uppermost brown packstone bed of the heterolithic succession is based on the last observed occurrences of *K. annulatus* and *C. indopacificus*, which are zonal markers for the end of Indo–Pacific letter stages Tf2 and Tf3, respectively (BouDagher-Fadel 2008). The Tf3/Tg boundary (11.63 Ma) of the Indo–Pacific Letter Stages is denoted by a significant faunal turnover in larger benthic foraminifera at the end of the Serravallian (Adams 1970; BouDagher-Fadel 2008). However, the faunal turnover at the Tf3/Tg boundary in New Guinea is denoted by complete facies change in planktonic foraminifera replacing shallow-water taxa, rather than a change in composition of shallow benthic assemblage. The uppermost periplatform facies was observed at three localities (Figs. 5, 6).

Pelagic facies

The argillaceous platform and periplatform facies are interbedded with grey marl within the heterolithic succession. Petrographic analysis of the marl shows them to be packstone comprising ~90% planktonic foraminifera and ~10% micrite matrix. Planktonic foraminifera within marl of the lower half of the heterolithic succession include *Catapsydrax dissimilis*, *D. altispira*, *D. baroemoensis*, *Globigerina*

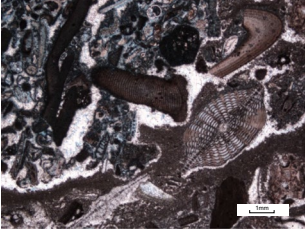
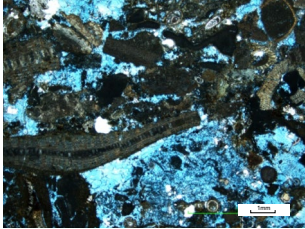
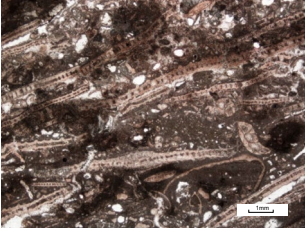
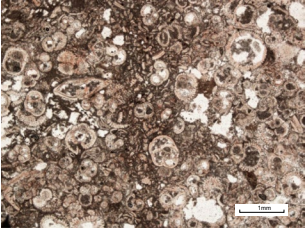
Shallow-water platform facies	Argillaceous platform facies	Periplatform facies	Pelagic facies
			
<p>Maruni Lst, Burdigalian (Zone M3)</p>	<p>Maruni Lst, Langhian (Zone M5)</p>	<p>Wafordori Fm, Langhian (Zone M7)</p>	<p>Napisendi Fm, Serravallian (Zone M10)</p>
<p>Dominant morphologies: Diverse assemblage, and robust rotaliid foraminifera</p>	<p>Dominant morphologies: Diverse assemblage, and robust rotaliid foraminifera</p>	<p>Dominant morphologies: Large, flat, rotaliid foraminifera</p>	<p>Dominant morphologies: Globular planktonic foraminifera</p>
<p>Key taxa in facies: <i>Lepidocyclina</i> <i>Eulepidina</i> <i>Miogypsina</i> <i>Miogypsinoidea</i> <i>Neorotalia</i> <i>Alveolinella</i> <i>Marginopora</i> <i>Miliolids</i></p>	<p>Key taxa in facies: <i>Lepidocyclina</i> <i>Eulepidina</i> <i>Miogypsina</i> <i>Miogypsinoidea</i> <i>Neorotalia</i> <i>Alveolinella</i> <i>Marginopora</i> <i>Miliolids</i></p>	<p>Key taxa in facies: <i>Heterostegina</i> <i>Planostegina</i> <i>Cyclocypeus</i> <i>Katacyclocypeus</i> <i>Operculina</i> <i>Spirocypeus</i> <i>Amphistegina</i></p>	<p>Key taxa in facies: <i>Globorotalia</i> <i>Truncorotalia</i> <i>Sphaeroidinellopsis</i> <i>Orbulina</i> <i>Globigerinoides</i> <i>Globigerina</i> <i>Candeina</i> <i>Pulleniatina</i> <i>Globoquadrina</i></p>
<p>Lithology: Benthic foraminiferal grainstones or mud-lean packstones, abundant non-ferroan calcite spar</p>	<p>Lithology: Mixed benthic/planktonic foraminiferal grainstones or mud-lean packstones, abundant ferroan calcite spar and clay material in matrix</p>	<p>Lithology: Mixed benthic/planktonic foraminiferal wackestones and packstones</p>	<p>Lithology: Planktonic foraminiferal wackestones and packstones</p>
<p>Water depth: Shallow</p>	<p>Water depth: Shallow</p>	<p>Water depth: Moderate</p>	<p>Water depth: Deep</p>

Fig. 4 Samples analyzed for petrography were grouped into four facies based on microfacies and faunal composition. 1 Shallow-water platform facies contain a diverse assemblage of organisms including robust rotaliid foraminifera often within grainstone or mud-lean packstone. 2 Argillaceous platform facies contain similar faunal composition to the shallow-water platform facies with the inclusion

of significant ferroan calcite cement. 3 Periplatform facies comprise packstone containing large, flat, rotaliid foraminifera. 4 Pelagic facies comprise packstone and wackestone that contain abundant planktonic foraminifera. Microfossil paleoenvironmental and bathymetric interpretations were made using Bé and Tolderlund (1971), Bé (1977), Hallock and Glenn (1986), and BouDagher-Fadel (2008)

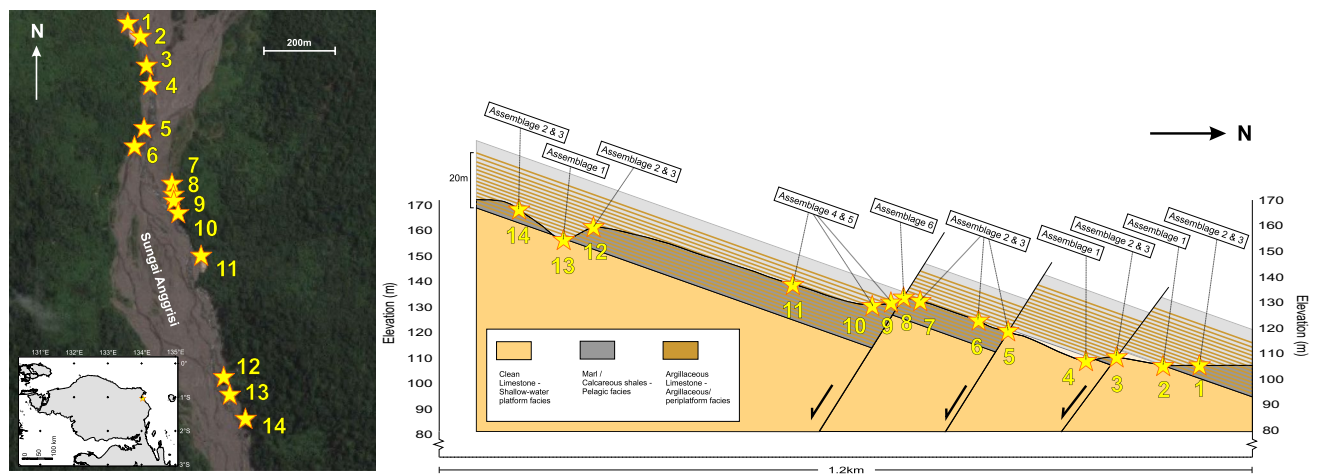


Fig. 5 Stratigraphic and structural cross section depicting the locations, foraminiferal assemblages, and facies encountered at 14 locations along the drowning succession observed at Sungai Anggrisi,

eastern Bird's Head. Satellite image credit: Google Earth (Map data: Google, DigitalGlobe)

occlusa, *Ga. praebulloides*, *Globigerinoides altiapertura*, *Gdes. obliquus*, *Gdes. quadrilobatus*, *Gdes. subquadratus*, *Gdes. trilobus*, *G. dehiscens*, *G. woodi* and *Praeorbulina glomerosa*. These taxa represent foraminiferal Assemblage 2, which suggest an Early Miocene, Burdigalian (sub-tropical zone M3, 19.3–17.54 Ma) to Middle Miocene, early Langhian (sub-tropical zone M5, 16.38–15.1 Ma) age for this unit (Iaccarino et al. 2007; BouDagher-Fadel 2015). Globular planktonic foraminifera are common in the basal grey marl of the succession and are indicative of ‘moderate’ water depths between 50 and 100 m (Bé and Tolderlund 1971; Bé 1977). The large number of globular planktonic foraminifera, with a planktonic:benthic ratio greater than 100:1, and micrite mud in this sample, indicate it was deposited as a pelagic limestone in a middle neritic environment. The lower pelagic facies was observed at seven localities (Figs. 5, 6).

The upper part of the heterolithic succession contains different assemblages of interbedded benthic and planktonic foraminifera. The upper grey marl contains the planktonic foraminifera *D. altispira*, *D. baroemoensis*, *D. globosa*, *D. quadrata*, *D. venezuelana*, *Globigerina decoraperta*, *Ga. eamesi*, *Ga. occlusa*, *Ga. praebulloides*, *Globigerinella prae-siphonifera*, *Globigerinoides quadrilobatus*, *Gdes. subquadratus*, *Gdes. trilobus*, *G. dehiscens*, *Globorotalia conoidea*, *Grot. menardii*, *Grot. miozea*, *Grot. peripheroronda*, *Grot. praemenardii*, *Grot. scitula*, *Globoturborotalita woodi*, *Orbulina bilobata*, *O. suturalis*, *O. universa* and *Sphaeroidinellopsis subdehiscens*. These taxa represent foraminiferal Assemblage 4 and suggest a Middle Miocene, middle Langhian (sub-tropical zone M6, 15.1–14.24 Ma) to latest Serravallian (sub-tropical zone M10, 11.79–11.63 Ma) age for this marl (Iaccarino et al. 2007; BouDagher-Fadel 2015). An increasing abundance of carinate planktonic foraminifera is observed up-section and indicates increasing water depths in excess of 100 m (Bé and Tolderlund 1971; Bé 1977). The upper pelagic facies was observed at three localities (Figs. 5, 6).

The heterolithic succession above the Maruni Limestone passes upwards to Befoor Formation siliciclastic strata (Figs. 2, 6). The Befoor Formation is dated to range from the Late Miocene, Tortonian, to Pliocene, Zanclean (sub-tropical zones M11–PL3, 11.63–3.59 Ma). However, along Sungai Anggrisi, only the lower Befoor Formation is observed, representing the Tortonian to Messinian (sub-tropical zones M11–M14, 11.63–5.72 Ma) based on the presence of the planktonic foraminifera *G. dehiscens*, *Globorotalia menardii*, *Pulleniatina primalis*, *Globigerinoides quadrilobatus*, *Gdes. trilobus*, *N. acostaensis*, and *S. subdehiscens* (Iaccarino et al. 2007; BouDagher-Fadel 2015). These taxa represent Assemblage 6 and are observed at a single locality along Sungai Anggrisi (Figs. 5, 6). This assemblage is recorded within a continuous succession of pelagic facies strata and an absence of interbeds of shallower-water carbonates. The Befoor Formation is observed as an uninterrupted succession of wholly

pelagic facies strata several tens of meters thick where it is widespread northwest of Manokwari, until carbonate deposition resumes once more in the Zanclean (Fig. 2).

Interpretation of the Sungai Anggrisi drowning succession

Heterolithic beds overlying the white reefal limestone along Sungai Anggrisi are interpreted here as a drowning succession, formed during approximately 7.5 My of Burdigalian–Serravallian carbonate platform drowning. Heterolithic successions were interpreted as drowning strata based on sedimentological and stratigraphic evidence marked by significant facies and faunal changes (e.g., Masse and Fenerci-Masse 2011). The initiation of platform drowning is interpreted to have occurred during sub-tropical planktonic foraminiferal zone M3 (19.3–17.54 Ma) due to the transition from shallow-water platform to argillaceous platform and pelagic facies likely occurring some time within this zone (Fig. 6). The faunal turnover observed between the periplatform and pelagic facies at the Tf3/Tg boundary (11.63 Ma) is interpreted to mark the age of completion of carbonate platform drowning as it represents the last bed of periplatform facies strata underlying wholly deep-water strata of uninterrupted pelagic facies (Fig. 6). The drowning unconformity therefore is situated above the uppermost bed containing Assemblage 5 taxa (Fig. 6) and is interpreted to correlate to the unconformity described by McAdoo and Haebig (1999) within the North Irian Basin, the ‘U’ Marker within the Salawati and Bintuni basins, and uppermost limestone unit of the Sekau Formation in the western Bird’s Head. In the eastern Bird’s Head, only c. 20 m of the heterolithic succession is exposed, consistent with descriptions of the 30–50-m-thick Sekau Member of the western Bird’s Head (Visser and Hermes 1962; Pieters et al. 1983). However, this is considerably less than the c. 200-m ‘U’ Interval found in the southern Salawati Basin.

The Early Miocene units of the NGLG, including the Maruni and Kais Limestones, are interpreted to represent typical carbonate platform deposits (e.g., Hallock and Schlager 1986). The Maruni Limestone is interpreted to have been deposited as part of a rimmed platform containing proximal back-reefs, passing basinward to fringing reefs and periplatform slope facies. The basal white Maruni Limestone of the Sungai Anggrisi succession represents healthy platform growth, deposited with high average surface water temperatures, high water transparency, and low levels of available nutrients (e.g., Föllmi et al. 1994). Low nutrient levels (oligotrophy) are indicated by the presence of large benthic foraminifera and absence of silt and coarser siliciclastic grains, indicating low levels of terrestrial runoff and therefore nutrient supply (Hallock 1988; Weissert 1989, 1990; Föllmi et al. 1994).

Throughout the Miocene, carbonate production rates were reduced in the Bird’s Head as denudation of land areas

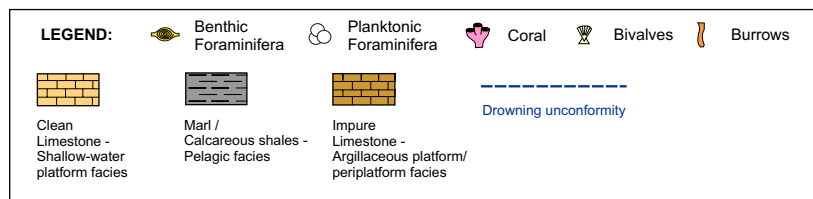
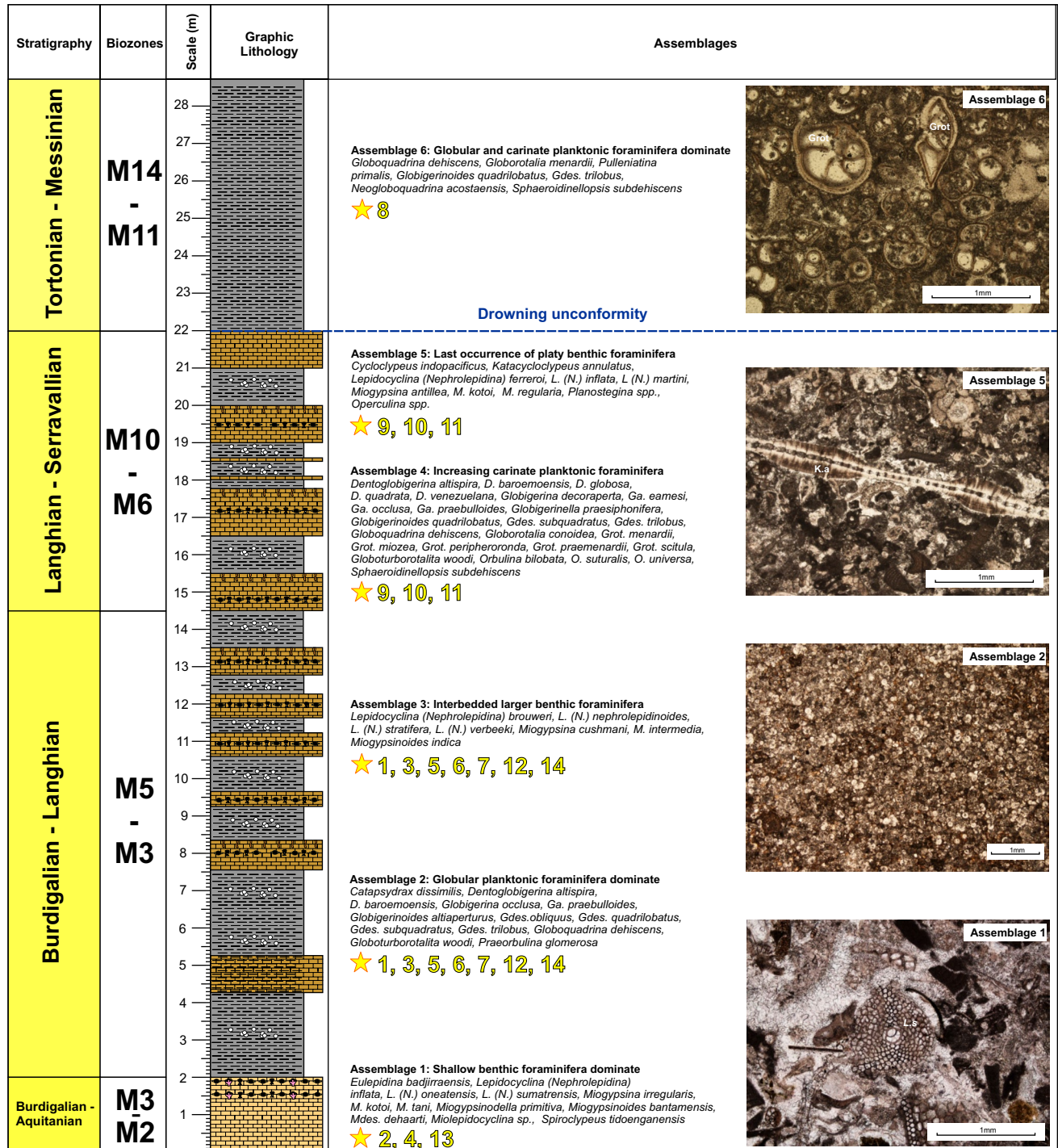


Fig. 6 Composite sedimentary log based on field data collected at 14 localities visited at the drowning succession exposed along Sungai Anggrisi. The drowning succession deepens upwards with increasing numbers of globular and carinate planktonic foraminifera replacing typical reefal shallow benthic foraminifera and moderate-water-depth larger platy foraminifera. The uppermost periplatform facies bed contains specimens of *Katacycloclypeus annulatus* (K.a) in the microfossil assemblage, indicating a Serravallian age to the completion of the drowning process along Sungai Anggrisi. Modified from Gold et al. (2014)

increased the amount of terrigenous material shed into off-shore areas (e.g., Wilson 2002). It is therefore interpreted that minor siliciclastic grains and iron within Fe-rich calcite cements in the argillaceous platform facies was sourced by terrestrial runoff. Iron was transported as particles of iron oxides from surrounding muddy, clay-rich sediment or adsorbed on to clays. The presence of Fe-enrichments in drowning and post-drowning sediments are often related to a primary higher nutrient content in the water column and a change from oligotrophic to meso- or eutrophic conditions (Zempolich 1993; Föllmi et al. 1994; Drzewiecki and Simo 1997; Blomeier and Reijmer 1999).

Carbonate accumulation rates are inversely related to nutrient availability at levels typically found in modern, low-latitude oceans (Hallock and Schlager 1986). One such nutrient involved in the primary productivity of oceans is iron. The link between increased iron and increased phytoplankton activity in the oceans is well established (Martin and Fitzwater 1988; Walsh and Steidinger 2001; Sellner et al. 2003; Winckler et al. 2008; Costa et al. 2016). The positive relationship between increasing amounts of iron and high phytoplankton growth rates occurs due to the role iron plays in photosynthesis, where iron–sulfur proteins, termed ‘ferredoxins’, mediate electron transfer in the metabolic reactions of photosynthesis (Geider and La Roche 1994). The increased amount of iron, and therefore phytoplankton, acted to reduce levels of transparency in the water column above the Maruni Limestone platform as conditions changed from that of oligotrophy to mesotrophy. Hallock and Schlager (1986) suggested that a reef cannot recover if submerged below 25% of surface light levels. This equates to only several meters or tens of meters of submergence required to drown present-day reefs in Aqaba, the Florida Keys, Hawaii, and Barbados (Hallock and Schlager 1986). The increasing abundance of carinate planktonic foraminifera replacing shallow-water benthos observed within drowning successions of the Bird’s Head indicate an increase in water depths to at least 100 m. This is more than sufficient to terminate carbonate production in waters where transparency is low. Burgess and Hovius (1998) demonstrated that where rates of fluvial discharge are high, such as areas of intense tropical denudation, many modern deltas are

able to prograde during marine transgressions and times of high relative sea-level. This scenario explains how iron was introduced into the depositional environment through terrestrial run-off while water depths began to increase above the Maruni Limestone platform.

Environmental preferences of organisms observed within foraminiferal Assemblage 1 suggest that water depths locally were no greater than 20 m during deposition of the reefal Maruni Limestone at the initiation of the drowning process during the middle Burdigalian. The increasing abundance of carinate planktonic foraminifera up-section indicates that water depths increased to greater than 100 m (e.g., Bé and Tolderlund 1971; Bé 1977) by the Tortonian, a duration of approximately 7.5 My. An increase in water depth of at least 75 m in approximately 7.5 My recorded from the entire Sungai Anggrisi succession equates to a rate of regional relative sea-level rise of approximately 10 m/My. The lowest accumulation rates for prograding carbonates are reported as 30 m/My (Schlager 1981; Handford and Loucks 1993; Wissler et al. 2003), therefore a long-term rise in relative sea-level of 10 m/My should not have been sufficient to outpace the rate of carbonate accumulation. This observation supports the ‘drowning paradox’ of Schlager (1981). However, Kim et al. (2012) demonstrated that the drowning paradox can be overcome if accumulation rates at the initiation of drowning are lower than the rate of relative sea-level rise. Evidence for mesotrophy in the carbonate production factory observed in the Sungai Anggrisi drowning succession may be one reason for initial carbonate accumulation rates to be low at this location. This may have been further exaggerated by the slow response time of benthic organisms of the periplatform and argillaceous platform facies to re-establish as relative sea-level rise continued to progress leading to drowning of the Maruni Limestone platform. The interpretation therefore is that a combination of relative sea-level rise and environmental deterioration is responsible for platform drowning in the Bird’s Head.

The already-foundering Maruni Limestone platform top under elevated trophic conditions may have finally succumbed to drowning during shorter, more rapid pulses of relative sea-level rise (e.g., Hallock and Schlager 1986). Incursions of pelagic facies over argillaceous platform and periplatform facies limestones are interpreted here as catch-up and short-term keep-up deposition (as defined in Kim et al. 2012) when drowning during longer-term relative sea-level rise was interrupted by high-frequency sea-level falls of approximately 20–30 m during the middle Miocene (Miller et al. 2005; Snedden and Liu 2010; Fig. 2). The amplitude of these oscillations was enough to shift the deposition of periplatform and pelagic facies up and down the platform slope, depositing the relatively shallow-water facies during intervening troughs in the short-term global sea-level trend (Fig. 2). The sharp increase in

Fig. 7 Gross depositional environment maps depicting position of migrating belts of shallow-water carbonate platform, argillaceous/periplatform, and pelagic facies in the Early, Middle and Late Miocene. Deeper-water facies encroach causing shallower-water areas to backstep to the north through time. Data points correspond to outcrop samples only

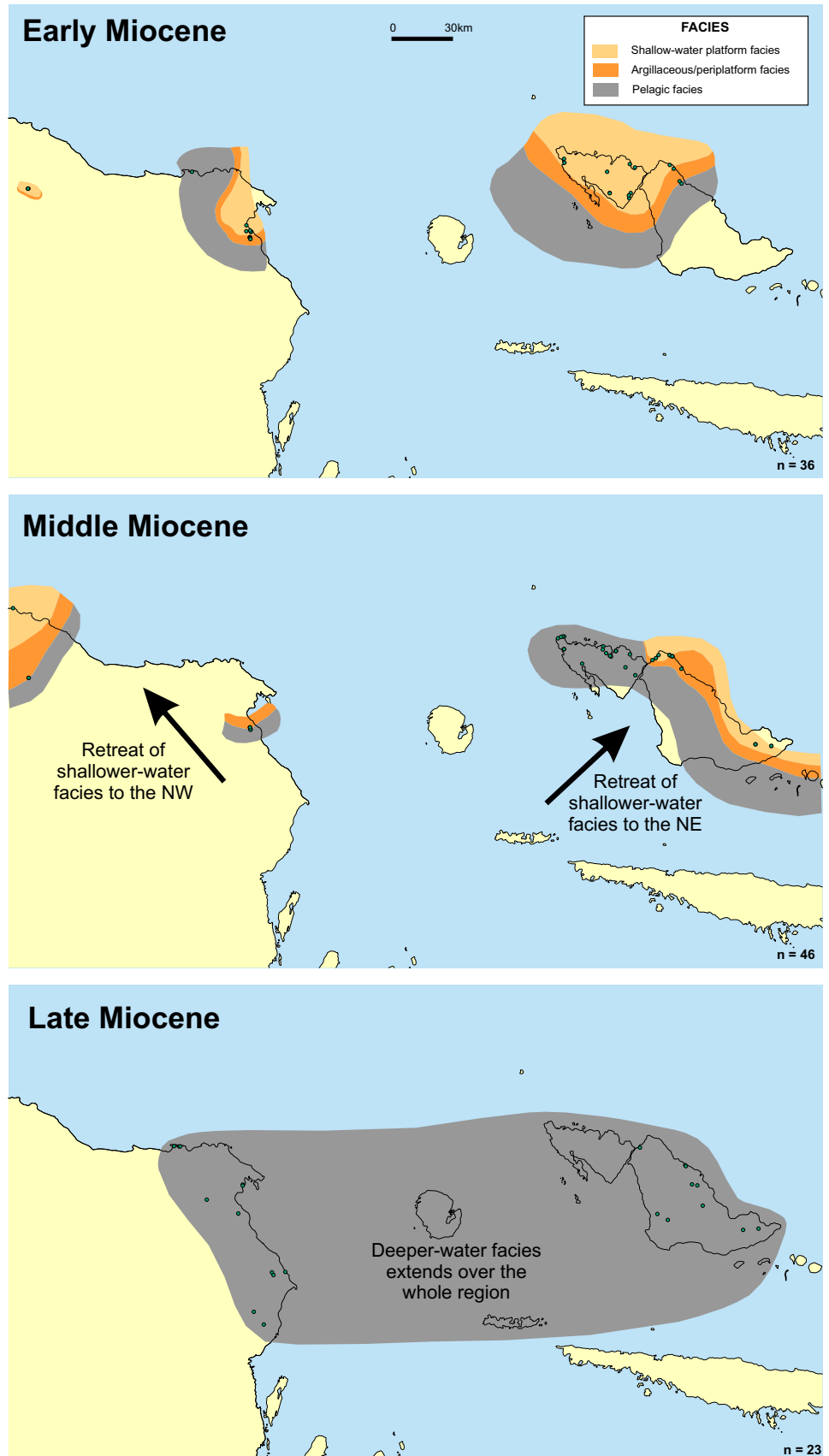


Table 1 References for public domain wells used in this study

Well	References
Agung-1	Fraser et al. (1993), Pairault et al. (2003)
ASA-1X	Darman and Sidi (2000)
Gunung-1	Fraser et al. (1993), Pairault et al. (2003)
Klalin-1	Livingstone et al. (1992)
Monie south-1	Pigram et al. (1982)
Roabiba-1	Fraser et al. (1993)
Sago-1	Chevallier and Bordenave (1986)
Tarof-2	Fraser et al. (1993)
TBA-1X	Livingstone et al. (1992)
TBA-2X	Livingstone et al. (1992)
TBE-1X	Koesoemadinata (1976), Chevallier and Bordenave (1986), Dolan and Hermany (1988), Fraser et al. (1993)
TBH-1X	Livingstone et al. (1992)
TBM-1X	Livingstone et al. (1992)
TBO-1X	Livingstone et al. (1992)
West Salawati	Satyana (2003)

the short-term global sea-level trend from approximately 20 m above present-day sea-level at the end of the Serravallian, to approximately 80 m above present-day sea-level at the Tor1 MFS, suggests a sudden sea-level rise of 60 m in 1 My (Fig. 2). This drastic increase in the rate of short-term sea-level rise may have outpaced rates of carbonate production on the Maruni Limestone platform, resulting in the termination of keep-up deposition and formation of the drowning unconformity by the end of the Middle Miocene. By deposition of the Befoor Formation in the Tortonian, water depths were high enough that carbonate deposition on the platform was unable to re-establish and it was draped by the deep-water sediments that overlie the drowning unconformity.

The heterolithic succession between the Maruni Limestone and Befoor Formation may represent an easterly lateral equivalent to the Sekau Formation (Fig. 2). Several scenarios have been suggested for the deposition of the Sekau Formation. These include struggling reef growth in mud-polluted waters and limestone interbeds within the heterolithic succession representing fore-reef deposits encroached by relatively deeper-water marls (Visser and Hermes 1962; Pieters et al. 1983). Along Sungai Anggrisi both scenarios are shown to be true. Deposition in mud-polluted waters is indicated by the argillaceous platform facies, and encroachment of deeper-water sediments over a shallow-water platform is indicated by the periplatform and pelagic facies.

Through a reduction in carbonate accumulation rates by excess nutrients, outpaced by the rate of relative sea-level rise, transgression of pelagic strata occurred across the top

of the Maruni Limestone platform. This is clearly shown in the gross depositional environment maps of Miocene age samples and outcrops from the eastern Bird's Head and islands of Biak and Supiori (Fig. 7). These delineate the shift in relative position of the migrating facies belts throughout the Miocene. The maps show the retreat of the shallow-water facies belts towards the north from the beginning to end of the Miocene (Fig. 7). The shallow-water carbonate platform facies (Fig. 4) is widespread across much of the Bird's Head during the Early Miocene (Fig. 7). During the Middle to Late Miocene, a faunal turnover of shallow-water benthic foraminiferal assemblages to deeper-water planktonic foraminiferal assemblages is observed as shallow-water facies are replaced by periplatform, platform, and pelagic facies (Fig. 4) across the study area (Fig. 7).

Evidence for drowning in New Guinea from well data

Further evidence for carbonate platform drowning was identified during review of biostratigraphic and wireline-log data from 15 public domain wells (Table 1). These wells are located within the known hydrocarbon producing Salawati and Bintuni Basins, and Arafura Sea (Fig. 1). The closest well to the outcrop analogues described in this paper is approximately 200 km to the west. However, the Kais Limestone and overlying Sekau, Klasafet and Klasaman formations are widely distributed in outcrop between the data points of this study, implying that carbonate platform drowning occurred over a large area during the Middle to Late Miocene.

Biostratigraphic data from 15 wells in western New Guinea record a turnover in fauna and facies of open-marine assemblages of planktonic foraminifera replacing assemblages of shallower-water benthic foraminifera (Fig. 8). The termination of stable carbonate platform growth of the NGLG is recorded in wells Agung-1, Klalin-1, Sago-1, and TBE-1X to occur during the early Langhian (upper Te5 letter stage, equivalent to sub-tropical zone M5, 16.38–15.1 Ma; Fig. 8). The shutdown of stable carbonate platform growth is often denoted by the last occurrences, base to top, of 'reefal' organisms such as corals and sponges, and typical shallow-water benthic foraminifera including lepidocyclinids, miogypsinids, and miliolids (Fig. 8). Several wells, Monie South-1, Roabiba-1, and Sago-1, record observable heterolithic successions of planktonic foraminiferal calcareous shales and marls interbedded with shallow-water benthic foraminiferal limestones. These successions were deposited between the middle Langhian and latest Serravallian (M6–M10, 15.1–11.63 Ma), marking the inception and completion of the drowning process (Fig. 8). The completion of

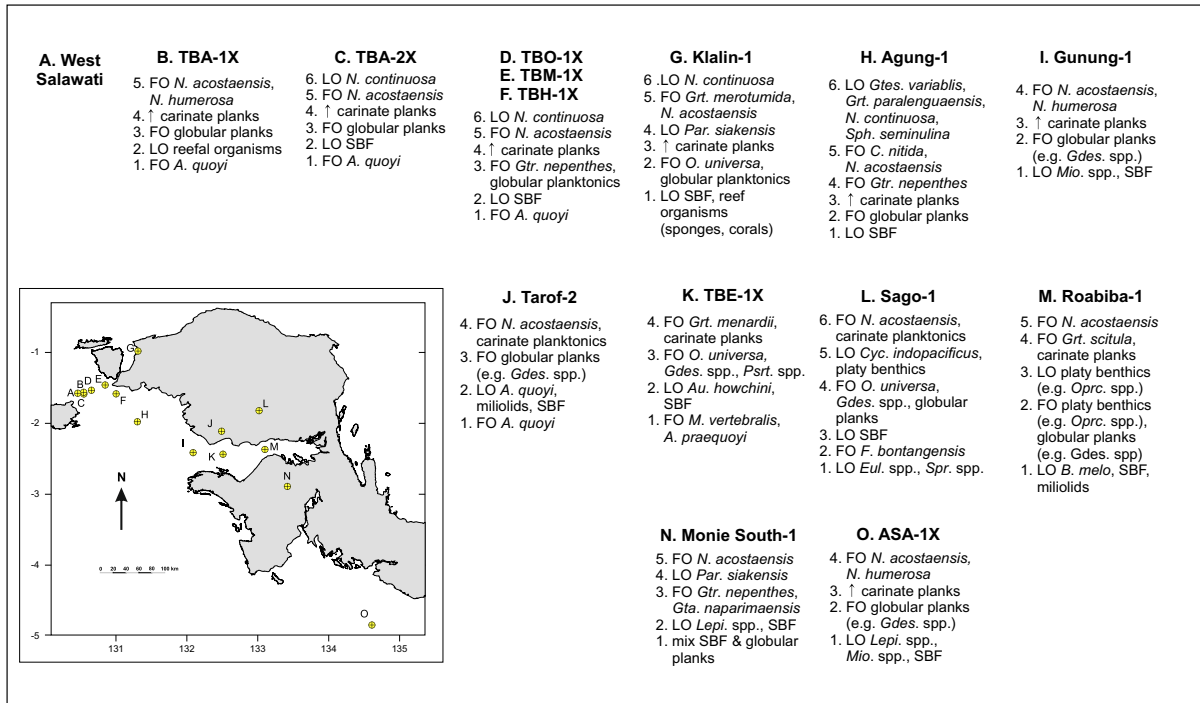
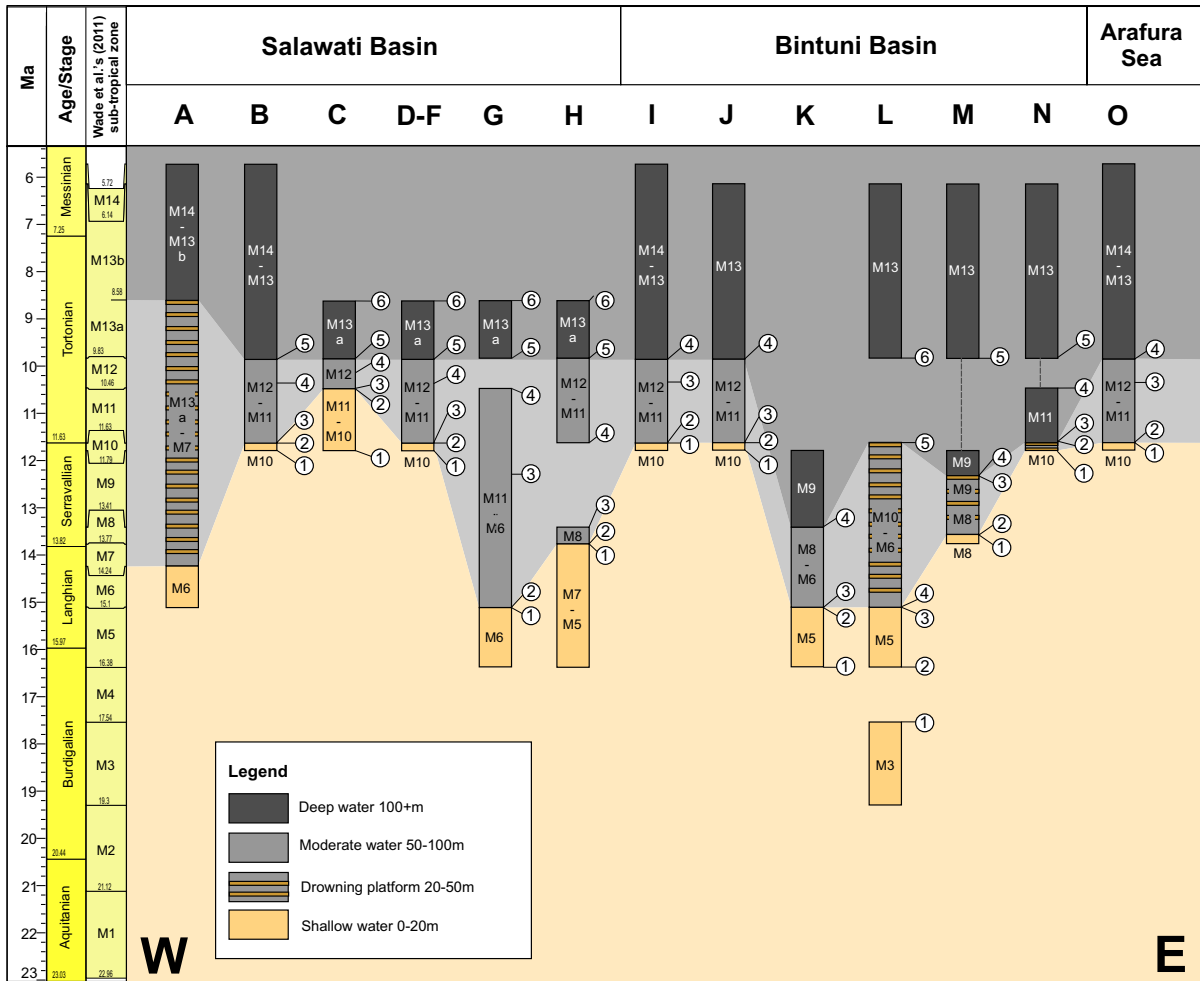


Fig. 8 Faunal turnover of shallow-water foraminifera replaced by increasingly deeper-water forms as recorded in 15 wells. Planktonic foraminiferal biozones based on Wade et al. (2011) zonation scheme. Zonal boundaries marked by changes in age or environmental diagnostic taxa. There is a general trend for stable platform deposition to cease by M5 (early Langhian), with the initiation of drowning occurring at M6 (mid Langhian), with the completion of the drowning process by M10 (latest Serravallian) marked by the disappearance of shallow-water or platy benthic foraminifera. Fully deep water in excess of 100 m, dominated by carinate planktonic foraminifera occurs by M13b (mid Tortonian). *FO* first occurrence, *LO* last occurrence, *SBF* shallow benthic foraminifera. *A. Alveolinella*, *Au. Austrorotulina*, *B. Borelis*, *C. Candaina*, *Cyc. Cycloclypeus*, *Eul. Eulepidina*, *F. Flosculinella*, *Ga. Globigerina*, *Gdes. Globigerinoides*, *Grt. Globorotalia*, *Gta. Globigerinita*, *Gtr. Globoturborotalia*, *Gtes. Globorotaloides*, *Lepi. Lepidocyclina*, *M. Marginopora*, *Mio. Miogypsina*, *N. Neogloboquadrina*, *O. Orbulina*, *Oprc. Operculina*, *Par. Paragloborotalia*, *Psrt. Pseudorotalia*, *Sph. Sphaeroidinellopsis*, *Spr. Spiroclypeus*)

the drowning process during the Serravallian is denoted by the last up-hole occurrences of moderate-water-depth larger benthic foraminifera, such as *Cycloclypeus* spp. and *Operculina* spp. (Fig. 8). The mixed assemblages of planktonic and shallow-water benthic foraminifera in these wells are interpreted to be analogous to the heterolithic beds observed above the Maruni Limestone in the Sungai Anggrisi drowning succession, and other locations observed in the eastern Bird's Head.

Other wells, ASA-1X, Gunung-1, Kalitami-1, Tarof-2, TBA-2X, TBH-1X, TBM-1X, and TBO-1X, record a sudden, rather than gradual, change in facies at the end of the Serravallian (M10, 11.79–11.63 Ma; Fig. 8). This is denoted by the presence of *Alveolinella quoyi* and last up-hole occurrences of other shallow-water foraminifera such as lepidocyclinids and miogypsinids, immediately overlain by facies with wholly planktonic foraminiferal assemblages (Fig. 8).

In many wells, basal planktonic foraminiferal assemblages are dominated by 'moderate-water' taxa (e.g., Bé and Tolderlund 1971; Bé 1977) becoming increasingly dominated by carinate and other 'deep-water' taxa such as *Globorotalia* spp., *Neogloboquadrina* spp. and *Sphaeroidinellopsis* spp. up-hole (e.g., Bé and Tolderlund 1971; Bé 1977). Fully deep-water conditions, with water depths in excess of 100 m, are interpreted to have occurred by the middle Tortonian (M13b, 8.58–6.14 Ma), often denoted by *Globorotalia merotumida*, *Neogloboquadrina acostaensis*, *N. continuosa*, and *N. humerosa* (Fig. 8).

In the Salawati Basin, pinnacle reefs at the top of the Kais Platform are younger towards the west of the basin than in the east (Fig. 8). The uppermost strata of the 'U' Interval are described to be N16 age in the west (recalibrated to Wade et al. (2011) sub-tropical zone M13a, 9.83–8.58 Ma), and N9–N16 (recalibrated to Wade et al. (2011) sub-tropical zones M6–M13a, 15.1–8.58 Ma), in the east (Satyana 2003),

marking the time at which platform drowning in the Salawati Basin was complete.

Data from outcrop collected by this study indicates that the drowning process was complete by the end of sub-tropical planktonic foraminiferal zone M10, and the Tf3/Tg boundary (11.63 Ma) of the Indo–Pacific Letter Stages, in areas farther to the east of the Bird's Head (e.g., Fig. 2, Columns C–F). However, in the Salawati Basin to the west, the drowning process is interpreted to be complete by the end of sub-tropical zone M13a (8.58 Ma). These observations suggest that the drowning unconformity formed diachronously, younging to the west from M10 (11.63 Ma) to M13a (8.58 Ma) times, over a period of approximately 3 Ma (Figs. 2, 8). This substantiates outcrop evidence from the mainland Bird's Head where migrating facies belts indicate the back-stepping of shallow-water facies to the northwest during the Early to Late Miocene (Fig. 7), further supporting diachronous formation of the drowning unconformity as younging towards the west. In the eastern Bintuni Basin, evidence from the Aroba, Jamusura and Besiri River wells (Rossetter 1976) indicates that termination of Kais platform growth occurred in the Middle Miocene (Pigram et al. 1982). However, this diachroneity is also shown in the western Bintuni Basin where the Kais Formation is indicated as continuing into the Late Miocene (Koesoemadinata 1976; Collins 1977; Chevallier and Bordenave 1986; Dolan and Hermany 1988). In the Bintuni Basin, diachroneity is interpreted to occur as subsidence migrated westwards away from the developing LFTB (Collins 1977).

Mechanisms for northern New Guinea regional relative sea-level rise

Gibson-Robinson and Soedirdja (1986) attributed the cause of the termination of Kais Platform and pinnacle reef growth in the Salawati Basin to a minor regression or still-stand that prevented further growth, leading to erosion. During subsequent renewed transgression, platform reef growth was not re-established (Gibson-Robinson and Soedirdja 1986). However, we consider that favorable conditions during renewed transgression following a minor regression or still-stand should have enabled the platform to re-colonize and continue to grow. Empirical evidence for platform drowning along Sungai Anggrisi is derived from facies and micropaleontological analyses of logged outcrop sections. This evidence indicates that a reduction in carbonate accumulation rates through environmental deterioration, outpaced by the rate of relative sea-level rise, is responsible for the formation of similar drowning successions observed in wells from the western Bird's Head, comparable to the 'U' Interval of the Salawati Basin.

The mechanism for the increase in rate of relative sea-level rise is unknown. However, rotation of the Philippine Sea Plate, initiating at about 20 Ma, and minor subduction close to the northern New Guinea margin related to extrusion of 11–12 Ma volcanic rocks in the Bird's Head (Hall and Spakman 2015), could have driven an acceleration in subsidence in extensional domains. This may have led to an increase in rate of relative sea-level rise contributing to platform drowning. In the Salawati Basin, Late Miocene termination of the Kais platform is attributed to tectonic subsidence commencing at about 6 Ma, with an estimated subsidence rate of approximately of 1 km per million years in the area of the Waipili-1 well (Livingstone 1992).

Extension is interpreted to have occurred across the Bird's Head during most of the Miocene, initiating in the Burdigalian and continuing into the Tortonian. This is supported by biostratigraphic data collected and reviewed by this study, and the presence of sequences, several km thick, of middle and late Miocene siliciclastic strata within the regional basins, including the Salawati Basin, described by Visser and Hermes (1962). However, the entire region has undergone subsequent compressional deformation, which overprints the original shallow-water sequences that are now well above sea-level. The timing of the onset of compressional deformation is controversial; however, we believe that it initiated with the onset of uplift of the LFTB of the Bird's Neck, and mountains of the Central Ranges (Fig. 1) across the Bird's Body during the latest Tortonian–Messinian (Hill and Gleadow 1989; Cloos et al. 2005; Spakman and Hall 2010; Bailly et al. 2011; Hall 2012), after the drowning of the NGLG carbonate platform.

Conclusions

Although outcrops studied in the eastern Bird's Head are several hundred kilometers away from wells studied in western New Guinea, the similar facies, paleobathymetry, and age of faunal turnover in both regions indicate that carbonate platform drowning was widespread across the entire Bird's Head region. The identification of drowning successions farther to the east of the Salawati and Bintuni Basins, and Sekau Formation type locality, is significant, as it implies the New Guinea drowning event was a widespread environmental crisis during the Middle and Late Miocene.

The interbedded deep-water marl and argillaceous carbonates described from the 'U interval' and Sekau Formation are similar to alternating beds of the same lithologies observed in Sungai Anggrisi. This is significant, as drowning successions such as the Sekau Formation, which contains commercial oil in the western Bird's Head (Visser and Hermes 1962), may extend to the east. Heterolithic successions of the 'U' Interval in the Salawati Basin record

the attempt of carbonate platform strata to keep-up with an increase in the rate of middle Miocene relative sea-level rise. Although the review of well data did not permit identification of Fe-enriched intervals, and thus interpretation of environmental deterioration through nutrification, the failure of platform accumulation to keep pace with the rate of relative sea-level rise is recorded by the formation of a drowning unconformity interpreted to be marked by the 'U' Marker.

The drowning process is interpreted to have lasted c. 9.5 My, initiating at approximately 18.0 Ma during subtropical planktonic foraminiferal zone M3 (19.3–17.54 Ma). Drowning unconformity formation occurred diachronously, younging towards the north and west over a period of c. 3 My. In areas to the east of the Bird's Head, the drowning unconformity is interpreted to have formed at 11.63 Ma; however, this surface is dated at 8.58 Ma in areas farther west.

A combination of a reduction in carbonate accumulation rates, through a change from oligo- to meso- trophic conditions (e.g., Brandano et al. 2016), and submergence through relative sea-level rise attributed to extensional tectonics, is interpreted to be responsible for the demise of the New Guinea Limestone platform (e.g., Hallock and Schlager 1986; Sattler et al. 2009). Evidence for submergence is recorded as a gradual increase in bathymetry from environmental preferences of foraminifera, either through short-term high-amplitude oscillations in the global glacio-eustatic sea-level trend or localized tectonic subsidence. Although regional tectonics and environmental deterioration are interpreted as a major control on the rise of relative sea-level in the Bird's Head, biostratigraphic dating of the exposed drowning successions in the study area permit a general correlation with recognized Miocene global cycles of sea-level change (e.g., Miller et al. 2005; Snedden and Liu 2010). These sequences correlate with a long-term trend of global sea-level rise from the late Oligocene, with several high-amplitude oscillations throughout the early to middle Miocene, reaching a peak of ca. 80 m above present sea-level in the early to middle Tortonian at the Tor1 maximum flooding surface (Snedden and Liu 2010). However, caution is required in attempting to correlate global sea-level trends with stratigraphy in Southeast Asia. This region has been the site of major tectonic events throughout the Neogene, resulting in localized areas of subsidence, uplift, and therefore relative sea-level changes. Nevertheless, onset of glacio-eustasy in the Miocene could well mean that a eustatic signal began to act as a significant control on Miocene strata, even in this tectonically active area.

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